

Article

Quantitative Analysis of Infrared Thermal Images in Rock Fractures Based on Multi-Fractal Theory

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Abstract: Rock disasters caused by rock fractures seriously threaten the safe and sustainable mining of deep coal resources. The infrared thermal imaging of rock fractures has a non-uniform distribution. This is important information for the monitoring and early warning of rock mass instability. In this paper, the multi-fractal theory was introduced to analyze the infrared thermal image obtained from the uniaxial loading of sandstone, which is expected to provide quantitative indicators for the infrared monitoring of rock failure. The results show that the multi-fractal parameters $\Delta\alpha$ (non-uniformity of temperature) and Δf (frequency diversity of minimum and maximum temperature) can be used to describe the distribution of the thermal field; they are sensitive to the rock macro fracture. Both $\Delta\alpha$ and Δf are constant during the initial loading stage. When the samples yield and there is a failure in the later stage of loading, the $\Delta\alpha$ and Δf change abruptly. The sudden change in $\Delta\alpha$ and Δf can be regarded as the precursor to rock failure. The research results preliminarily show the feasibility and potential of multi-fractal analysis in rock mass disaster monitoring and early warning.

Keywords: quantitative analysis; infrared thermal image; multi-fractal theory; sandstone; fracture



Citation: Miao, B.; Wang, X.; Li, H. Quantitative Analysis of Infrared Thermal Images in Rock Fractures Based on Multi-Fractal Theory. *Sustainability* **2022**, *14*, 6543. <https://doi.org/10.3390/su14116543>

Academic Editor: William Frodella

Received: 16 April 2022

Accepted: 22 May 2022

Published: 27 May 2022

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1. Introduction

China's coal resources have gradually entered the deep stratum, due to the depletion of shallow resources. In the process of deep mining, coal and rock mass are in a more severe geomechanical environment of high geostress, high ground temperature and strong dynamic disturbance. Moreover, rock mass disasters (e.g., rockburst, roof caving) caused by rock fracture occur often [1–4], which seriously threatens the safety and sustainability of mining. To predict these disasters, various monitoring methods have been proposed and applied, such as acoustic emission (AE) [5], electromagnetic radiation (EMR) [6,7] and infrared radiation (IRR) [8–11]. Among them, infrared technology is considered to be a potential method for monitoring rock mass instability; it assumes that the infrared radiation temperature (IRT) of the rock surface changes in the process of rock deformation and fracture [12–16]. Infrared technology has the characteristics of being non-destructive, non-contact and observable. It is also widely used and studied in different rock engineering fields, such as roadway excavation [17] and landslides [18].

A large number of scholars have studied the characteristics of IRR in the rock loading process. Deng et al. [19,20] found that the IRT of rock increases with an increase in stress; an abnormal high or low temperature region on the infrared thermal image appears when the rock fractures. Wu et al. [21–24] explored the characteristics of the temperature–time curve and infrared thermal image in the rock loading process through different loading experiments (uniaxial compression, shear, impact, etc.). They took the abrupt change in average infrared radiation temperature (AIRT) and infrared thermal image as the precursor to rock failure. Moreover, the shear fracture of rock causes the IRT to rise and the tensile

fracture causes the IRR to decrease. Freund [25] studied the IRR during the loading process of granite and put forward the P-hole theory; he considered the electron flow in the rock as the source of IRR. Cao et al. found that the AIRT rate of rock before and after dilatancy changes [26,27]. Sun et al. [16] compared the AE and IRR characteristics of rocks during the true triaxial loading and unloading experiments. Li et al. [28] investigated the characteristics of IRR during the loading of coal containing gas. However, all the above research focused on the application and mechanism of infrared technology, while most descriptions of IRT and infrared thermal images are qualitative, which is not conducive to the further practical application of such technology to engineering.

Therefore, some work has been done to put forward the quantitative description method of IRR and to try to find the effective indicators for the early warning signs of rock mass instability. For instance, Liu et al. [29] and Ma et al. [30,31] used variance, entropy, feature roughness and *b*-value to capture the evolution characteristics of infrared thermal images. Shen et al. [8] adopted the method of critical slowing down to quantitatively analyze the maximum infrared radiation temperature (MaxIRT). In fact, when the rock fracture occurs, the infrared thermal image clearly shows areas of either abnormally high temperature (shear fracture zone) or abnormally low temperature (tensile fracture zone), and the temperature distribution is uneven. Moreover, the simultaneous existence of a high-temperature anomaly area and low-temperature anomaly area may cause little change in AIRT. Therefore, a better description of this infrared thermal image feature is very important and necessary. On the other hand, the multi-fractal theory can be used to describe the complexity and heterogeneity of data and it has been widely used in different disciplines [32–35]. Thus, it is possible to introduce the multi-fractal theory to the analysis of complicated infrared thermal images.

In this paper, the uniaxial compression test of sandstone samples was carried out and the IRR was collected simultaneously. The multi-fractal theory was introduced to quantitatively describe the infrared thermal image characteristics of the rock loading process, which is expected to provide new indicators and guidance for the monitoring and early warning signs of rock disasters.

2. Experimental Process and Analysis Method

2.1. Materials and Experimental System

According to the standards of the international society of rock mechanics (ISRM), standard samples of $\Phi 50 \text{ mm} \times 100 \text{ mm}$ were prepared to carry out the uniaxial compression tests. The samples were made of red sandstone from the Linfen mining area, Shanxi Province. The density of the rock used is 2454.65 kg/m^3 . In order to ensure the homogeneity of the samples, the P-wave velocity of the samples was tested. The wave velocities of all samples are about 3.51 km/s and the homogeneity between samples is good.

The test system consists of two main parts: a loading system and an IRR data acquisition system. The loading system adopted the new SANS microcomputer-controlled electro-hydraulic servo pressure testing machine. The maximum load capacity is 3000 kN . The IRR data acquisition system adopted the Optris PI450 high-resolution infrared thermal imager. The optical resolution is 288×382 pixels; the temperature measurement range is $-20 \sim 100 \text{ }^\circ\text{C}$; the spectral range is $7.5 \sim 13 \text{ }\mu\text{m}$; the frame frequency is 80 Hz ; the field-of-view angle is $30^\circ \times 23^\circ$; and the thermal sensitivity is 0.04 k .

To reduce the interference of the external environment, all tests were carried out in an AFGP-II high-efficiency electromagnetic-shielding room. The infrared thermal imager was arranged 50 cm away from the sample and the infrared thermal image acquisition rate was set at 10 Hz . Before the test, the infrared thermal imager was turned on for about 30 min to preheat and the loading was started after the imager was stable. The loading rate was controlled as 400 N/s . Once loading was completed, the force, displacement and the IRR data were synchronously measured and stored. The test was finished upon the failure of the sandstone.

2.2. IRR Principle of Rock Fracture

Any object with a temperature higher than absolute zero ($-273.15\text{ }^\circ\text{C}$) emits waves in a very wide wavelength range in the form of electromagnetic radiation. IRR is an electromagnetic wave in the wavelength range of $0.7\text{--}1000\text{ }\mu\text{m}$ and the part with a wavelength of $2.5\text{--}15\text{ }\mu\text{m}$ has a thermal effect. The IRR wave is collected by an infrared thermal imager and converted into temperature information according to the Stefan–Boltzmann law [16,36]:

$$W_e = \varepsilon\alpha T^4 \tag{1}$$

where W_e is the radiant energy intensity; ε is the emissivity; α is a constant; and T is the absolute temperature (K). The principle is shown in Figure 1. Under the action of stress, the internal energy of rock transforms and causes IRR. The IRR can reflect the stress state of rock. The change in infrared radiation temperature is the comprehensive result of the thermoelastic effect, pore gas desorption and escape effect, fracture effect, friction thermal effect, heat transfer effect and environmental radiation effect [24,37].

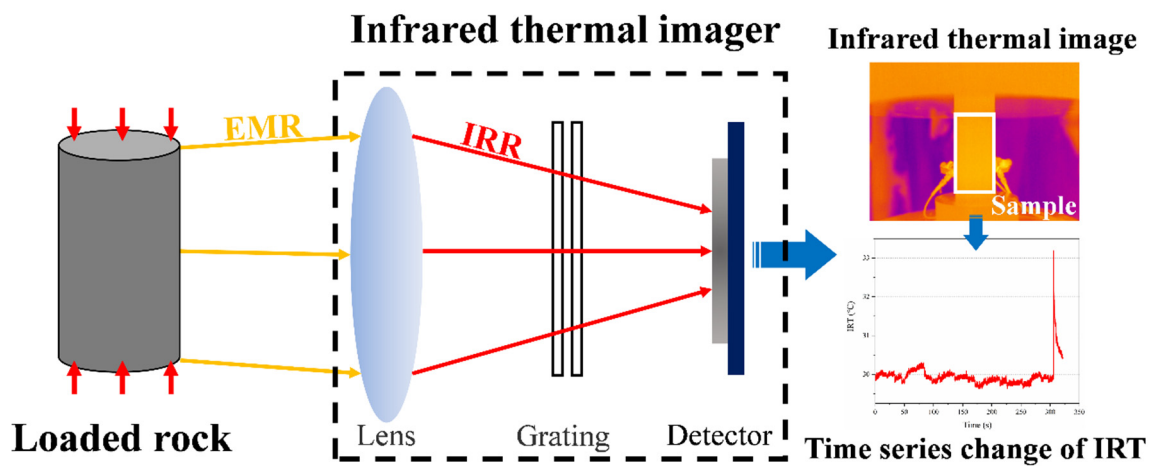


Figure 1. Schematic diagram of infrared technology.

2.3. Multi-Fractal Theory

The upper left of Figure 2 shows the original infrared thermal image, which is essentially a matrix composed of IRT data. Each pixel corresponds to the IRT data of the corresponding position. If the optical resolution of the instrument is 288×382 pixels, the output result will be an IRT matrix $T_{288 \times 382}$ with the size of 288×382 . The IRT matrix $x_{m \times n}$ at the sample position (valid area) is extracted by MATLAB software, and the multi-fractal characteristics of the data in the matrix are analyzed according to the literature [38,39].

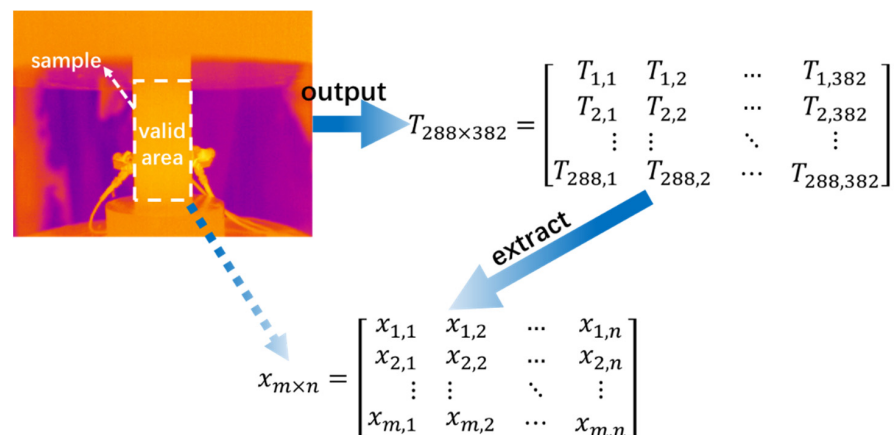


Figure 2. Infrared thermal image data extraction process.

The matrix $x_{m \times n}$ is divided into boxes of size $k \times k$. The probability density of the box subset (i, j) is $\{P_{ij}(k)\}$. The probability distribution $\{P_{ij}(k)\}$ is calculated for each subset:

$$X_q(k) \equiv \sum P_{ij}(k)^q \sim k^{\tau(q)} \quad (2)$$

$$\tau(q) = \lim_{k \rightarrow 0} \frac{\ln X_q(k)}{\ln k} \quad (3)$$

$$\alpha = \frac{d(\tau(q))}{dq} = \frac{d}{dq} \left(\lim_{k \rightarrow 0} \frac{\ln X_q(k)}{\ln k} \right) \quad (4)$$

$$f(\alpha) = \alpha q - \tau(q) \quad (5)$$

Among them, $X_q(k)$ is the defined partition function, that is, the statistical moment. $\tau(q)$ is the quality index, with $-\infty < q < +\infty$. α is a constant, called the singularity index, which controls the singularity of $\{P_{ij}(k)\}$. It reflects the non-uniformity of the probability subset under different k . $f(\alpha)$ means the frequency of the subset represented by α in all subset sets and it is also the fractal dimension of subset α .

The curve $\alpha - f(\alpha)$ is the multi-fractal spectrum of the calculated matrix. The width of the spectrum $\Delta\alpha = \alpha_{max} - \alpha_{min}$ reflects the differences among the distribution of temperature data. The larger the value of $\Delta\alpha$, the greater the non-uniformity of temperature and the more obvious the multi-fractal characteristics. The difference between the frequency of the lowest temperature data and that of the highest temperature data can be obtained by $\Delta f = f(\alpha_{max}) - f(\alpha_{min})$. A smaller value of Δf represents a greater proportion of high temperature data and vice versa. $\Delta f > 0$ indicates that the low temperature data dominates, while $\Delta f < 0$ signifies the domination of high temperature data.

Therefore, the values of $\Delta\alpha$ and Δf can be used to describe the differences in infrared thermal images and the dominant temperature data, respectively. The evolution of these two multi-fractal indicators during the loading can be used to identify the characteristics of the thermal field.

3. Analysis and Discussion of Experimental Results

3.1. Time Series Change in IRT

Two typical samples were taken as examples for subsequent analysis, and their numbers were S-1 and S-2, respectively. Figure 3 shows the stress and strain of the two samples during the loading process, as well as the changes in MaxIRT, AIRT and the minimum IRT (MinIRT) on the surface of the samples. The loading process of rock can be roughly divided into three stages. I (Compaction stage): the primary pores are closed, the deformation is large and the deformation rate decreases gradually. II (Elastic deformation stage): the strain increases linearly with the increase in stress. III (Yield and failure stage): the strain increases unsteadily, and a large number of micro fractures gather to form macro fractures. When the stress reaches the peak value, the failure occurs.

In the early stages (I and II) of loading, the changes in MaxIRT, AIRT and MinIRT are highly synchronous, which indicates that the surface temperature changes are uniform and integral. The initial MaxIRT, AIRT and MinIRT of sample S-1 are about 31.19 °C, 30.58 °C and 20.8 °C, respectively, and then the curves show a slow decline trend, but little change. The initial MaxIRT, AIRT, and MinIRT of sample S-2 are about 30.04 °C, 29.63 °C, and 29.02 °C, respectively, and then the curves are slightly inflected, but they are also basically unchanged. In most experiments, the IRT of the rock increases during the loading process, but occasionally the IRT continues to decrease, which may be caused by the exhaust gas after the rock pore is compressed [8]. The abnormal temperature fluctuation in the early stages of loading may be caused by changes to the microstructure inside the sample, such as the closure and propagation of microcracks caused by stress adjustment [14].

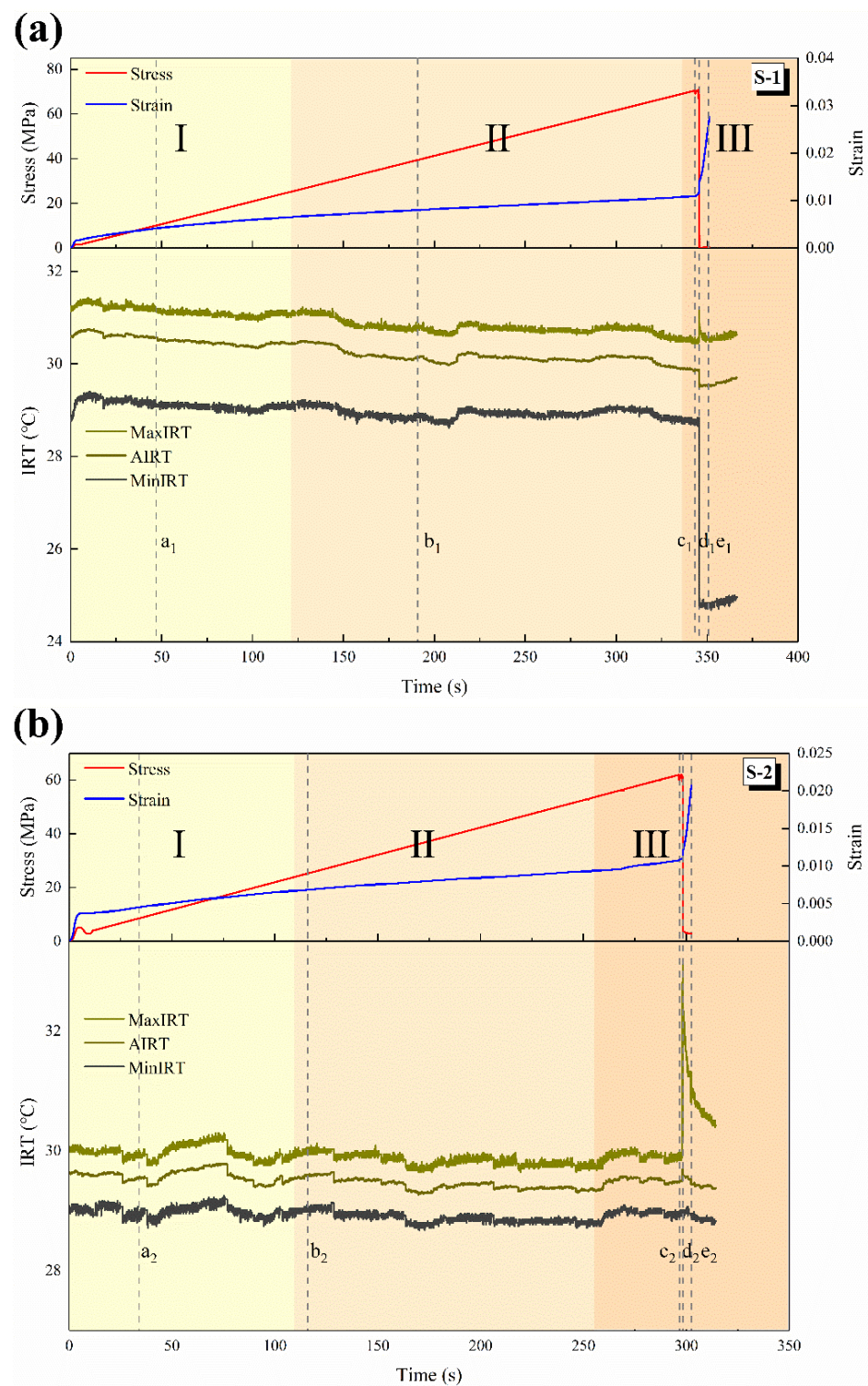


Figure 3. Time series change in IRT: (a) S-1; (b) S-2. (Different color blocks distinguish stage I, II and III).

In stage III, the specimens begin to show obvious macro fractures and the surface infrared temperature fields show local abnormal changes. In the moment of failure, the MaxIRT of sample S-1 increases from 30.54 °C to 31.16 °C and then gradually drops, while AIRT drops and MinIRT drops significantly. For sample S-2, MaxIRT significantly increased from 29.94 °C to 33.19 °C and then gradually cooled down, with AIRT being slightly increased, while MinIRT remained basically unchanged. This is related to the failure modes of different specimens, which will be discussed in more detail in the next section.

3.2. Multi-Fractal Analysis of Infrared Thermal Images

The infrared thermal images of two samples at several typical moments (a_1 - e_1 , a_2 - e_2) were selected. The positions of different time points are shown in Figure 3. The multi-fractal characteristics of IRT field data at each time were calculated. Figures 4 and 5a show the calculation results of sample S-1. The infrared thermal images in all the figures were redrawn by Surfer software. At the time of a_1 (47.03 s), the $\Delta\alpha$ is low, meaning that the degree of non-uniformity of the infrared temperature field data distribution is small, and Δf is close to 0. With the loading process, the infrared temperature field changes as a whole. At the time of b_1 (190.88 s), the values of $\Delta\alpha$ and Δf show little change. At c_1 (343.45 s), the loading is close to the peak stress of the specimen, and a macro fracture occurs. The temperature field of the surface is in the abnormally high-temperature region and develops continuously. At d_1 (345.73 s), the sample is destroyed, and there is an obvious shear fracture zone on the surface and some rock blocks fall off in the tensile fracture zone. The MaxIRT of the fracture zone is 31.16 °C and the temperature of the falling off area is the temperature of the background environment. In the process of c_1 - d_1 , the difference in infrared temperature distribution on the surface of samples gradually increases and $\Delta\alpha$ increases. The performance of high temperature data is gradually dominant, so Δf decreases. e_1 (350.79 s) is the moment after the failure of the sample, at which time the temperature decreases, but the difference in the temperature distribution is still significant.

Figures 5b and 6 show the calculation results of sample S-2. Local shear failure occurs in the sample and the changing trend of $\Delta\alpha$ and Δf is similar to that of sample S-1. However, during the failure of the sample, the increase in $\Delta\alpha$ and the decrease in Δf are much lower than those of S-1. The reason may be that the falling off of some rocks increases the difference in the infrared temperature field distribution of S-1. The shear fracture area of S-1 is more serious, so the high temperature data performance is better, even if the increase in MaxIRT of S-2 is greater than that of S-1.

In general, $\Delta\alpha$ and Δf are very sensitive to the macroscopic fracture of the rock. In the early stages of loading, the IRT of the sample surface changes uniformly as a whole and $\Delta\alpha$ and Δf are basically unchanged. Entering the yield and failure stage, the macro-fracture of the sample results in changes to $\Delta\alpha$ and Δf . This shows the validity of the indicators. The sudden change in $\Delta\alpha$ and Δf can be used as a precursor to the occurrence of rock disasters.

In addition, the experimental results show that the variations in $\Delta\alpha$ and Δf are closely related to the failure mode of the rock. Different fractures lead to different infrared temperature distributions. For example, due to the frictional thermal effect, the more severe the shear fracture, the better the high temperature data performance and the more uneven the overall temperature distribution. This results in an increase in $\Delta\alpha$ and a decrease in Δf . However, the failure mode of the rock is related to rock type, loading mode, water bearing state and so on. Therefore, we intend to carry out more experiments to establish a deeper relationship between rock fracture and multi-fractal indicators.

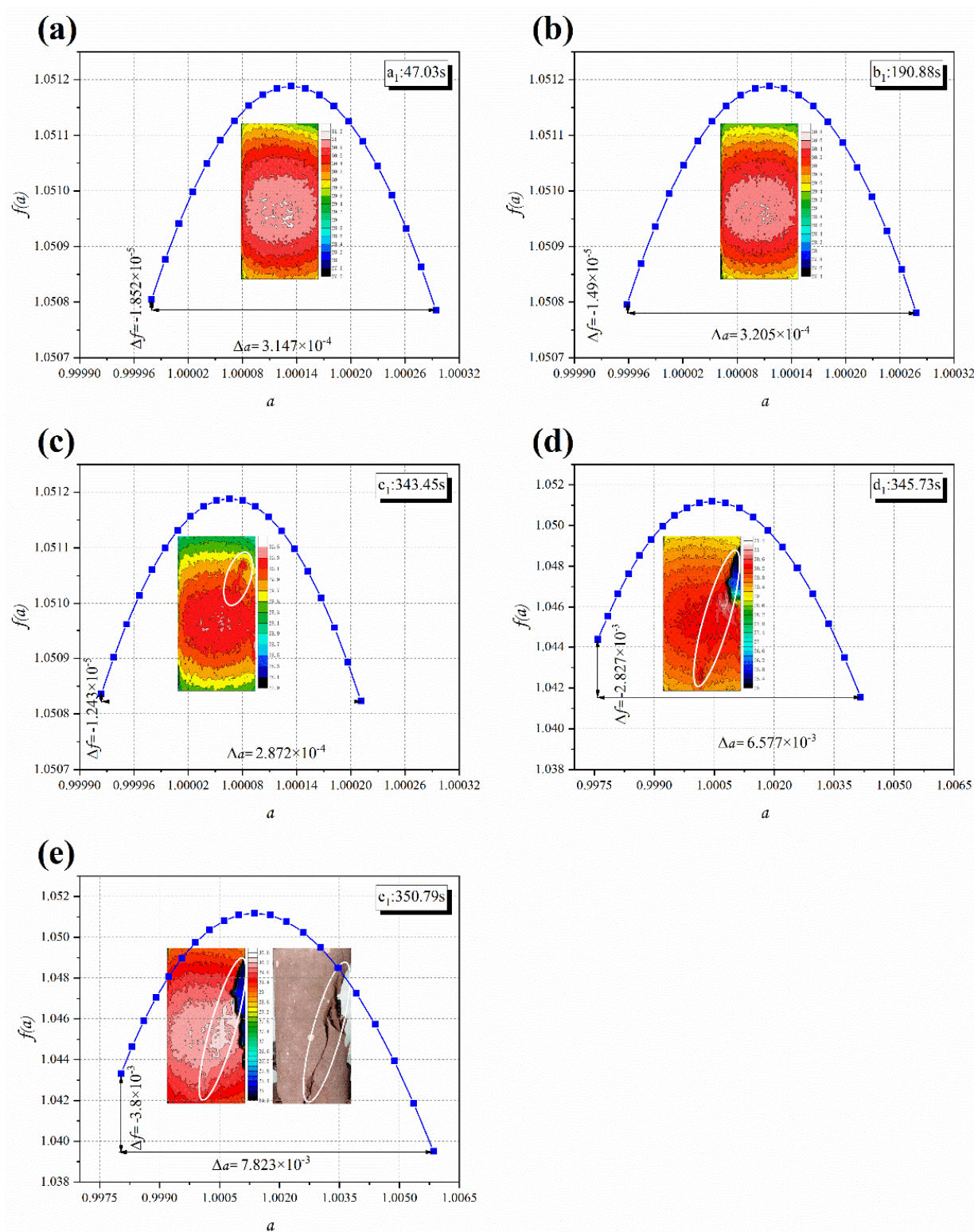


Figure 4. Multi-fractal features of infrared thermal images at different time points of S-1: (a) a_1 ; (b) b_1 ; (c) c_1 ; (d) d_1 ; (e) e_1 .

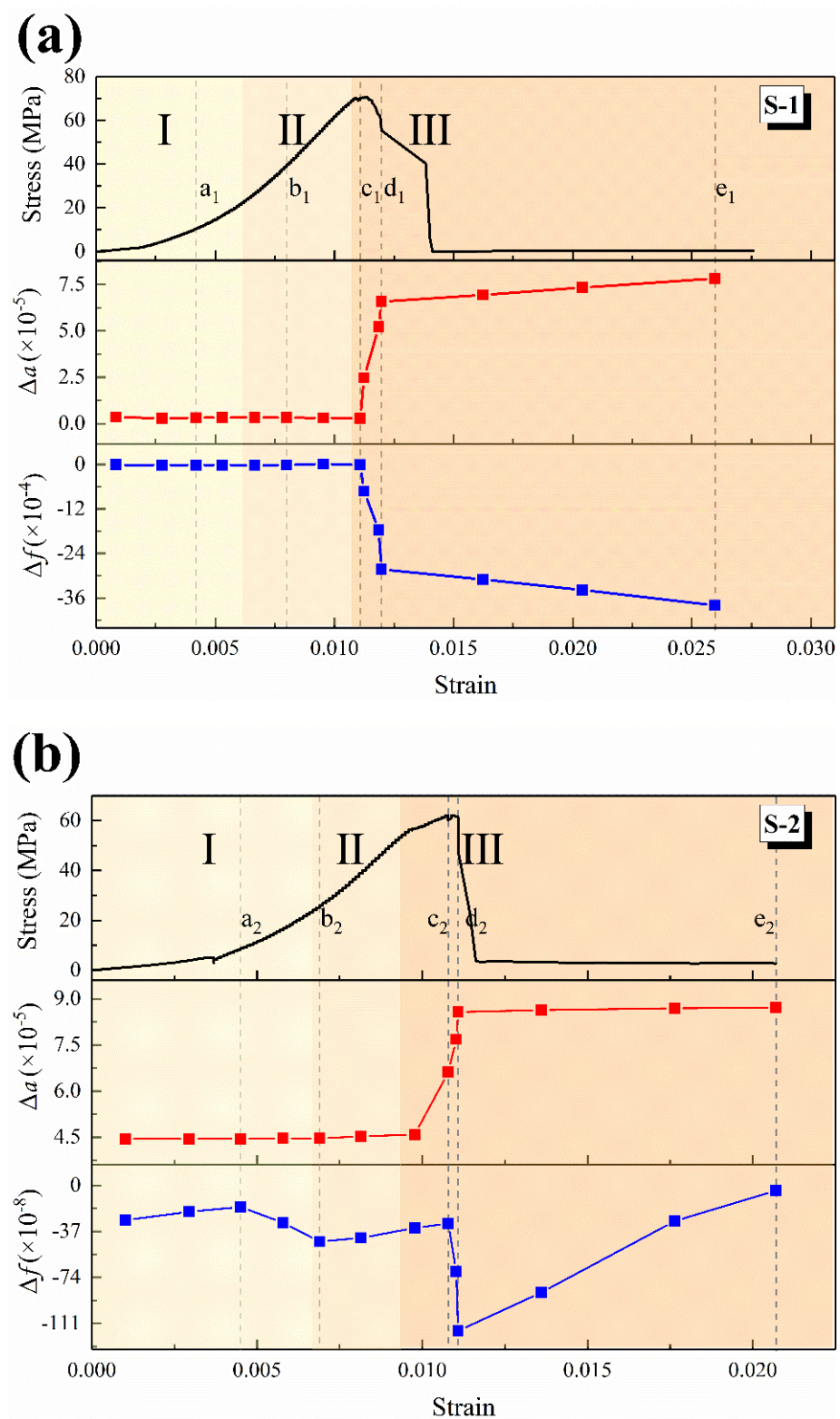


Figure 5. Dynamic changes of multi-fractal indicators of infrared thermal image: (a) S-1; (b) S-2. (Different color blocks distinguish stage I, II and III).

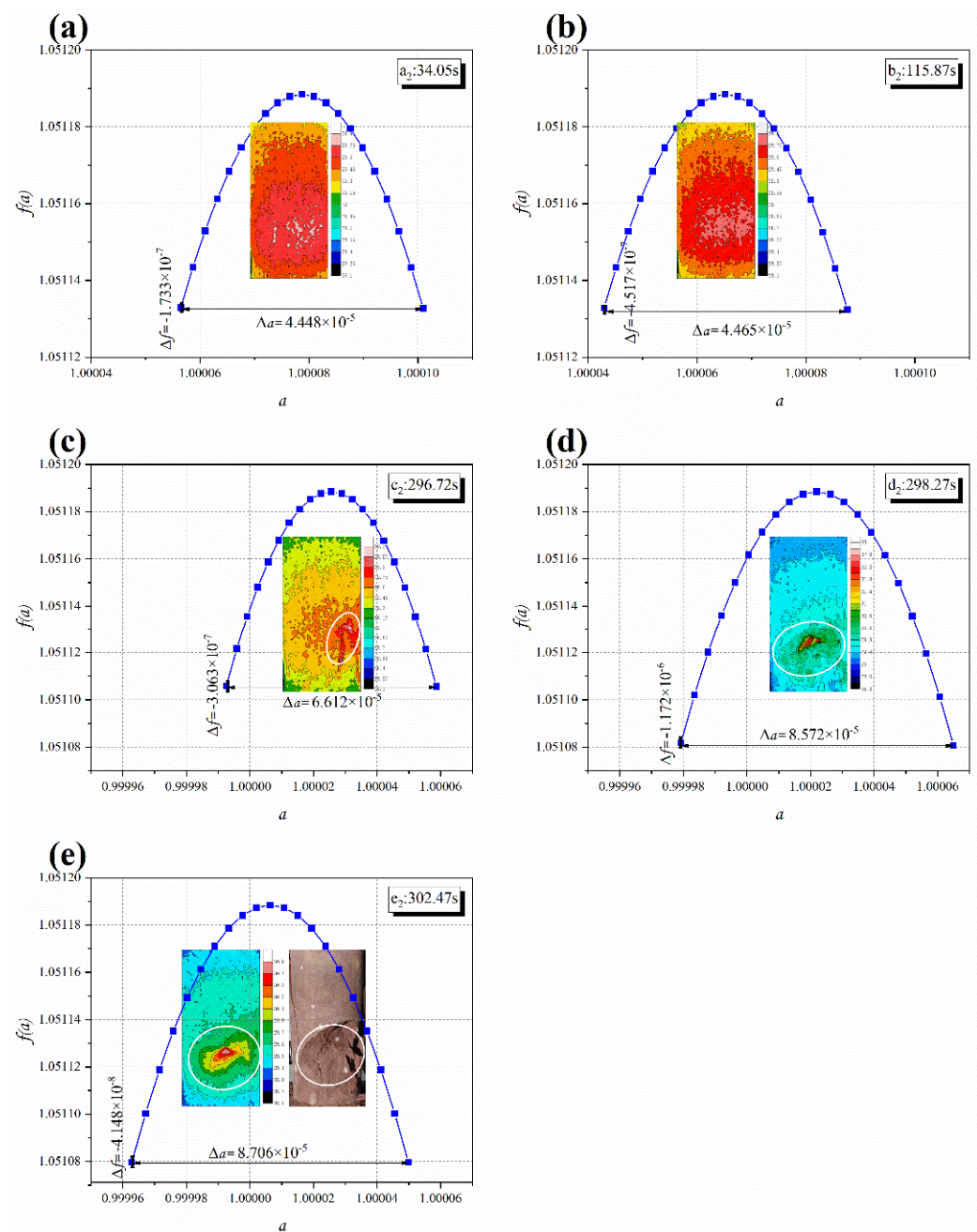


Figure 6. Multi-fractal features of infrared thermal images at different time points of S-2: (a) a_2 ; (b) b_2 ; (c) c_2 ; (d) d_2 ; (e) e_2 .

4. Conclusions

In this paper, a uniaxial compression experiment of sandstone is carried out and the infrared radiation is monitored at the same time. The multi-fractal theory was introduced to analyze infrared thermal images. The infrared thermal image of rock during loading has multi-fractal characteristics. The multi-fractal parameters $\Delta\alpha$ (non-uniformity of temperature) and Δf (frequency diversity of minimum and maximum temperature) are not sensitive to changes to the infrared thermal image in the early stage of loading, but are more sensitive to the macro fracture of rock. When the samples yield and fail in the later stage of loading, they have an obvious response. The sudden change to $\Delta\alpha$ and Δf can be regarded as the precursor of rock failure. The research results preliminarily show that the multi-fractal analysis of infrared thermal images can provide reference for the monitoring and early warning signs of rock mass disasters, which has great application value. In the

future, we intend to establish a deeper relationship between multi-fractal indicators and rock fracture.

Author Contributions: Conceptualization, X.W.; writing—original draft preparation, B.M.; writing—review and editing, X.W.; methodology, H.L.; software, H.L. All authors have read and agreed to the published version of the manuscript.

Funding: This research received no external funding.

Conflicts of Interest: The authors declare no conflict of interest.

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