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Evaluating the Effects of Human Activity over the Last Decades on the Soil Organic Carbon Pool Using Satellite Imagery and GIS Techniques in the Nile Delta Area, Egypt

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Received: 14 February 2019; Accepted: 6 May 2019; Published: 8 May 2019



Abstract: The study aims to clarify the relationship between soil organic carbon (SOC) and human activity under arid conditions, in the east area of the Nile Delta, Egypt. SOC is one of the critical factors in food production and plays an important role in the climate change because it affects the physio-chemical soil characteristics, plant growth, and contributes to sustainable development on global levels. For the purpose of our investigations, 120 soil samples (0–30 cm) were collected throughout different land uses and soil types of the study area. Multiple linear regressions (MLR) were used to investigate the spatiotemporal relationship of SOC, soil characteristics, and environmental factors. Remote sensing data acquired from Landsat 5 TM in July 1995 and operational land imager (OLI) in July 2018 were used to model SOC pool. The results revealed significant variations of soil organic carbon pool (SOCP) among different soil textures and land-uses. Soil with high clay content revealed an increase in the percentage of soil organic carbon, and had mean SOCP of 6.08 ± 1.91 Mg C ha⁻¹, followed by clay loams and loamy soils. The higher values of SOCP were observed in the northern regions of the study area. The phenomenon is associated with the expansion of the human activity of initiating fish ponds that reflected higher values of SOC that were related to the organic additions used as nutrients for fish. Nevertheless, the SOC values decreased in southeast of the study area with the decrease of soil moisture contents and the increase in the heavy texture profiles. As a whole, our findings pointed out that the human factor has had a significant impact on the variation of soil organic carbon values in the Eastern Nile Delta from 1995 to 2018. As land use changes from agricultural activity to fish ponds, the SOCP significantly increased. The agriculture land-use revealed higher SOCP with 60.77 Mg C ha⁻¹ in clay soils followed by fish ponds with 53.43 Mg C ha⁻¹. The results also showed a decrease in SOCP values due to an increasing in land surface temperature (LST) thus highlighting that influence of temperature and ambient soil conditions linked to land-use changes have a marked impact on surface SOCP and C sequestration.

Keywords: Spatiotemporal Evaluation; Remote sensing indices; land use/land cover; GIS

1. Introduction

Soil contains the largest stock of organic carbon compared to vegetation and the atmosphere. It is considered a main factor among environment elements that affects C reservoir of carbon stock totaling



2500 Gt (1 Gt $\frac{1}{4}$ 109 t), of which about 20% are stored in the top soil layers [1]. Soil organic carbon contents and storage are a global issue and are significantly linked to global warming. About 10% of the Earth's land surface is characterized by the presence of agro-ecosystems, and, consequently, top soil organic carbon storage is directly influenced by human action [2].

Soil organic carbon is associated to soil characteristics, climate, topographic factors, and crop management. The emission of CO_2 is not only originated from combustion, but also from soil organic carbon release due to a decreasing in soil organic sequestration equal to 1047.5 Gg CO_2 [3]. The development of soils as an organic carbon reservoir is controlled by different soil characteristics, such as, clay content, moisture level, aeration, soil depth, slope degree, and local aspects [4–6]. These factors affect the soil biological activities and, in turn, the transformation of organic matter. The soil organic carbon pool (SOCP) varies among different soils: Clay soils in general contain relatively high mean SOCP with 4.08 ± 1.41 kg C m², while highest SOCP with 7.07 kg C m² was observed in a clay loam soil associated with soil type and soil water retention [7]. The lowest SOCP of 2.57 kg C m² was observed in sandy clay loam soil that was associated with bare soil.

In arid and semi-arid regions, low precipitation with strong seasonality is one of the most important controlling factors of the decay and transformation of soil organic matter. In these arid conditions, to increase agricultural production (according to the economic and social situations of the settlers), irrigation and mineral additive practices are commonly used, producing as a consequence a reduction of carbon stocks in the soil layers [8]. Therefore, in arid to semi-arid areas, agriculture has a significant influence on the stock of soil organic carbon at the local level. However, the annual rate of loss of organic matter can vary greatly, depending on the specific cultivation practices and the type of plant/crop cover, whereupon special focus has to be put on soil melioration measures such as irrigation and drainage [9]. To better understand the impacts of land use and climate change on the carbon cycle processes at a local scale [10,11], investigations on the spatiotemporal distributions of SOC pools and their changing dynamics are required.

The estimation of the impact of human activity on the environment is a critical issue as clearly highlighted in the 2030 Agenda on Sustainable Development ratified by the UN General Assembly in September 2015. Indeed, the international community is facing critical challenges in addressing the promotion of prosperity and people's well-being while protecting the environment. It is expected that the 2030 agenda will drive the global sustainable development until 2030 and beyond, monitoring the process on the basis of 17 Sustainable Development Goals (SDGs), 169 targets translated into 232 indicators that intend to provide a management tool for countries to implement development strategies. The 2030 Agenda for Sustainable Development clearly suggested the use of Geospatial Information and Earth Observations (EO) to monitor progress and achieve the SDG targets. Our effort is a contribution in the definition of the best practices and experiences regarding the production of critical information necessary for the preservation of natural resources and environment. Our approach is based on the joint use of (i) historical and updated data available from satellite technologies at a global scale free of charge, and (ii) data processing to extract useful information from the investigated dataset. In other words, the use of diverse satellite based proxy indicators are herein proposed as tools to detect, compare, and capture the impacts of human activity over time in a quantitative, objective, and reliable manner. Updated maps and related information are important for decision makers to support and design sustainable development plans following the Agenda 2030 recommendations. In particular, the use of satellite technologies today can suitably support the decision makers and the development plan processes providing suitable free of charge tools useful for multi-temporal analysis based on the past and the current conditions capturing changes at diverse temporal and spatial scales from a global down to a local level.

The estimation and modelling of SOC contents using remote sensing technologies and geographical information systems has increasingly gained importance as reliable tools useful to outline the spatial variability of various soil phenomena, including soil organic carbon [12–14]. In particular, satellite technologies can provide useful information to estimate SOC, allowing quantitative assessments

of SOC contents using proxy indicators, such as spectral indices like the Normalized Difference Vegetation Index (NDVI), Normalized Difference Wetness Index (NDWI), and Enhanced Vegetation Index (EVI). According to several studies [15–24], the reliability of a quantitative assessment of SOC contents depends on statistical prediction models such step-wise linear regression, principal component regression, and partial least squares regression support vector regression (SVR), artificial neural networks (ANN), and random forest (RF) used to infer the volume-dependent SOC content of the soil body [15–25]. Remote sensing technologies and statistical analysis can enable us to overcome the limitation of methods only based on field surveys and laboratory measurements, which provide information limited to the sites where the measurement was done.

In this study we focused on the use of satellite data, statistical analysis and geographical information systems in order to assess: (i) the spatiotemporal changes of SOC in selected areas (located in the eastern Nile Delta), and (ii) the effects on the spatiotemporal changes of SOC.

2. Materials and Methods

2.1. Experimental Site

The study area is located in the eastern Nile Delta (31°, 55 to 32°, 26 E and 30°, 22 to 31°, 05 N; Figure 1a,b). The elevation of the study area decreases from 45 m in the south to sea level in the north. The area is characterized by Mediterranean climate with strong seasonal changes, and annual rainfall averages of around 35 mm year⁻¹ [26], annual temperature average of around 21.7 °C, and the maximum temperatures in July reach 45 °C.



Figure 1. Location of the study area and soil samples.

The study area is quite homogeneous and characterized by two major landscapes: Fluvio-marine flats shaped by fluvial and deltaic processes and river terraces shaped exclusively by alluviation

processes [27]. Considering different land use types and soil types, 120 samples from top soils (0–30 cm) were collected from the study area during the summer of 2018 year. In order to assess the changes in the SOC pool since 1995, the soil sampling points correspond to the sampling locations as introduced in the study by [28] (Figure 1). Field survey included the recording of natural landscape characters as well as farming practices registering farm management, different agricultural operations, tillage, organic additions, irrigation methods, etc. During field survey, each sampling point land-use was coded and allocated to soil type; an attribute table was created including map unit ID, current land use, soil type crop types, agricultural practices, tillage conditions, and soil erosion evidences. The analysis of physical and chemical soil characteristics was achieved based on [29]. Landsat 5 TM images acquired in July 1995 and Operational Land Imager (OLI) with spatial resolution of 30 m acquired in July 2018 were used to distinguish the changes of land use and land cover (LULC) between 1995 and 2018. The support vector machine (SVM) has been applied to discriminate the spatial variation of LULC, where utilized on the images acquired in 1995 and 2018 according to [30].

As the vegetation cover is always indicative of soil fertility and simultaneously indicates the availability of SOC and soil nutrients, the NDVI was used. NDVI maps were generated from the Operational Land Imager (OLI). The spatial heterogeneity in soil organic carbon contents reflects the differentiation of soil properties of both physical and chemical characteristics. Therefore, geomorphological characteristics and surface features were described based on remote sensing and topographic 1: 10,000 map.

2.2. Soil Organic Carbon Pool Calculation

The evaluation of SOC has been relied upon the estimation of SOCP using (Mg C ha⁻¹) according to [31]. The current study focusses on the estimation of SOC changes from 1995 to 2018 with implementing different soil characteristics (soil texture, bulk density, pH, and OM) as well as climate data.

Soil organic matter (SOM) percentage was transformed to SOC percentage referring to [32]. Soil organic carbon pool was calculated for each region, depending on the analysis of the surface soil samples at depth 0–30 cm (kg C m⁻²) using the equation suggested by [33] (see Equation (1));

$$SOCP = \left[L * SOC * B.D * \left(1 - \frac{F}{100}\right)\right] / 10 \tag{1}$$

where SOCP in 0–30 cm depth (kg m⁻²), SOC (wt. %), L thickness of the soil layer in cm, B.D soil bulk density, usually at 33 kPa suction (Mg m⁻³), and F (fine soil fraction) < 2 mm coarse fragment (wt. %).

2.3. Land Surface Temperature Estimation

Land surface temperature was estimated based on thermal remote sensing data (Thermal band 10 of OLI) according to [34] as follows (see Equation (2);

$$Ts = \gamma \left(\varepsilon 10 - 1(\psi 1 * L 10 \text{ sensor } + \psi 2) + \psi 3\right) + \delta$$

$$Ts = \gamma (\varepsilon 10(\psi 1 * L 10 \text{ sensor } + \psi 2) + \psi 3) + \delta$$
(2)

 $\gamma \approx \mathrm{T}^2 10 \ sens(b\gamma * L10 \ sens$

$$\delta \approx T10 \ sens - (T^2 \ 10/b\gamma)$$

Ts is land *surface temperature* (LST) in (Kelvin); $\varepsilon 10$ *is emissivity of Band 10* (unitless). L10sens *is Band 10 radiance* ($\psi 1$, $\psi 2$, $\psi 3$) $A = \pi r^2$ in (W. m⁻². sr⁻¹. μ m⁻¹); T10sens is Band 10 at-sensor brightness, temperature in (K). λ 10 Band 10 effective wavelength in μ m = 10.9000 μ m; b γ = c2/ λ = (1324K) for TIRS⁻¹, Band 10; ($\psi 1$, $\psi 2$, $\psi 3$) *are* three *atmospheric functions*, w is the atmospheric water vapor content in 4.11 g/cm².

2.4. Spatial Distribution Mapping and Validation

Thematic maps have been produced applying the Inverse Distance Weighting (IDW). This method depends on the mean values of a phenomenon by the information of the nearby known points where the neighboring points have more weights than distant points and vice versa [35] (see Equation (3));

$$Z(x_0) = \sum_{i=1}^{n} w_i Z(x_i) w_i = \frac{1/d_i^2}{\sum_{i=1}^{n} 1/d_i^2}$$
(3)

where $Z(x_0)$ is the value of SOC at the un sampled location x_0 , $Z(x_i)$ is the measured values of SOC at the sampled location x_i , n is the number of sample points, w_i is the weight assigned to each $Z(x_i)$, and d_i is the distance among observed and the estimated point.

The maps' accuracy was achieved by comparison of the actual values and the predicted values. As Root Mean Square Error (RMSE) was used where one-third of the total soil samples were selected to verify the model accuracy based on the agreement between the observed and predicted [36].

IDW was chosen according to the recommendation of the study conducted by [37] for mapping soil and sediment parameters. This study suggested that IDW is suitable and reliable and easier to implement compared to kriging, which more time consuming and cumbersome. Even if kriging provides a more accurate description of the data spatial structure and produces valuable information about estimation error distributions, IDW is regarded as a good compromise between the reliability and the complexity of implementation and this is a key issue for the generalization of this type of investigation. The interpolation techniques commonly used in agriculture include inverse distance weighting and kriging [38,39]. Several other studies, however, found inverse distance weighing to be more accurate than kriging. Some studies found that squared inverse distance weighing produced better interpolation results than any other method, including kriging [40].

Evaluating the effects of the human activities on the SOCP using Satellite imagery and GIS techniques is presented by the frame work of the followed methodology, as shown in Figure 2.



Figure 2. Flowchart of the effects parameters on soil organic carbon pool (SOCP) (methodology flowchart).

3. Results and Discussion

3.1. Land Use Changes and Anthopogenic Activities

Figure 3a,b illustrates the magnitude of the land use and land cover (LULC) changes from 1995 to 2018 along with the extension of the new reclamation areas in the eastern Nile Delta. The cultivated area increased around 219,283 hectares from 1995 to 2018. This increment was linked to the necessity to cope with the increased population growth in the same period [41,42].

The results of the change detection from 1995 to 2018, with overall accuracy 93.7% and Kappa Coefficient 0.87, show that along with an increasing agricultural activities in the study area, there also was observed an increasing in urban areas of around 18,585 hectares during the 23 years between 1995 and 2018 (Table 1). In addition, the results of the change detection pointed out that this urban sprawl significantly affected the green surface areas and, thus, increased the emission of carbon dioxide to the surrounded atmosphere [42]. The results of the change detection from 1995 to 2018 show that with increasing agricultural activities in the study area, the urban areas increased by 18,585 hectares during the 23 years between 1995 and 2018 (Table 1). Nevertheless, significant increase in agricultural expansion into the desert areas (on the east of the Suez Canal) reflects an increase in human activity. Additionally, the vegetation cover increased around 37,407.7 hectares between 1995 and 2008, mostly due to the spread of orchard areas for the cultivation of olive and mango trees [43].



Figure 3. Spatial distribution of the land-use and land cover (LULC) between 1995 and 2018: (**a**) spatial distribution of the land-use and land cover in 1995; (**b**) spatial distribution of the LULC in 2018.

Class	1995	2018	Changes		
	(ha)	(ha)	(± ha)		
Agriculture	221,696	240,890	19,194		
Urban	15,929	34,514	18,585		
Sabkha	14,574	18,993	4419		
Bare area	165,791	16,326	-149,465		
Fish farms	25,717	103,554	77,837		
Orchards	17,331	45,901	28,570		
Water bodies	11,553	12,413	860		

Table 1. Land-use change detection from 1995 to 2018 of the study area by hectare.

3.2. NDVI Changes and Human Activity

NDVI is considered an efficient index to describe the vegetation density and photosynthetic activity status [44]. The spatial changes of the vegetation cover between 1995 and 2018 are linked to the agricultural development in the southern parts of the investigated area. This portion of farms mainly depends on modern irrigations systems. The results of the NDVI distribution in 1995 and 2018 (Figure 4) illustrate that in 1995 the highest values of NDVI appeared in the western part of the study area ranging between 0.4 and 0.69. Moreover, since 1995, the complete area is characterized by moderate NDVI values of less than 0.5, corresponding to a mixture of bare soils and areas covered by photosynthetic active vegetation [45,46]. Nevertheless, the spatial distribution of NDVI values in 2018 reveal a distinct increase of the NDVI values in several locations of the study area as a result of human activities during the previous 23 years. The results showed an increase in the agricultural activity in the study area, especially in the new reclamation sites located east and west of the Suez Canal [47]. Therefore, the NDVI values have increased in those areas reaching levels higher to be more than 0.5. These variability changes in the NDVI values were attributed to some factors, including human influences on management methods of land resource, type and quality of fertilization

(minerals or organic), plant status, growth stage, and biomass and photosynthesis intensity [48]. Contrariwise, the effect of the urban sprawl at the north and west of the study area revealed a decrease of the NDVI values, mainly occurring in scattered areas. Moreover, decreasing NDVI values from 0.5–0.69 in 1995 to 0.4–0.62 in 2018 occurred in areas that were affected by salinization processes. These observations are consistent with the findings of [49], who also pointed out the negative impact of soil degradation due to the deterioration of the land cover on the NDVI values.



Figure 4. Normalized Difference Vegetation Index (NDVI) for the study area between 1995 and 2018: (a) NDVI for the study area in 1995; (b) NDVI for the study area in 2018.

3.3. Soil Organic Carbon and Human Activities

The value of soil organic carbon pool (SOCP) in 1995 varied from 0.12 to 60.77 Mg C ha⁻¹, changing gradually to between 0.10 and 65.64 Mg C ha⁻¹ in the year 2018 (Figure 5). In 1995, the high values were found in the northern west of the study area, ranging from 23.11 to 60.77 Mg C ha⁻¹, while in 2018 overall the soil organic carbon contents increased to 24.99 to 65.64 Mg C ha⁻¹. Highest values were observed in the fish farming areas due to additive of organic materials for fish feeding [50]; the expansion of the fish ponds effected an increase of soil organic matter contents. Meanwhile, the results of soil organic carbon values increased in 2018 from 24.99 to 65.64 Mg C ha⁻¹ in 2018.



Figure 5. Spatial distribution of soil organic carbon (SOC) pool for the study area between 1995 and 2018: (a) spatial distribution of SOC pool in 1995; (b) spatial distribution of SOC pool in 2018.

The spatial distribution of the SOC pool was significantly affected by human activity, in particular:

- (i) The effect appeared with the urban sprawl on the cultivated lands characterized by high soil quality; this urban encroachment led to a decrease of areas suitable for agricultural land-use. Nevertheless, the increasing of the population density in Nile Delta initiated a pressure on the needs of increasing the agriculture sector yields [51];
- (ii) An increase in SOC pool values in 2018, which can be explained by the enhancement of the agricultural practices and managements especially in the western parts of the study area. In contrast, in 1995 the agricultural practices predominantly occurred in the western parts of Suez Canal and Wadi Altumilatt [52].

3.4. Effect of Soil Characteristics on SOCP

The dynamic change of SOCP is subjected to several factors besides human actions, including soil properties and exogenous factors such as global climate change. In spite of the homogenous mono cultivation agricultural practices that were applied in the study area region [53], the SOCP values underlie notable variations that can be attributed to the climate effects (Figure 6).



Figure 6. Spatial distribution of the land surface temperature of the study area in 2018.

The spatial distribution of the SOCP in 2018 revealed interesting phenomena paralleled with the effect of the land surface temperature (LST). SOCP values decreased in particular locations to less than 30 Mg C ha⁻¹ compared to 1995 where at the same locations' SOCP totaled more than 30 Mg C ha⁻¹. While variations in SOCP values can be due to several reasons, such as land surface temperature, vegetation cover, human activity, and soil characteristic [54], the spatial correlation with satellite images pointed out that the areas with lower SOCP values in 2018 than 1995 were characterized by increased land surface temperature (LST), exceeding 38 °C in 2018 (Figure 7). An increase of land surface temperature will accelerate decomposition of SOC by fauna and microorganisms, resulting in SOC degradation and release the OC into the atmosphere as CO_2 [55]. Comparable results for both SOCP and LST document the heterogeneity distribution in the study area, as the areas with lower LST values still implement the conventional cultivation systems which are often irregular in agricultural practices. In addition, organic carbon values are observed in the bare areas. Figure 7 shows the negative correlation between the spatial distribution of soil organic carbon pool and the land surface temperature in 2018 as R² was 0.77, as shown in Figure 6. These results are convenient with the finding of [56].



Figure 7. Correlation between soil organic carbon pool (SOCP) and land surface temperature (LST). (a) Correlation between SOCP and NDVI. (b).

Moreover, the increase in the density of vegetative cover has direct effects on SOC due to the increase of crop residue available for the composition of soil organic carbon. Figure 8 shows a positive correlation between NDVI and SOCP with a reasonable correlation coefficient (R^2 0.63). Furthermore, the results point out that a relation between SOCP values and different soil characters: In clay soils the SOCP varied from 1.7 to 53 Mg C ha⁻¹, while in the sandy loam and sand clay loam soils are characterized by minimum values of SOCP around 8 Mg C ha⁻¹, and in clay loam soils minimum SOCP totaled 5.59 Mg C m² (Figure 8).



Figure 8. Effect of different soil types (e.g., silty, sand loam, sand clay, sand, etc.) on the soil organic carbon pool (SOCP).

Sandy soils have relatively low SOCP values ranging from 0.12 to 25 Mg C ha⁻¹. Nevertheless, in the northern part of the study area which is characterized by clayey soils, the SOCP values reduced between 1995 and 2018 due to the deterioration of the physical properties such as the bulk density (Figure 4). The results show that, in 1995 the values were found to be ranging from a minimum of

1.35 gcm³ to maximum of 1.9 gcm³ (with mean of 1.65 gcm³ in 1995), whereas in 2018 were found to be ranging from a minimum of 1.25 gcm³ to maximum of 1.7 gcm³ with mean of 1.67 gcm³. The observed degradation of soil organic carbon pool mainly refers to the relationship between organic matter contents, soil types, soil porosity and bulk density, which well fit with [57,58].

The decrease of SOCP predominantly refers to an increase of land degradation processes in the east of Nile Delta, and corresponds to the effects of management practices in terms of the use of heavy engines for tillage affecting the soil physical properties such structure and soil porosity in the northern Nile Delta [59–61] in the study area reflected by the decomposition of organic material and SOCP (Table 2).

Table 2. Effect of different soil types (e.g., silty, sand loam, sand clay, sand, etc.) on the soil organic carbon pool (SOCP).

	Silty Loam	Sand Loam	Sand Clay Loam	Sand Clay	Sand	Loamy Sand	Loam	Clay Loam	Clay
Min	18.7	36.4	18.7	13.8	4.9	19.4	36.6	25.0	34.6
Mean	29.3	50.9	25.2	28.6	19.4	26.5	41.8	31.7	44.9
Max	35.0	61.1	29.0	50.2	25.9	34.0	49.7	44.8	54.3
Std.	7.3	11.4	4.6	15.4	9.7	5.9	5.6	8.9	8.1

4. Conclusions

The outputs from our investigation clearly highlighted that the use of satellite data can provide useful proxy indicators which (i) enable us to satisfactorily model the SOC pool and (ii) make the approach re-applicable to other study areas and geographic regions also considering that our approach is based on the use of free of charge satellite data (Landsat 5 TM) available systematically for the whole globe from the NASA and USGS website. In particular, our analysis has been based on remote sensing images Landsat 5 TM (acquired in July 1995 and operational land imager (OLI) acquired in July 2018). The results revealed significant variations of soil organic carbon pool (SOCP) among different soil textures and land-uses. Spatiotemporal changes of soil organic carbon pool in the east of the Nile Delta between 1995 and 2018 document the heterogeneity and complexity of processes affecting SOCP values. With our research we document the positive and negative impacts of human activity on SOCP values in agricultural systems of Egyptian drylands. In particular, the change from agriculture to fish farming caused an increase SOCP in the northern parts of the study area as a result of organic additions. Furthermore, the agricultural expansion of new reclaimed areas between 1995 and 2018 caused an increase in vegetation cover due to an increase of cultivation—finally reflected in increasing of soil organic matter contents east of the Suez Canal and some of the southern parts of the study area. In contrast, SOCP values were decreased in some areas in the north and west of the study site as a result of the negative management practices causing increased land degradation processes such as increasing soil bulk density and soil salinity that in turn negatively affects biological activity in the soils.

As a whole, our results pointed out that the use of historical archives as well as the past and current space data today available, even updated weekly, provides an excellent data source suitable to monitor human activity and support systematic low cost monitoring over large areas. In actuality, satellite data provide both (i) historical time-series data set and (ii) timely updated information related to the current and past human activities in terms of land use/land cover and their characteristics as well as parameters (proxy indicators) suitable to assess, characterize, and better understand and manage the impact of the human activities on environment. Updated maps and related information are very important for planners to design sustainable development plans following the Agenda 2030 recommendations. In particular, the use of satellite technologies today can suitable support the decision maker and the development plan processes providing suitable free of charge tools useful for multi-temporal analysis based on the past and the current conditions capturing changes at diverse temporal and spatial scales from global, down to a local level.

Author Contributions: The research article included three main contributions; field survey, satellite imagery interpretation, and writing the text. M.A. and M.A.E.A. provided the field survey. The satellite image data have been analyzed by M.E. and M.A.E.A. The text of article was written by M.A. and B.S. The last version of the article was revised by R.L. and B.S.

Acknowledgments: The manuscript presented an efficient scientific participation between the scientific institutions in three countries (Egypt, German and Italy). The authors would like to thank the University of Basilicata and Italian National Research Council (CNR) at Potenza, Italy, and the institute of Geographic Sciences, FU Berlin, Germany, for supporting the research activities. Special thanks to the National Authority for Remote Sensing and Space Science (NARSS) for funding the satellite data and the field survey.

Conflicts of Interest: The authors would like to hereby certify that no conflict of interest in the data collection, analyses, and the interpretation; in the writing of the manuscript, and in the decision to publish the results. Authors would like also to declare that the funding of the study has been supported by the authors' institutions.

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