



## Article

# CO<sub>2</sub> Flux over the Contiguous United States in 2016 Inverted by WRF-Chem/DART from OCO-2 XCO<sub>2</sub> Retrievals

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**Citation:** Zhang, Q.; Li, M.; Wang, M.; Mizzi, A.P.; Huang, Y.; Wei, C.; Jin, J.; Gu, Q. CO<sub>2</sub> Flux over the Contiguous United States in 2016 Inverted by WRF-Chem/DART from OCO-2 XCO<sub>2</sub> Retrievals. *Remote Sens.* **2021**, *13*, 2996. <https://doi.org/10.3390/rs13152996>

Academic Editor: Sang-Hyun Lee

Received: 19 June 2021

Accepted: 27 July 2021

Published: 30 July 2021

Corrected: 14 March 2022

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**Abstract:** High spatial resolution carbon dioxide (CO<sub>2</sub>) flux inversion systems are needed to support the global stocktake required by the Paris Agreement and to complement the bottom-up emission inventories. Based on the work of Zhang, a regional CO<sub>2</sub> flux inversion system capable of assimilating the column-averaged dry air mole fractions of CO<sub>2</sub> (XCO<sub>2</sub>) retrieved from Orbiting Carbon Observatory-2 (OCO-2) observations had been developed. To evaluate the system, under the constraints of the initial state and boundary conditions extracted from the CarbonTracker 2017 product (CT2017), the annual CO<sub>2</sub> flux over the contiguous United States in 2016 was inverted (1.08 Pg C yr<sup>-1</sup>) and compared with the corresponding posterior CO<sub>2</sub> fluxes extracted from OCO-2 model intercomparison project (OCO-2 MIP) (mean: 0.76 Pg C yr<sup>-1</sup>, standard deviation: 0.29 Pg C yr<sup>-1</sup>, 9 models in total) and CT2017 (1.19 Pg C yr<sup>-1</sup>). The uncertainty of the inverted CO<sub>2</sub> flux was reduced by 14.71% compared to the prior flux. The annual mean XCO<sub>2</sub> estimated by the inversion system was 403.67 ppm, which was 0.11 ppm smaller than the result (403.78 ppm) simulated by a parallel experiment without assimilating the OCO-2 retrievals and closer to the result of CT2017 (403.29 ppm). Independent CO<sub>2</sub> flux and concentration measurements from towers, aircraft, and Total Carbon Column Observing Network (TCCON) were used to evaluate the results. Mean bias error (MBE) between the inverted CO<sub>2</sub> flux and flux measurements was 0.73 g C m<sup>-2</sup> d<sup>-1</sup>, was reduced by 22.34% and 28.43% compared to those of the prior flux and CT2017, respectively. MBEs between the CO<sub>2</sub> concentrations estimated by the inversion system and concentration measurements from TCCON, towers, and aircraft were reduced by 52.78%, 96.45%, and 75%, respectively, compared to those of the parallel experiment. The experiment proved that CO<sub>2</sub> emission hotspots indicated by the inverted annual CO<sub>2</sub> flux with a relatively high spatial resolution of 50 km consisted well with the locations of most major metropolitan/urban areas in the contiguous United States, which demonstrated the potential of combing satellite observations with high spatial resolution CO<sub>2</sub> flux inversion system in supporting the global stocktake.

**Keywords:** regional CO<sub>2</sub> flux inversion; OCO-2; WRF-Chem/DART; ensemble adjustment Kalman filter; the contiguous United States

## 1. Introduction

Since the industrial revolution, the significant increase of carbon dioxide (CO<sub>2</sub>) in the atmosphere caused by anthropogenic activities is believed to be one of the main driving forces of climate change [1]. The atmospheric CO<sub>2</sub> concentration has increased from 277 parts per million (ppm) in 1750 to above 409 ppm in 2019 [2].

In order to mitigate climate change, the Paris Agreement [3], a legally binding international treaty on climate change, was adopted at the 21st Conference of the Parties of the United Nations Framework Convention on Climate Change (UNFCCC) and entered into force on November 4, 2016. The Paris Agreement [3] requires all parties to submit their emission reduction plans in the form of nationally determined contributions (NDC) and requires regular global stocktake processes to verify whether all parties have fulfilled their commitments.

Currently, the statistical bottom-up methodology developed by the Inter-governmental Panel on Climate Change (IPCC) is used to prepare the emission inventory reports required by UNFCCC [4]. This methodology uses statistical activity data and source-specific emission factors to quantify the anthropogenic CO<sub>2</sub> emission of a country or a region. However, due to large uncertainties in the emission factors and in the statistical activity data, as well as biases caused by unaccounted sources, the bottom-up emission inventories can have significant uncertainties, especially for countries with underdeveloped statistical infrastructure. Therefore, with the progress in observations and inverse models, the Subsidiary Body for Scientific and Technological Advice (SBSTA) of UNFCCC [5,6] along with the IPCC Task Force [4] acknowledged the complementary capability offered by the top-down observation approach, which links emissions with atmospheric concentration observations by atmospheric transport and chemistry models, i.e., the inverse models.

Before satellite-based observations of atmospheric CO<sub>2</sub> concentrations were widely available, the research of top-down approach mainly focused on inverting CO<sub>2</sub> flux from ground-based observations [7–12]. Compared with ground-based observations, satellite observations have the advantage of providing much better spatiotemporal coverage. With the development of technology, in the past 10 years or so, several CO<sub>2</sub> monitoring satellites—including the Greenhouse gases Observing SATellite (GOSAT) [13], GOSAT-2 [14], the Orbiting Carbon Observatory 2 (OCO-2) [15], OCO-3 [16,17], and TanSat [18], had been launched. Several others, such as MicroCarb [19,20], geoCARB [21,22], and a CO<sub>2</sub>M constellation [23,24]—are currently under development.

Promoted by the CO<sub>2</sub> monitoring satellites, recent studies used the column-averaged dry air mole fractions of CO<sub>2</sub> (XCO<sub>2</sub>) products retrieved from satellite observations to inverse CO<sub>2</sub> flux [25–30]. Among them, Deng et al. [31] used the four-dimensional variational (4D-Var) approach coupled with GEOS-Chem model to demonstrate the significant impact of GOSAT observations on the flux inversion, Crowell et al. [32] utilized different atmospheric transport models, data assimilation methods and prior fluxes to inverse the global CO<sub>2</sub> fluxes from OCO-2 observations and evaluate their uncertainties, Chevallier, et al. [33] and Wang et al. [34] assimilated observations from GOSAT and OCO-2 respectively to conduct global CO<sub>2</sub> flux inversion. Although these global-scale inversion studies can deepen the understanding of the global carbon cycle mechanism that affects climate change, due to their coarse spatiotemporal resolution, their results are not enough to support the global stocktake required by the Paris Agreement.

To complement the bottom-up emission inventories, regional flux inversion systems with high spatiotemporal resolution are needed. Villalobos et al. [35] used the Community Multiscale Air Quality (CMAQ) transport-dispersion model coupled with 4D-Var to study the impact of OCO-2 observations on CO<sub>2</sub> flux inversions and on their uncertainty reduction over Australia through a series of Observing System Simulation Experiments (OSSEs). Zheng et al. [36] and Peng et al. [37] developed regional CO<sub>2</sub> flux inversion systems based on WRF and CMAQ modeling system coupled with 4D-Var and ensemble Kalman filter approach, and then evaluated the system performances in the continental United States and East Asia through OSSEs, respectively. However, real XCO<sub>2</sub> retrievals from OCO-2 were not yet utilized in these studies.

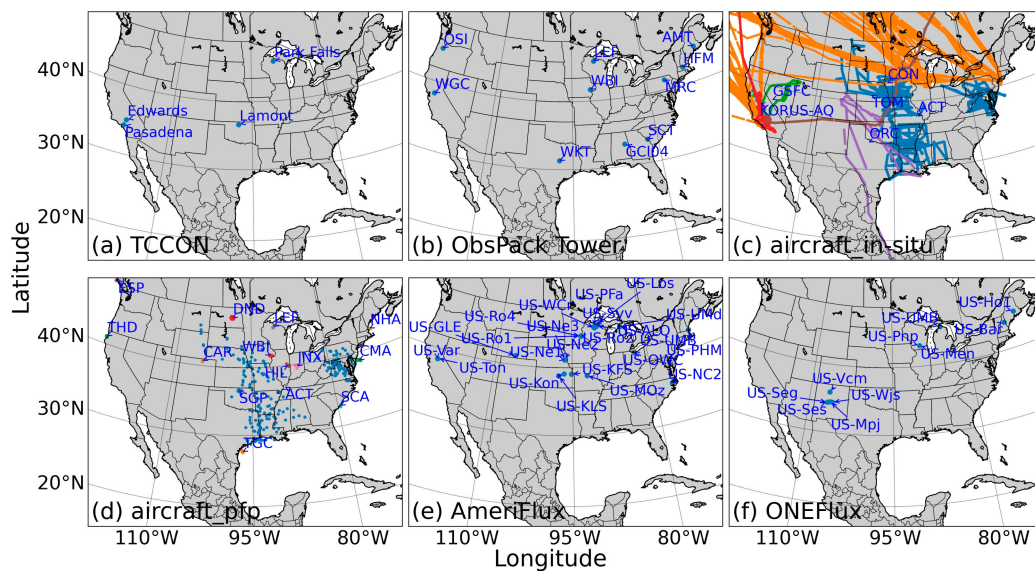
In this study, based on the work of Zhang et al. [38], a regional CO<sub>2</sub> flux inversion system had been developed from WRF-Chem model [39] coupled with Data Assimilation Research Testbed (DART) [40] using the ensemble adjustment Kalman filter (EAKF) [41,42] assimilation method to invert regional CO<sub>2</sub> flux from OCO-2 XCO<sub>2</sub> retrievals. The contigu-

ous United States was selected as the study area because there were relatively abundant ground-based observations available for evaluation. The study period was the entire year of 2016. Different types of CO<sub>2</sub> concentration observations, CO<sub>2</sub> flux measurements, as well as the products from the OCO-2 model intercomparison project (MIP) [32] and CT2017 [10], were used to evaluate the performance of the inversion system over the contiguous United States.

## 2. Materials and Methods

### 2.1. CO<sub>2</sub> Transport Model

WRF-Chem model version 3.9.1 [39] was used as the CO<sub>2</sub> transport model. As shown in Figure 1, the model domain, having 103 (west–east) × 82 (south–north) grid points with a spatial resolution of 50 km, covered the contiguous United States. In the vertical direction, the atmosphere was decomposed into 45 terrain-following levels from the surface to the upper boundary at 50 hPa. The physical and chemical options of the model, listed in Table 1, were the same as those used in Zhang, et al. [38]. The chemical option “chem\_opt = 16” means that CO<sub>2</sub> is treated as a passive tracer with no feedback on meteorological variables and is not part of any chemical reactions.



**Figure 1.** The WRF-Chem domain and locations of ground-based and aircraft observations for evaluation. (a) TCCON, (b) towers, (c) aircraft with in-situ measurements, (d) aircraft with Programmable Flask Package (PFP) measurements, (e) Eddy covariance measurement stations from AmeriFlux network, (f) Eddy covariance measurement stations from ONEFlux (Open Network-Enabled Flux processing pipeline).

The Final (FNL) Operational Global Analysis data [43] with  $1^\circ \times 1^\circ$  spatial resolution and 6-h interval from National Centers for Environmental Prediction (NCEP), were used as the meteorological initial and boundary conditions of the model. The initial and boundary conditions of CO<sub>2</sub> concentrations were interpolated from the CO<sub>2</sub> mole fraction product of the NOAA’s CarbonTracker, version 2017 (CT2017) [10]. Prior CO<sub>2</sub> fluxes used by CarbonTracker in CT2017, including anthropogenic emissions, fire emissions, biogenic fluxes, and ocean fluxes were used as the prior fluxes input to the model.

**Table 1.** WRF-Chem configurations.

Model Options	Configurations
Domain center	34.939 °N, −96.275 °W
Grid resolution	50 km
nx, ny, nz	103, 82, 45
Time step	240 s
Microphysics process	WSM 5-class simple ice scheme
Cumulus parameterization	Kain-Fritsch scheme
Longwave atmospheric radiation	RRTM scheme
Shortwave atmospheric radiation	Dudhia scheme
Planetary boundary layer scheme	MYNN 2.5 level TKE
Surface layer scheme	MYNN
Land surface scheme	Unified Noah Land surface model
Chemical option	chem_opt = 16 (CO <sub>2</sub> only)

### 2.2. OCO-2 XCO<sub>2</sub> Retrievals

OCO-2 is an atmospheric CO<sub>2</sub> observing satellite launched by the National Aeronautics and Space Administration (NASA) in 2014 [44,45]. OCO-2 is routinely making around 1 million observations each day, and more than 10% of these observations are sufficiently cloud-free to allow for the retrieval of XCO<sub>2</sub> via an inversion algorithm [46]. The XCO<sub>2</sub> retrieved from OCO-2 observations have a median difference of less than 0.5 ppm and root mean square differences typically below 1.5 ppm compared with Total Carbon Column Observing Network (TCCON) [47].

The Version 9r of Level 2 Lite products of the OCO-2 XCO<sub>2</sub> retrievals [48,49] were used in this study. In order to be consistent with the OCO-2 MIP [32], in the lite products, only the retrievals from nadir observation mode with good quantity based on the “xco2\_quality\_flag” were selected. Rather than assimilating each XCO<sub>2</sub> retrieval falling inside a model grid cell separately, a representative mean XCO<sub>2</sub> retrieval and its associated uncertainty of a model grid cell were calculated and assimilated according to the strategy of Crowell et al. [32].

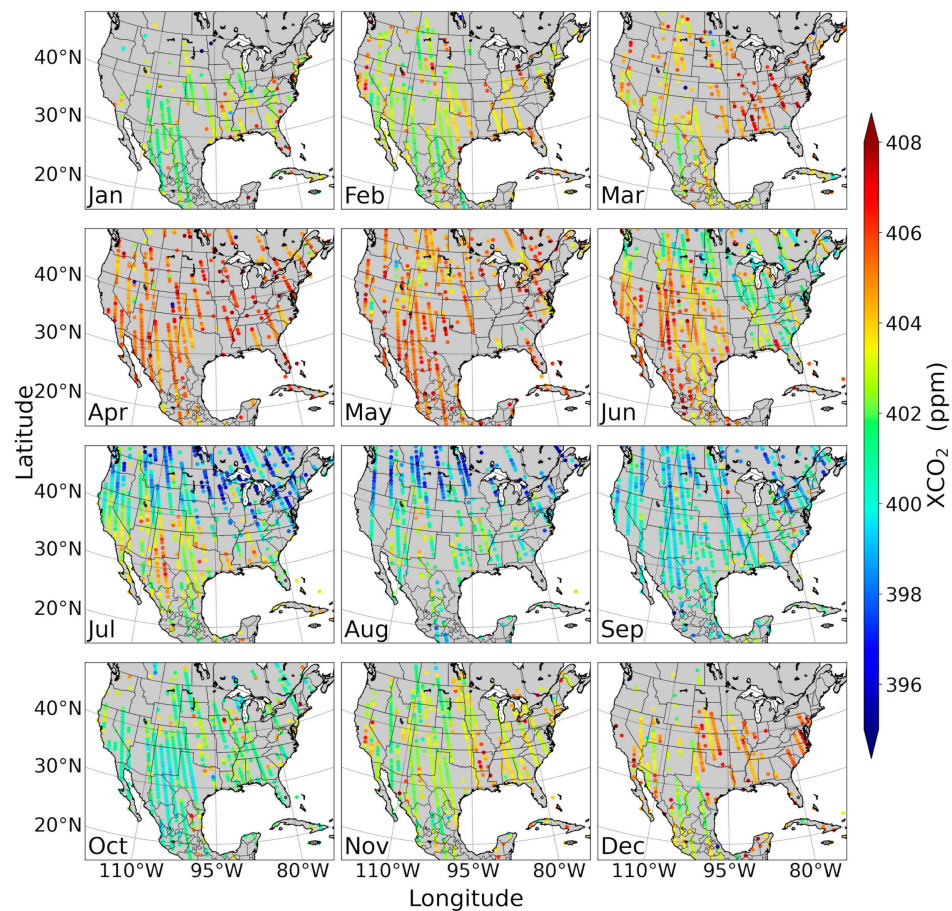
The OCO-2 XCO<sub>2</sub> retrievals over the contiguous United States in each month of 2016 were processed by this strategy and the results are shown in Figure 2.

### 2.3. Observation Operator

The CO<sub>2</sub> concentrations forecasted by WRF-Chem in the model space were converted into XCO<sub>2</sub> values in the OCO-2 observation space by an observation operator adopted from Connor et al. [50]. The observation operator  $H$  was defined as

$$XCO_2^m = XCO_2^b + \sum_j h_j a_{CO_2,j} (x^m - x^b)_j \quad (1)$$

where  $XCO_2^b$ ,  $h$ ,  $a_{CO_2}$ , and  $x^b$  are the prior XCO<sub>2</sub> value, the pressure weighting function, the column averaging kernel, and the prior CO<sub>2</sub> concentration profile used by OCO-2 XCO<sub>2</sub> retrieval processes, respectively. All these parameters are provided in the OCO-2 Level 2 Lite products and require pretreatments similar to the representative mean XCO<sub>2</sub> retrievals.  $x^m$  is the vertical profile of CO<sub>2</sub> concentrations interpolated from WRF-Chem forecasts to the pressure levels of the OCO-2 XCO<sub>2</sub> retrieval.  $XCO_2^m$  is the converted result of the WRF-Chem forecasts by the observation operator  $H$ .



**Figure 2.** Representative mean XCO<sub>2</sub> retrievals of OCO-2 in each month of 2016 over the contiguous United States.

#### 2.4. Regional CO<sub>2</sub> Flux Inversion System

Zhang et al. [38] had implemented the assimilation of the OCO-2 XCO<sub>2</sub> retrievals by extending DART [40] to improve the estimation of regional CO<sub>2</sub> concentrations. DART [40] is a widely used open-source software framework for ensemble data assimilation research [51–54].

On the basis of Zhang’s assimilation system [38], a regional flux inversion system had been further developed in this study by augmenting the state vector to include the CO<sub>2</sub> fluxes [55,56]. The state vector  $x$  was defined as

$$x = [CC, CF]^T \quad (2)$$

where  $CC$  represent the CO<sub>2</sub> concentration states in the model space, which was advanced by WRF-Chem and  $CF$  represent the CO<sub>2</sub> fluxes input into WRF-Chem.  $CF$  was the non-observable part of the state vector and could not be forwarded by WRF-Chem model.

The initial value of  $CC$  for the  $k$ -th ensemble member, defined as  $CC_k^{init}$ , was generated in the same way as that of Zhang, et al. [38] by imposing zero-mean Gaussian distributions to the average CO<sub>2</sub> concentrations interpolated from CT2017 [10]. The standard deviations of these Gaussian distributions were assumed to be 1.5% of the average CO<sub>2</sub> concentrations.

The update process of  $CC$  was the same as that of Zhang et al. [38]. Let  $CC_{k,i}^f$  and  $CC_{k,i}^a$  represent the forecast of  $CC$  advanced by WRF-Chem starting from  $CC_k^{init}$  and the analysis result of  $CC$  for the  $k$ -th ensemble member at time  $i$  ( $i \geq 1$ ), respectively. The  $CC_{k,i}^f$  was converted into the observation space by the observation operator  $H$  defined

by Equation (3), i.e.,  $y_{k,i}^f = H(CC_{k,i}^f)$ . Then the analysis result  $y_{k,i}^a$  of the  $k$ -th ensemble member at time  $i$  was calculated by

$$y_{k,i}^a = \left[ \sqrt{\frac{\widetilde{\sigma_{XCO_2}^2}}{\widetilde{\sigma_{XCO_2}^2} + \bar{\sigma}_i^{f,2}}} \right] \left[ y_{k,i}^f - \bar{y}_i^f \right] + \left[ \frac{\bar{y}_i^f}{\bar{\sigma}_i^{f,2}} + \frac{\widetilde{XCO_2}}{\widetilde{\sigma_{XCO_2}^2}} \right] \left[ \frac{1}{\bar{\sigma}_i^{f,2}} + \frac{1}{\widetilde{\sigma_{XCO_2}^2}} \right]^{-1} \quad (3)$$

where  $\bar{y}_i^f$  and  $\bar{\sigma}_i^f$  were the forecast mean and the spread of all ensemble members in the observation space at time  $i$  ( $i \geq 1$ ), respectively.  $\widetilde{XCO_2}$  and  $\widetilde{\sigma_{XCO_2}^2}$  were the representative mean  $XCO_2$  retrieval and its associated uncertainty of a model grid cell calculated from the OCO-2 retrievals according to the strategy of Crowell et al. [32].

The analysis result of  $CC_{k,i}^f$  was updated as

$$CC_{k,i}^a = CC_{k,i}^f + \alpha \frac{\sigma(CC_i^f, y_i^f)}{\bar{\sigma}_i^{f,2}} (y_{k,i}^a - y_{k,i}^f) \quad (4)$$

where  $\alpha$  was a covariance localization coefficient to compensate the sampling error caused by the limited ensemble size [57],  $\sigma(CC_i^f, y_i^f)$  is the covariance of the forecasts of  $CC$  and their corresponding conversion results in the observation space across all ensemble members at time  $i$  ( $i \geq 1$ ).

The initial value of  $CF$  for the  $k$ -th ensemble member at time  $i$  ( $i \geq 1$ ), defined as  $CF_{k,i-1}^{init}$ , was generated by imposing zero-mean Gaussian distributions to the average  $CO_2$  fluxes interpolated from the prior  $CO_2$  fluxes used by CT2017 [10]. The standard deviations of these flux Gaussian distributions were assumed to be 50% of the average  $CO_2$  fluxes similar to Chevallier et al. [30].

Let  $CF_{k,i-1}^{prior}$  and  $CF_{k,i-1}^{post}$  represent the prior and posterior  $CO_2$  fluxes of the model grids for the  $k$ -th ensemble member at time ( $i \geq 1$ ), respectively. Let  $\lambda_{k,i-1}^{prior}$  and  $\lambda_{k,i-1}^{post}$  represent the prior and posterior scaling factors between the pairs of  $CF_{k,i-1}^{prior}$ ,  $CF_{k,i-1}^{init}$  and  $CF_{k,i-1}^{post}$ ,  $CF_{k,i-1}^{prior}$ , respectively. The  $CF_{k,i-1}^{prior}$  was calculated as

$$CF_{k,i-1}^{prior} = \lambda_{k,i-1}^{prior} \times CF_{k,i-1}^{init}, i \geq 1 \quad (5)$$

where the prior scaling factor  $\lambda_{k,i-1}^{prior}$  was evolved forward with time  $i$  by

$$\lambda_{k,i-1}^{prior} = \begin{cases} 1, & i = 1 \\ \lambda_{k,i-2}^{post}, & i \geq 2 \end{cases} \quad (6)$$

The posterior  $CO_2$  flux and scaling factor were updated by

$$CF_{k,i-1}^{post} = CF_{k,i-1}^{prior} + \alpha \frac{\sigma(CF_{i-1}^{prior}, y_i^f)}{\bar{\sigma}_i^{f,2}} (y_{k,i}^a - y_{k,i}^f), i \geq 1 \quad (7)$$

$$\lambda_{k,i-1}^{post} = CF_{k,i-1}^{post} / CF_{k,i-1}^{prior}, i \geq 1 \quad (8)$$

where  $\alpha$  and  $\sigma(CF_{i-1}^{prior}, y_i^f)$  have the similar definitions as those of in the Equation (4).

## 2.5. Experiment Design

Two experiments were conducted to evaluate the performance of the flux inversion system. The first one, named "DA\_FLUX", used the flux inversion system to optimize both the atmospheric  $CO_2$  concentrations and  $CO_2$  fluxes during the assimilation process. The

second experiment, named “SIM”, only used WRF-Chem to simulate the atmospheric CO<sub>2</sub> concentrations and did not assimilate any OCO-2 XCO<sub>2</sub> retrievals.

#### 2.5.1. The DA\_FLUX Experiment

In the DA\_FLUX experiment, the CO<sub>2</sub> concentrations and CO<sub>2</sub> fluxes over the contiguous United States in 2016 were estimated by the flux inversion system through the assimilation of OCO-2 XCO<sub>2</sub> retrievals.

The ensemble size of the system was 20. The initial and boundary conditions of CO<sub>2</sub> concentrations were interpolated from CT2017 [10]. A strategy similar to CarbonTracker was taken to quantify the uncertainty due to prior flux selection ([https://www.esrl.noaa.gov/gmd/ccgg/carbontracker/CT2019B\\_doc.php](https://www.esrl.noaa.gov/gmd/ccgg/carbontracker/CT2019B_doc.php), last access: 18 June 2021), by interpolating two sets of prior fluxes from two different combinations of terrestrial fluxes, wildfire emissions, air–sea CO<sub>2</sub> exchanges, and fossil fuel emissions from CT2017. The mean posterior CO<sub>2</sub> fluxes of two inversions, each using one set of prior fluxes, were the final CO<sub>2</sub> flux results of the inversion system. The final CO<sub>2</sub> concentration results were estimated by applying the mean posterior fluxes to the forward model WRF-Chem once again and without assimilation.

The study period was the entire year of 2016. The period of five weeks before 1 January 2016 was used as the spin-up phase of the simulation. The simulation cycle was 6-h at 00:00, 06:00, 12:00, and 18:00 UTC, respectively. A short assimilation window of one day was used according to Liu et al. [58].

#### 2.5.2. The SIM Experiment

The SIM experiment was a control experiment, using only WRF-Chem to simulate the CO<sub>2</sub> concentrations, and not assimilating any OCO-2 XCO<sub>2</sub> retrievals. The prior CO<sub>2</sub> flux of the SIM experiment was the average of the two sets of prior fluxes used in the DA\_FLUX experiment. All other initial conditions and configurations were the same as those of the DA\_FLUX experiment.

### 2.6. Evaluation

The level-4 flux product version 7 (<https://www.esrl.noaa.gov/gmd/ccgg/OCO2/>, last access: 18 June 2021) from the OCO-2 model intercomparison project (OCO-2 MIP) was used to compare with the posterior CO<sub>2</sub> flux obtained by the regional flux inversion system. The OCO-2 MIP [32] is a collaboration organized by the OCO-2 Science Team to study the impact of assimilating OCO-2 XCO<sub>2</sub> retrievals into atmospheric inversion models. Currently, there are nine models in the OCO-2 MIP version 7, they differ in the sources of prior CO<sub>2</sub> flux, the types of atmospheric chemical transport model, the methods of data assimilation, and so on. The detailed information of these models is listed in Table S1.

Independent ground and aircraft observations were used to evaluate the performance of the regional flux inversion system. The CO<sub>2</sub> flux observations from AmeriFlux [59] and ONEFlux [60], including 18 and 10 eddy covariance flux measurement sites respectively, were used to evaluate the CO<sub>2</sub> flux results of the DA\_FLUX experiment. The detailed information of these flux observation sites is listed in Tables S2 and S3, and their locations are shown in Figure 1e,f, respectively.

The XCO<sub>2</sub> retrievals of four TCCON [61] sites, including Park Falls [62], Lamont [63], Edwards [64], and Caltech [65], CO<sub>2</sub> concentration observations of 10 tower sites and 20 aircraft campaigns from the GLOBALVIEWplus project [66] with Observation Package (ObsPack) [67] data products “obspack\_co2\_1\_GLOBALVIEWplus\_v5.0\_2019-08-12” were used to evaluate the CO<sub>2</sub> concentration results of the two experiments. The detailed information of these sites and campaigns is listed in Tables S4 and S5, and their locations are shown in Figure 1a,d.

The posterior CO<sub>2</sub> concentrations and posterior fluxes from CT2017 were also compared with the independent observations to serve as references for the evaluation.

The experimental results were evaluated by mean bias error (MBE), root-mean-square error (RMSE), and correlation coefficient (CORR), same as the Equations (6), (8), and (9) in Zhang's work [38].

The uncertainty reduction between the posterior and prior CO<sub>2</sub> fluxes is a useful method to evaluate the impact of assimilation of OCO-2 XCO<sub>2</sub> retrievals in constraining CO<sub>2</sub> flux. The uncertainty reduction achieved by the flux inversion system was evaluated by the equation [68]

$$UR = \left(1 - \frac{\sigma_{post}}{\sigma_{prior}}\right) \times 100\% \quad (9)$$

where  $\sigma_{post}$  and  $\sigma_{prior}$  were the standard deviations of the posterior and prior CO<sub>2</sub> fluxes, respectively

### 3. Results and Discussion

#### 3.1. Regional CO<sub>2</sub> Flux Inversion Results

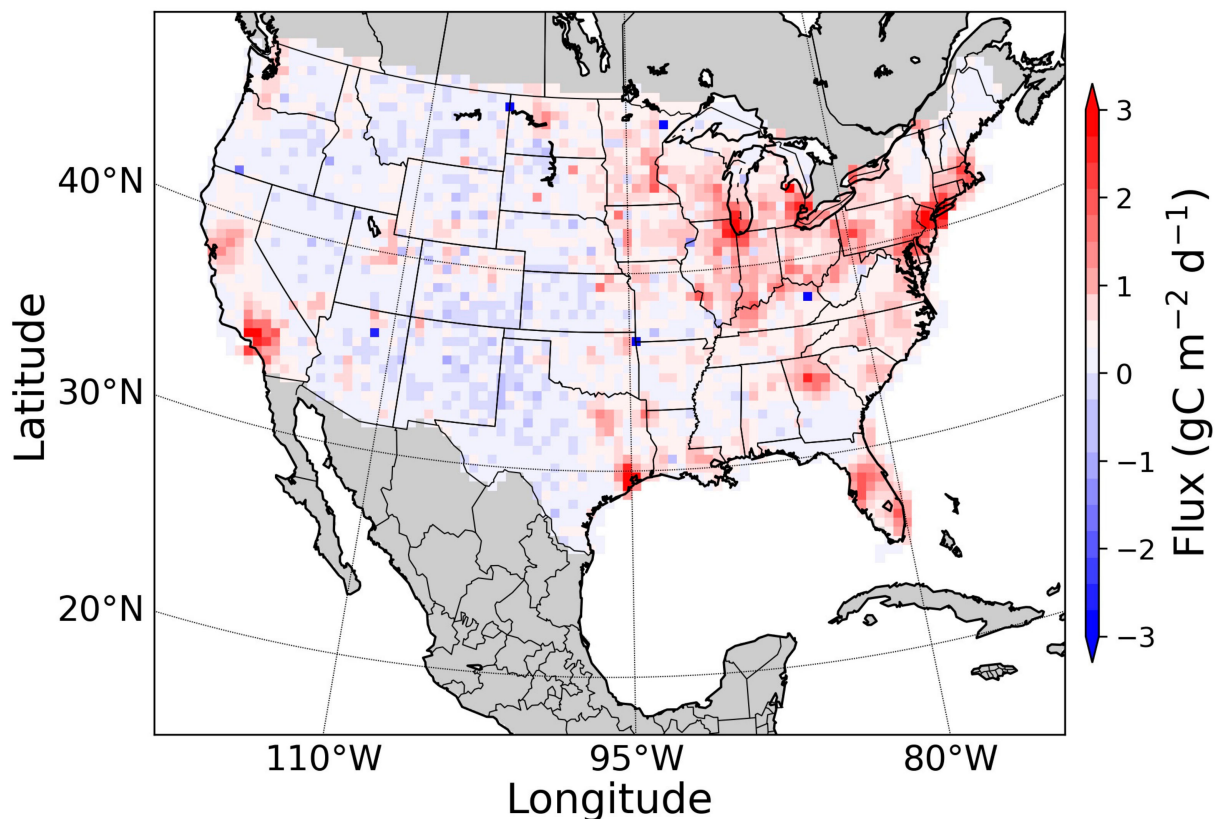
By assimilating the OCO-2 XCO<sub>2</sub> retrievals, the regional CO<sub>2</sub> flux inversion system in the DA\_FLUX experiment optimized the annual mean CO<sub>2</sub> flux over the contiguous United States in 2016 from the prior flux of  $1.51 \pm 0.11$  Pg C yr<sup>-1</sup> to the posterior flux of  $1.08 \pm 0.03$  Pg C yr<sup>-1</sup>, which was more in line with the fluxes estimated by CT2017 and the models from OCO-2 MIP. The prior and posterior monthly mean CO<sub>2</sub> fluxes of the DA\_FLUX experiment are listed in Table 2. The results of DA\_FLUX indicated that in 2016 the contiguous United States was a carbon sink during the four-month period from May to August due to the vegetation growing season, and was a carbon source in the remaining eight months.

**Table 2.** Prior and posterior monthly mean CO<sub>2</sub> fluxes of the DA\_FLUX experiment over the contiguous United States in 2016.

Month	Prior Fluxes (Pg C)	Posterior Fluxes (Pg C)
1	0.35 ± 0.03	0.33 ± 0.01
2	0.32 ± 0.008	0.27 ± 0.001
3	0.28 ± 0.02	0.19 ± 0.02
4	0.12 ± 0.05	0.05 ± 0.03
5	−0.05 ± 0.06	−0.1 ± 0.05
6	−0.21 ± 0.05	−0.17 ± 0.05
7	−0.20 ± 0.05	−0.07 ± 0.02
8	−0.08 ± 0.002	−0.02 ± 0.001
9	0.04 ± 0.03	0.07 ± 0.02
10	0.21 ± 0.03	0.16 ± 0.03
11	0.35 ± 0.01	0.19 ± 0.03
12	0.37 ± 0.02	0.19 ± 0.04
Annual	1.51 ± 0.11	1.08 ± 0.03

The spatial distribution of the annual mean posterior CO<sub>2</sub> flux of DA\_FLUX, shown in Figure 3, indicates the areas with higher CO<sub>2</sub> fluxes roughly coincided with the densely populated urban areas—i.e., the metropolis areas of New York, New Jersey, Detroit of Michigan, Chicago of Illinois, Atlanta of Georgia, Tampa and Orlando of Florida, Miami of Florida, Houston of Texas, Los Angeles, and San Francisco of California all showed up as CO<sub>2</sub> sources of varying magnitude. It implies the majority of the CO<sub>2</sub> emission in the contiguous United States comes from the entire northeast, midwest, and a few hotspots in the south and west because these are the largest metropolitan/urban areas. Figure 3 also indicates almost the entire Great Plains and Rocky Mountain West were a weak CO<sub>2</sub> sink because with a few exceptions there are virtually no large cities. The spatial distributions of the monthly mean posterior fluxes are shown in Figure S1.





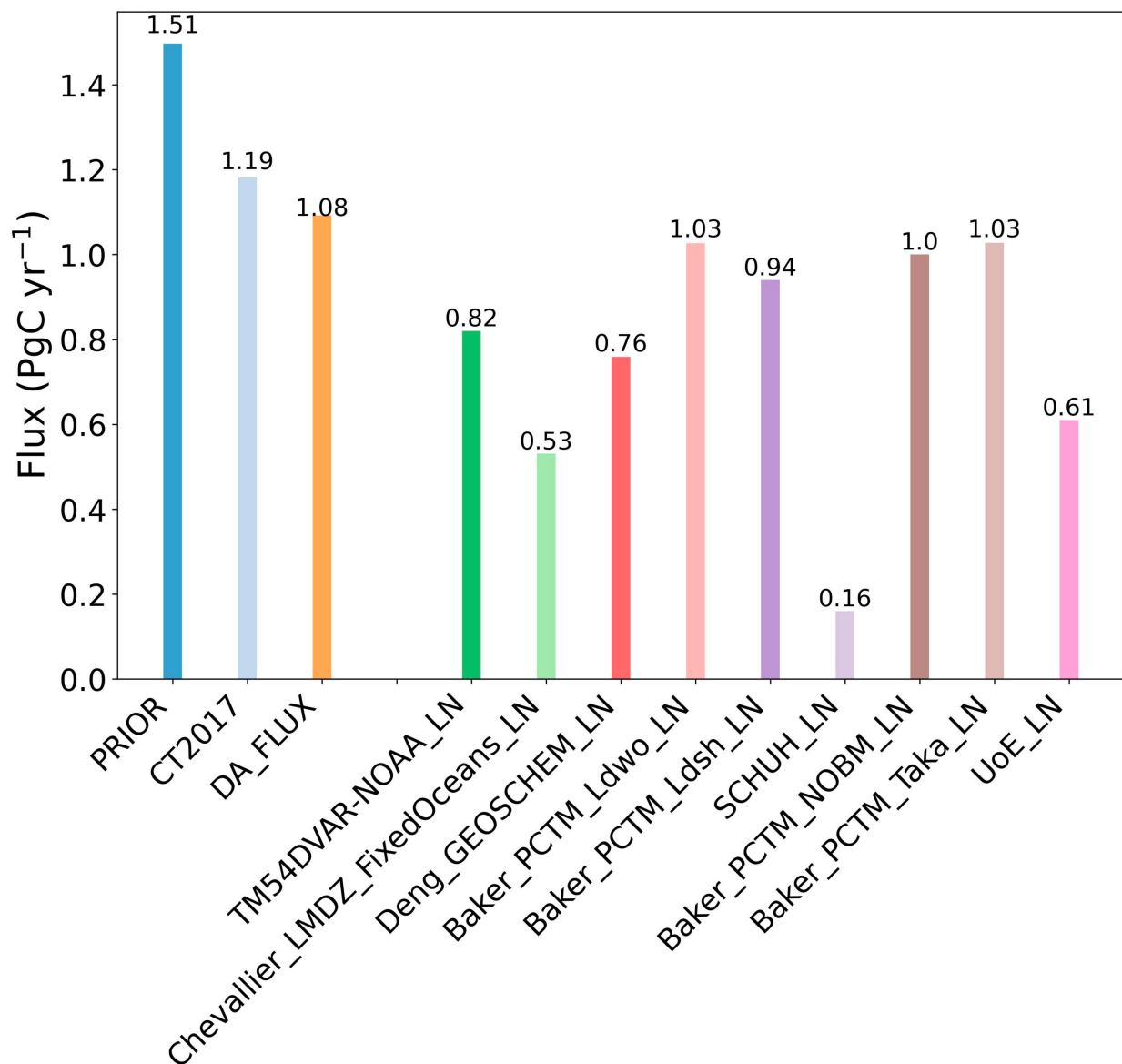
**Figure 3.** Spatial distribution of the annual mean posterior CO<sub>2</sub> flux estimated by the DA\_FLUX experiment over the contiguous United States in 2016.

### 3.1.1. Compared with CT2017 and OCO-2 MIP Models

Figures 4 and 5 show the comparison of the annual and monthly posterior CO<sub>2</sub> flux of the DA\_FLUX experiment with the CO<sub>2</sub> fluxes estimated by CT2017 and OCO-2 MIP models, and with the prior flux used in the experiment, respectively. The detailed information of the OCO-2 MIP models shown in Figures 4 and 5 is listed in Table S1. The monthly mean values of each flux in Figure 5 are listed in Table S6.

The flux of 1.19 Pg C yr<sup>-1</sup> of CT2017 was extracted from the posterior CO<sub>2</sub> flux of CT2017 in 2016 according to the border of the contiguous United States. The fluxes of OCO-2 MIP models were extracted from the “TransCom 02” region of the OCO-2 MIP level 4 flux products, according to the area ratio of the contiguous United States and the “TransCom 02” region. The “TransCom 02” region covers the North American temperate zone, which is slightly larger than the contiguous United States. The mean flux and standard deviation of all the nine models from OCO-2 MIP were 0.76 Pg C yr<sup>-1</sup> and 0.29 Pg C yr<sup>-1</sup>, respectively.

Figure 4 shows the annual posterior CO<sub>2</sub> flux of the DA\_FLUX experiment was 0.11 Pg C yr<sup>-1</sup> lower than that of CT2017, and was 0.05 Pg C yr<sup>-1</sup> higher than the highest flux from OCO-2 MIP estimated by the model of Baker\_PCTM\_Ldwo\_LN and Baker\_PCTM\_Taka\_LN.

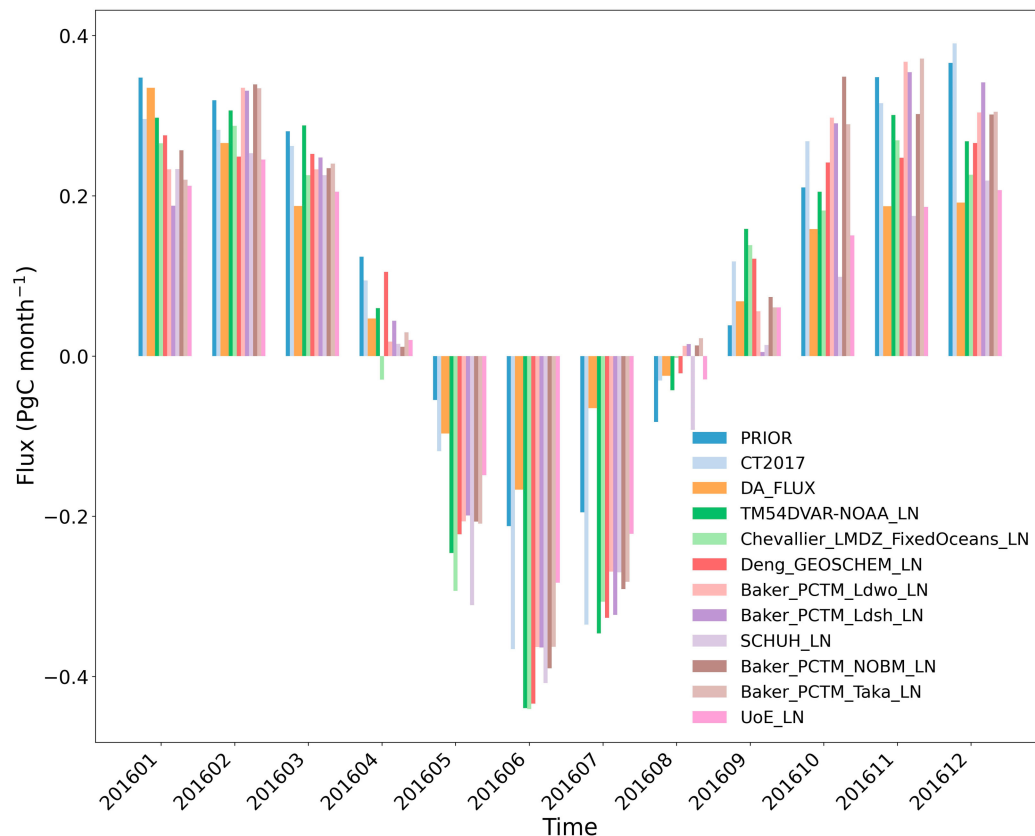


**Figure 4.** Comparison of the annual posterior CO<sub>2</sub> flux over the contiguous United States in 2016 estimated by the DA\_FLUX experiment with those estimated by CT2017 and the models from OCO-2 MIP, and with the prior CO<sub>2</sub> flux used in the DA\_FLUX experiment.

Figure 5 shows that the monthly mean fluxes of DA\_FLUX from May to July, when the contiguous United States was a carbon sink, were significantly smaller than that of OCO-2 MIP models, i.e., less carbon uptake. This explains why the annual flux of DA\_FLUX was higher than that of OCO-2 MIP models. By analyzing the prior flux, the CO<sub>2</sub> concentrations simulated by WRF-Chem before assimilation and the value of representative mean XCO<sub>2</sub> retrievals of OCO-2, it was found that the small posterior fluxes in May and June were mainly caused by the significantly smaller prior fluxes which implied less CO<sub>2</sub> uptake compared with those of OCO-2 MIP models. The prior fluxes in May and June were only 24.3% and 54.8% of the mean fluxes of OCO-2 MIP models, respectively. The small posterior flux in July was caused by the combination of the relatively high representative mean XCO<sub>2</sub> retrievals of OCO-2 and a small prior flux. The average value of the representative mean XCO<sub>2</sub> retrievals in July was 401.21 ppm and was 0.35 ppm higher than the monthly mean CO<sub>2</sub> concentrations simulated by WRF-Chem before assimilation. The prior flux in July was 66.5% of the mean fluxes of OCO-2 MIP models. Figure 5 also shows CT2017 and

OCO-2 MIP models were inconsistent in terms of whether the contiguous United States were the sink or source of CO<sub>2</sub> in the two months of April and August. On this issue, the results of the DA\_FLUX experiment were consistent with the majority.

Based on the above annual and monthly comparison results with other models, the CO<sub>2</sub> flux estimated by the inversion system presented here can be considered reasonable.

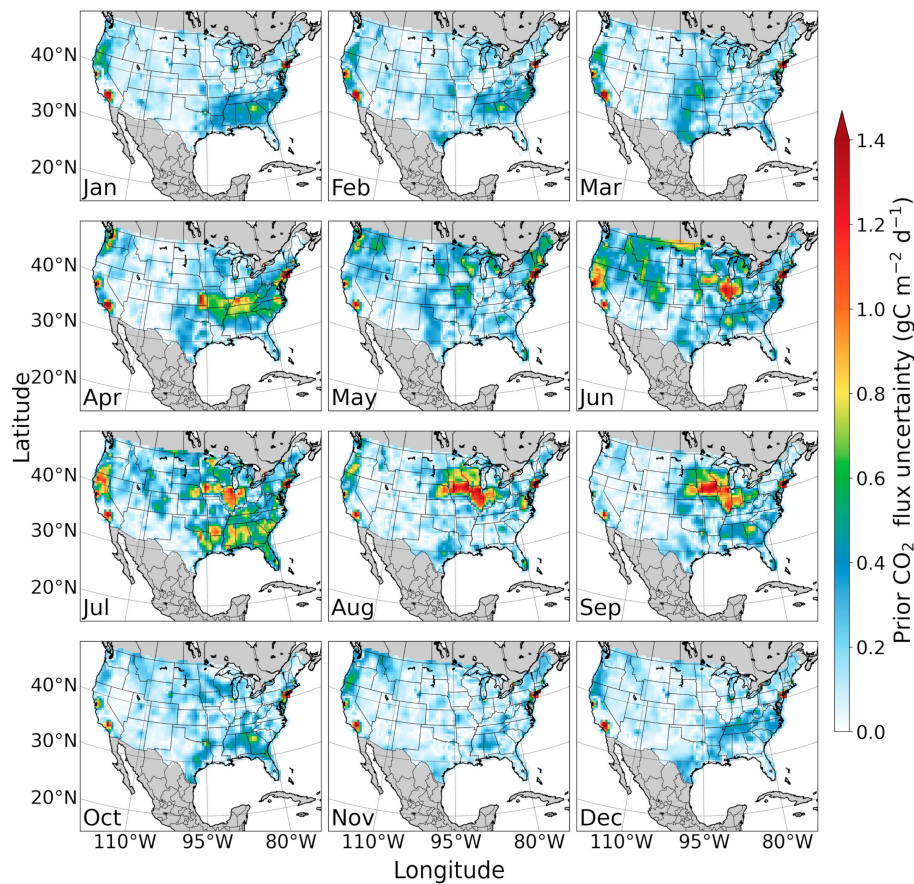


**Figure 5.** Comparison of the monthly posterior CO<sub>2</sub> flux over the contiguous United States in 2016 estimated by the DA\_FLUX experiment with those estimated by CT2017 and the models from OCO-2 MIP, and with the prior CO<sub>2</sub> flux used in the DA\_FLUX experiment.

### 3.1.2. Uncertainty Reduction by the Inversion System

The uncertainty of the inverted annual CO<sub>2</sub> flux in the DA\_FLUX experiment was reduced by 14.71% compared to the prior flux.

Figures 2 and 6 show the spatial distribution of the representative mean XCO<sub>2</sub> retrievals of OCO-2 and the mean uncertainty of the prior CO<sub>2</sub> flux of each month in 2016 used in the DA\_FLUX experiment, respectively. The monthly mean uncertainty of the prior CO<sub>2</sub> flux, the monthly mean uncertainty reduction percentage over all grid cells, and the number of the representative mean XCO<sub>2</sub> retrievals of OCO-2 of each month in 2016 are listed in Error! Reference source not found. Figure 7 shows the spatial distribution of monthly mean uncertainty reduction between the posterior and the prior CO<sub>2</sub> flux in the DA\_FLUX experiment. In Figure 7, areas with negative uncertainty reduction were masked as zero, since such uncertainty increases simply resulted from the small number of realizations [35].

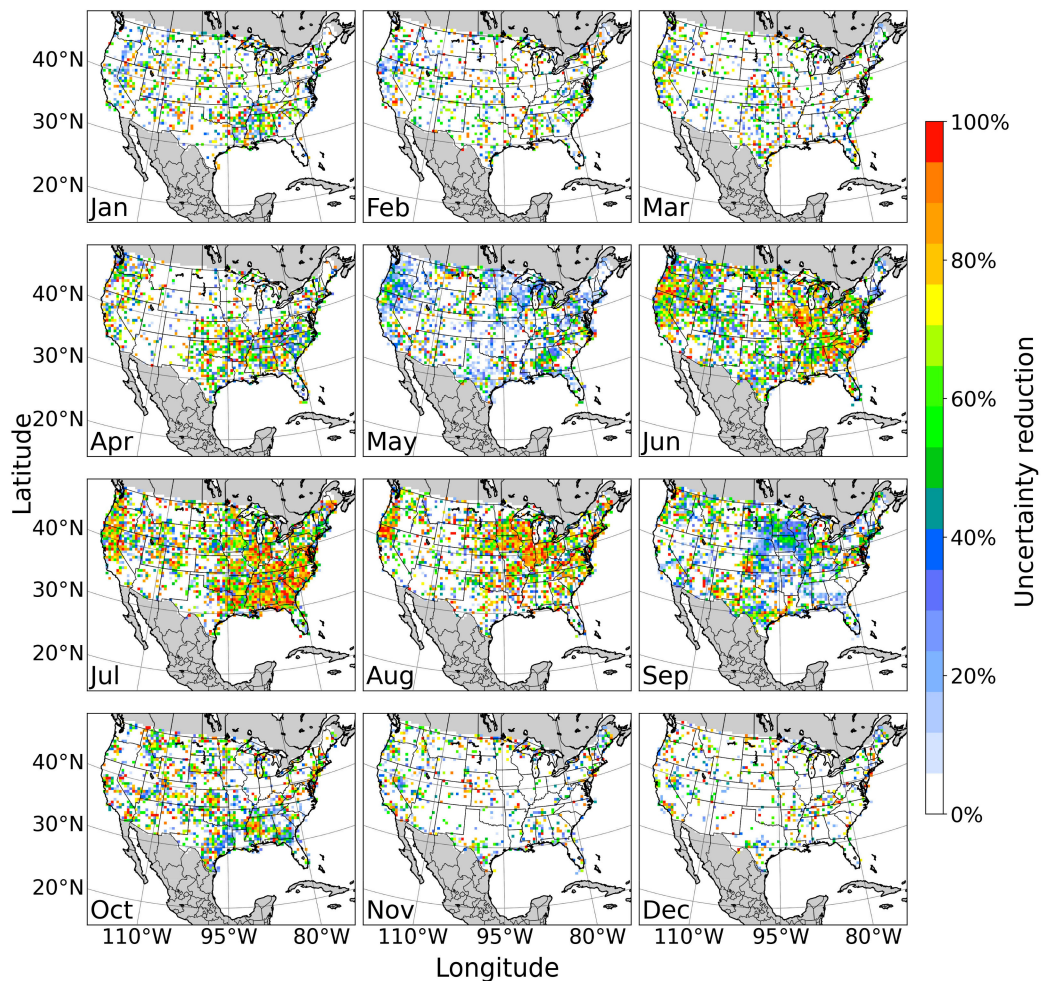


**Figure 6.** Spatial distribution of monthly mean uncertainty of the prior CO<sub>2</sub> flux over the contiguous United States in 2016 used in the DA\_FLUX experiment.

From Table 3, the uncertainty reduction had a strong positive correlation with the uncertainty of the prior CO<sub>2</sub> flux and the number of the representative mean XCO<sub>2</sub> retrievals of OCO-2. The correlations between the uncertainty reduction and the prior flux uncertainty, the number of the representative mean XCO<sub>2</sub> retrievals were 0.93 and 0.72, respectively. The largest uncertainty reductions of 36.17% and 38.42% occurred in June and July, which coincided with the period that had the largest prior flux uncertainties of 0.31 g C m<sup>-2</sup> d<sup>-1</sup> and 0.32 g C m<sup>-2</sup> d<sup>-1</sup>, respectively. June and July are the middle of the growing season in the contiguous United States, and their mean prior flux uncertainties over all the grid cells are driven up by vegetation photosynthesis.

**Table 3.** Monthly mean uncertainty of the prior CO<sub>2</sub> flux, the monthly mean uncertainty reduction percentage over all grid cells and the number of the representative mean XCO<sub>2</sub> retrievals of OCO-2 of each month in 2016.

Month in 2016	Percentage of Mean Uncertainty Reduction over All Grid Cells	Number of Representative Mean XCO <sub>2</sub> Retrievals of OCO-2	Mean Uncertainty of the Prior CO <sub>2</sub> Flux over All Grid Cells (g C m <sup>-2</sup> d <sup>-1</sup> )
1	14.82%	303	0.16
2	12.91%	458	0.16
3	12.80%	347	0.17
4	19.59%	449	0.22
5	18.30%	504	0.22
6	36.17%	695	0.31
7	38.42%	727	0.32
8	30.18%	429	0.22
9	24.42%	620	0.21
10	21.72%	575	0.17
11	10.33%	566	0.14
12	9.53%	344	0.14
Annual	14.71%	6017	0.085



**Figure 7.** Spatial distribution of monthly mean uncertainty reduction between the posterior and the prior CO<sub>2</sub> flux in 2016 in the DA\_FLUX experiment.

In the spatial dimension, except for September, the regions with larger uncertainty reduction in Figure 7 roughly coincided with the locations with higher prior uncertainty in Figure 6. In August and September, Error! Reference source not found. shows that the areas, mainly near the southwest of the Great Lakes of North America, had similar distributions of high prior uncertainty. This “upper midwest” region of the contiguous United States is also referred to as the “corn belt”, it is a known mega-sink of CO<sub>2</sub> during the agricultural growing season [69]. Figure 7 shows this region had a larger uncertainty reduction in August, which was in line with the expectation, while it only had a limited uncertainty reduction in September. By comparing the distribution of OCO-2 representative XCO<sub>2</sub> retrievals in August and September shown in Figure 2, it was found that there were no OCO-2 XCO<sub>2</sub> retrievals in this “upper midwest” region in September. This explains why the region had a limited uncertainty reduction in September, and also demonstrated the impact of OCO-2 XCO<sub>2</sub> retrievals on the flux constraints.

### 3.1.3. Compared with CO<sub>2</sub> Flux Measurements

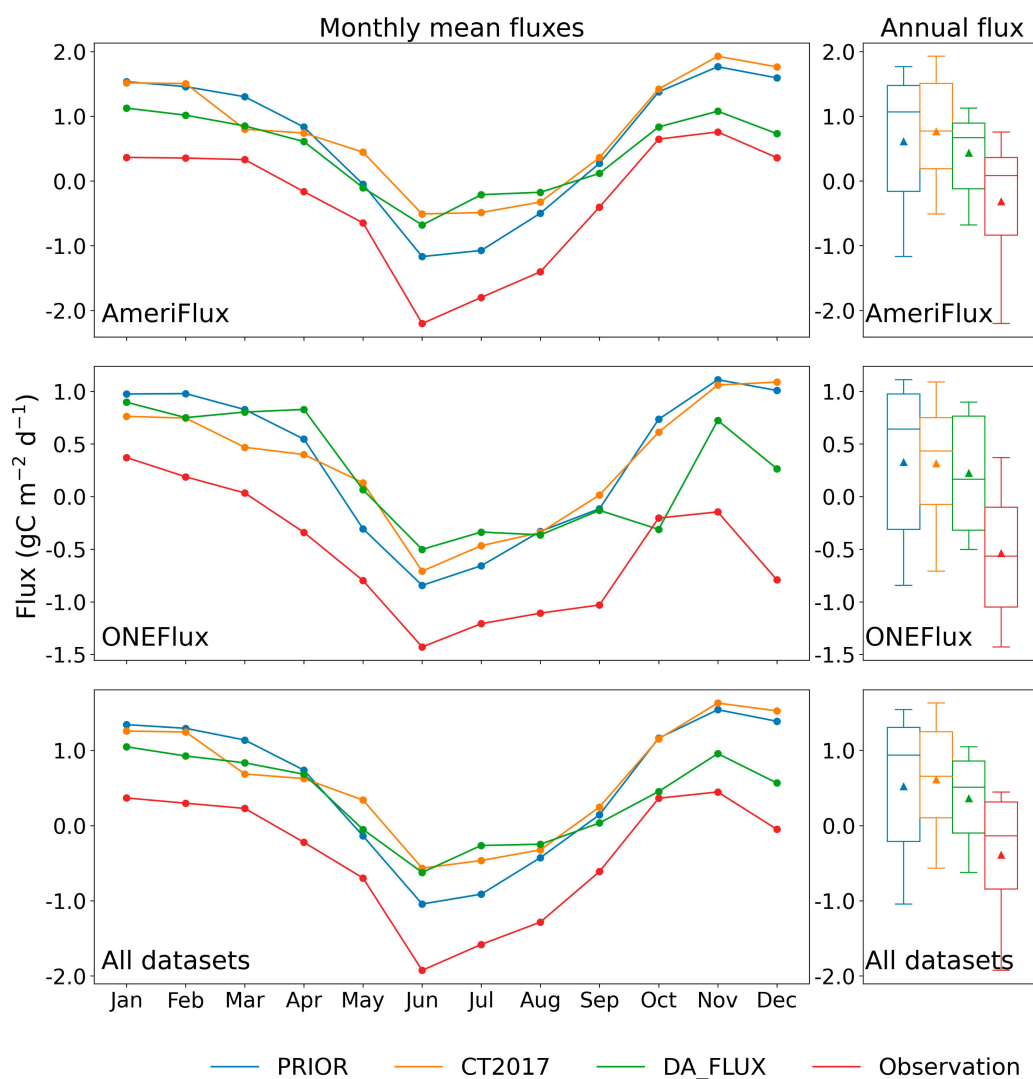
Although the spatial representativeness of the CO<sub>2</sub> flux tower measurements is usually less than 1 km<sup>2</sup>, and the grid cell size in this study was 50 km, meaning that there was some mismatch of the spatial scale, the statistical results of comparing the CO<sub>2</sub> flux measurements with the DA\_FLUX experimental results were still meaningful.

The annual mean of the prior, posterior CO<sub>2</sub> flux of DA\_FLUX and the posterior flux of CT2017 interpolated to the locations of flux measurement sites are listed in Table 4, and

box-plotted on the right side of Figure 8. The overall annual mean flux measurement of AmeriFlux [59] and ONEFlux [60] was  $-0.38 \text{ g C m}^{-2} \text{ d}^{-1}$ . The corresponding annual mean posterior flux of DA\_FLUX was  $0.35 \text{ g C m}^{-2} \text{ d}^{-1}$ , which were  $0.19 \text{ g C m}^{-2} \text{ d}^{-1}$  and  $0.25 \text{ g C m}^{-2} \text{ d}^{-1}$  smaller than that of prior flux and CT2017, respectively, and closer to the measurements.

**Table 4.** Annual mean value of the prior (PRIOR), posterior (DA\_FLUX) CO<sub>2</sub> flux of the DA\_FLUX experiment, the posterior CO<sub>2</sub> flux of CT2017 interpolated to the locations of flux measurement sites, and annual mean CO<sub>2</sub> flux measurements of AmeriFlux and ONEFlux.

Dataset	Flux Measurements ( $\text{g C m}^{-2} \text{ d}^{-1}$ )	Flux		
		PRIOR ( $\text{g C m}^{-2} \text{ d}^{-1}$ )	CT2017 ( $\text{g C m}^{-2} \text{ d}^{-1}$ )	DA_FLUX ( $\text{g C m}^{-2} \text{ d}^{-1}$ )
AmeriFlux	-0.25	0.71	0.85	0.46
ONEFlux	-0.52	0.36	0.35	0.24
All	-0.38	0.54	0.60	0.35



**Figure 8.** Monthly and annual mean of the prior (PRIOR), posterior (DA\_FLUX) CO<sub>2</sub> flux of the DA\_FLUX experiment, the posterior CO<sub>2</sub> flux of CT2017, and the CO<sub>2</sub> flux measurements from AmeriFlux and ONEFlux. For the boxplots on the right, the lower and upper edges of the boxes represent the first (Q1) and third (Q3) quartiles, the lower and upper whiskers represent the minimum and maximum values within 1.5 times the interquartile range from Q1 and Q3, respectively, solid lines and triangles represent median and mean value.

The line charts on the left side of Figure 8 show the monthly mean of the prior and posterior CO<sub>2</sub> fluxes of DA\_FLUX, of the posterior CO<sub>2</sub> flux of CT2017, and of the CO<sub>2</sub> flux measurements. The detailed monthly mean CO<sub>2</sub> flux values are listed in Table S7.

The statistical evaluation results of comparing the CO<sub>2</sub> flux measurements from AmeriFlux [59] and ONEFlux [60] with the prior, posterior CO<sub>2</sub> fluxes of the DA\_FLUX experiment, and with the posterior CO<sub>2</sub> flux of CT2017 are listed in Table 5. The RMSE and MBE of the posterior CO<sub>2</sub> flux in DA\_FLUX against all the flux measurements from AmeriFlux and ONEFlux were 0.79 g C m<sup>-2</sup> d<sup>-1</sup> and 0.73 g C m<sup>-2</sup> d<sup>-1</sup>, respectively, which were 0.16 g C m<sup>-2</sup> d<sup>-1</sup> (16.84%) and 0.21 g C m<sup>-2</sup> d<sup>-1</sup> (22.34%) smaller than that of the prior flux and were 0.27 g C m<sup>-2</sup> d<sup>-1</sup> (25.47%) and 0.29 g C m<sup>-2</sup> d<sup>-1</sup> (28.43%) smaller than that of CT2017, respectively.

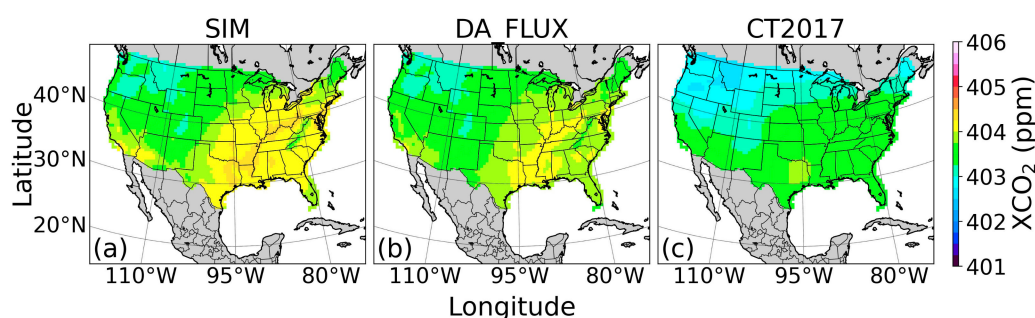
**Table 5.** Statistical results of comparing the CO<sub>2</sub> flux measurements from AmeriFlux and ONEFlux with the prior (PRIOR), posterior (DA\_FLUX) CO<sub>2</sub> flux of the DA\_FLUX experiment, and with the posterior CO<sub>2</sub> flux of CT2017.

Dataset	Flux	RMSE (g C m <sup>-2</sup> d <sup>-1</sup> )	MBE (g C m <sup>-2</sup> d <sup>-1</sup> )	CORR
AmeriFlux	PRIOR	0.98	0.97	0.97
	CT2017	1.15	1.10	0.92
	DA_FLUX	0.80	0.71	0.95
ONEFlux	PRIOR	0.94	0.88	0.85
	CT2017	0.94	0.86	0.75
	DA_FLUX	0.83	0.76	0.83
All dataset	PRIOR	0.95	0.94	0.97
	CT2017	1.06	1.02	0.92
	DA_FLUX	0.79	0.73	0.94

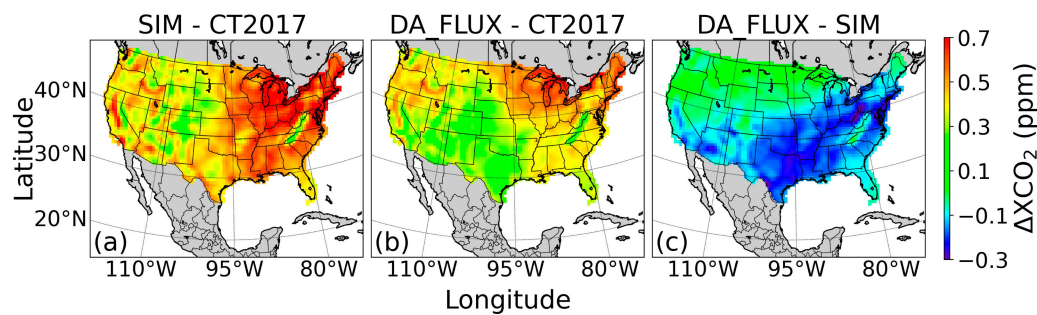
The evaluation results indicated that the posterior CO<sub>2</sub> fluxes of the inversion system were significantly more consistent with the flux measurements than the prior CO<sub>2</sub> fluxes and the posterior CO<sub>2</sub> flux of CT2017.

### 3.2. CO<sub>2</sub> Concentrations Results

The annual mean XCO<sub>2</sub> over the contiguous United States in 2016 estimated by the SIM, DA\_FLUX experiment, and CT2017 were 403.78 ppm, 403.67 ppm, and 403.29 ppm, respectively, and their distributions are shown in Figure 9. The annual mean XCO<sub>2</sub> of DA\_FLUX was 0.11 ppm smaller than that of SIM and closer to CT2017. The differences of annual mean XCO<sub>2</sub> between SIM and CT2017, between DA\_FLUX and CT2017, between DA\_FLUX and SIM were 0.49 ppm, 0.38 ppm, −0.11 ppm, respectively. Their distributions are shown in Figure 10. It is obvious that the result of DA\_FLUX was more in line with CT2017 than that of SIM.

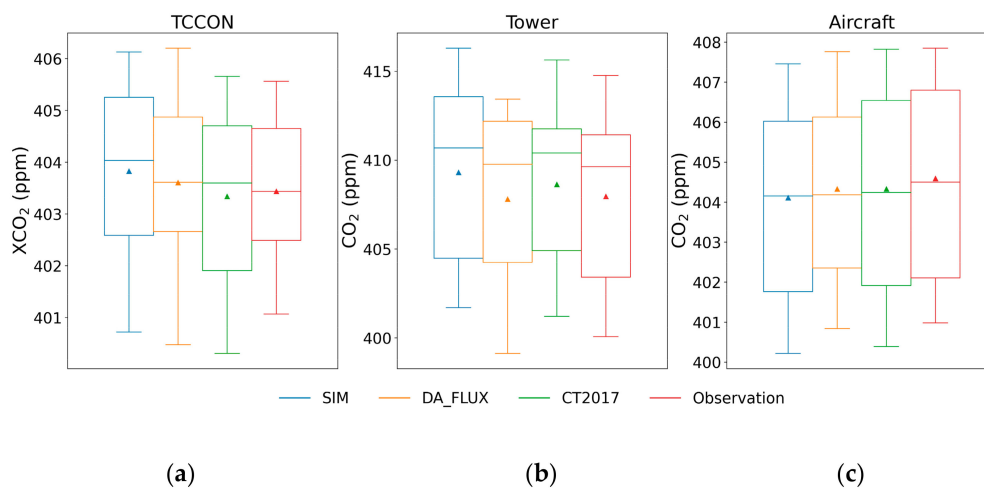


**Figure 9.** Annual mean XCO<sub>2</sub> distribution over the contiguous United States in 2016 from the SIM (a), DA\_FLUX (b) experiments, and CT2017 (c).



**Figure 10.** Difference in the annual mean XCO<sub>2</sub> distribution over the contiguous United States in 2016, (a) between SIM and CT2017, (b) between DA\_FLUX and CT2017, (c) between DA\_FLUX and SIM.

The boxplots of the monthly mean CO<sub>2</sub> concentrations of the SIM, DA\_FLUX experiment and CT2017 compared with monthly mean observations of TCCON, tower, and aircraft are shown in Figure 11, the corresponding data of Figure 11 are listed in Table S8. The detailed evaluation results of statistical metrics are shown in Table 6. The comparison of the monthly mean CO<sub>2</sub> concentrations of the SIM, DA\_FLUX experiment, and CT2017 with monthly mean observations of TCCON, tower, and aircraft are shown in Figures S2–S4, respectively.



**Figure 11.** Boxplots of the different types of monthly mean observations, (a) for TCCON observations, (b) for tower observations, (c) for aircraft observations, and the corresponding monthly mean XCO<sub>2</sub> or CO<sub>2</sub> concentrations from SIM, DA\_FLUX and CT2017. The lower and upper edges of the boxes represent the first (Q1) and third (Q3) quartiles, the lower and upper whiskers represent the minimum and maximum values within 1.5 times the interquartile range from Q1 and Q3, respectively, solid lines and triangles represent median and mean value.

**Table 6.** Evaluation results between the monthly mean CO<sub>2</sub> concentration or XCO<sub>2</sub> observations and the corresponding results from SIM, DA\_FLUX and CT2017, where the CORRs are statistically significant at the level of  $p < 0.01$ .

Observation Type	Experiment	RMSE (ppm)	MBE (ppm)	CORR
TCCON	SIM	0.70	0.36	0.95
	DA_FLUX	0.81	0.17	0.92
	CT2017	0.69	−0.11	0.94
Tower	SIM	2.83	1.41	0.92
	DA_FLUX	2.63	−0.05	0.89
	CT2017	2.71	0.70	0.89
Aircraft	SIM	0.95	−0.20	0.94
	DA_FLUX	0.68	−0.05	0.97
	CT2017	0.65	−0.10	0.97



The RMSE between the monthly mean CO<sub>2</sub> concentrations of DA\_FLUX and the observations of tower and aircraft were 2.63 ppm and 0.68 ppm, which were 0.20 ppm and 0.27 ppm smaller than those of SIM, respectively. The RMSE between the monthly mean XCO<sub>2</sub> of DA\_FLUX and the TCCON observations was 0.81 ppm, which was 0.11 ppm larger than that of SIM. The reason is, shown in Figure S2, the monthly mean XCO<sub>2</sub> values estimated by DA\_FLUX spread more widely but symmetrically around the TCCON observations than those of SIM. The MBE between the monthly mean CO<sub>2</sub> concentrations of DA\_FLUX and the observations of TCCON, tower, and aircraft were 0.17 ppm, −0.05 ppm, and −0.05 ppm, which were 0.19 ppm, 1.36 ppm, and 0.15 ppm smaller than those of SIM, respectively. Figure S3 and S4 show the comparison of the monthly mean tower and aircraft observations with the modeled CO<sub>2</sub> concentrations from the SIM, DA\_FLUX experiment, and CT2017, respectively.

#### 4. Conclusions

A regional CO<sub>2</sub> flux inversion system had been developed from WRF-Chem model coupled with DART using the EAKF assimilation method to invert regional CO<sub>2</sub> flux from OCO-2 XCO<sub>2</sub> retrievals. The effectiveness of the inversion system was verified by the DA\_FLUX experiment over the contiguous United States in 2016. The posterior CO<sub>2</sub> fluxes of the DA\_FLUX experiment were compared with the corresponding fluxes extracted from OCO-2 MIP and CT2017 and was evaluated with the independent ground-based flux measurements. The inverted annual CO<sub>2</sub> flux over the contiguous United States in 2016 was 1.08 Pg C yr<sup>−1</sup>, which was 0.05 Pg C yr<sup>−1</sup> higher than the largest flux from two models in OCO-2 MIP and was 0.11 Pg C yr<sup>−1</sup> lower than the posterior CO<sub>2</sub> flux of CT2017. The uncertainty of the inverted annual CO<sub>2</sub> flux was reduced by 14.71% compared to the prior flux, which demonstrated the impact of OCO-2 XCO<sub>2</sub> retrievals on the flux constraints. The evaluation with the flux measurements showed that RMSE and MBE between the posterior CO<sub>2</sub> flux of the DA\_FLUX experiment and the ground observations were reduced by 16.84% and 22.34% compared with that of the prior flux and were reduced by 25.47% and 28.43% compared with that of CT2017, respectively. The CO<sub>2</sub> concentrations estimated by the DA\_FLUX experiment were more in line with CT2017 than that of the SIM experiment which did not assimilate the OCO-2 XCO<sub>2</sub> retrievals. The annual mean XCO<sub>2</sub> of the DA\_FLUX experiment over the contiguous United States in 2016 was 403.67 ppm which was 0.11 ppm smaller than that of the SIM experiment (403.78 ppm) and closer to CT2017 (403.29 ppm). MBE between the inverted CO<sub>2</sub> flux and flux measurements was 0.73 g C m<sup>−2</sup> d<sup>−1</sup>, was reduced by 22.34% and 28.43% compared to those of the prior flux and CT2017, respectively. MBEs between the CO<sub>2</sub> concentrations of the DA\_FLUX experiment and concentration measurements from TCCON, towers, and aircraft were reduced by 52.78%, 96.45%, and 75%, respectively, compared to those of the SIM experiment. These experiments and evaluations indicated that the regional CO<sub>2</sub> flux inversion system had a reasonable performance.

Finally, the significant drop in uncertainty reduction of the inverted CO<sub>2</sub> flux of the “Upper Midwest” region in September 2016 due to the lack of OCO-2 XCO<sub>2</sub> retrievals, and the fact that CO<sub>2</sub> emission hotspots indicated by the inverted annual CO<sub>2</sub> flux with a relatively high spatial resolution of 50 km consisted well with the locations of most major metropolitan/urban areas in the contiguous United States, demonstrated the indispensable impact of OCO-2 XCO<sub>2</sub> retrievals on the regional flux constraints and the potential of combining satellite observations with high spatial resolution CO<sub>2</sub> flux inversion system in supporting the global stocktake.

**Supplementary Materials:** The following are available online at <https://www.mdpi.com/article/10.3390/rs13152996/s1>, Figure S1: Spatial distributions of monthly mean posterior CO<sub>2</sub> fluxes estimated by the DA\_FLUX experiment over the contiguous United States in 2016; Figure S2: Comparison of the monthly mean TCCON observations and the modeled XCO<sub>2</sub> from the SIM, DA\_FLUX experiment and CT2017. The top part shows scatter plots of the modeled XCO<sub>2</sub> against the TCCON observations. In the lower part, dot represents the mean bias error between the modeled XCO<sub>2</sub> and the TCCON

observations, bar represents the standard deviation of the bias errors; Figure S3: Comparison of the monthly mean tower observations with the modeled CO<sub>2</sub> concentrations from the SIM, DA\_FLUX experiment, and CT2017. The top part shows scatter plots of the modeled CO<sub>2</sub> concentrations against the tower observations. In the lower part, dot represents the mean bias error between the modeled CO<sub>2</sub> concentrations and the tower observations, bar represents the standard deviation of the bias errors; Figure S4: The comparison of the monthly mean aircraft observations with the modeled CO<sub>2</sub> concentrations from the SIM, DA\_FLUX experiment, and CT2017. The top part shows scatter plots of the modeled CO<sub>2</sub> concentrations against the aircraft observations. In the lower part, dot represents the mean bias error between the modeled CO<sub>2</sub> concentrations and the aircraft observations, bar represents the standard deviation of the bias errors; Table S1: The information of inversion models in OCO-2 MIP version 7; Table S2: The information of CO<sub>2</sub> fluxes tower sites from AmeriFlux; Table S3: The information of CO<sub>2</sub> fluxes tower sites from ONEFlux; Table S4: The information of ground-based CO<sub>2</sub> concentration observation sites; Table S5: Aircraft observations from ObsPack products used for evaluation; Table S6: The monthly prior (PRIOR) and posterior (DA\_FLUX) CO<sub>2</sub> flux of the DA\_FLUX experiment, the monthly posterior fluxes from CT2017 and the models in OCO-2 MIP, over the contiguous United States in 2016; Table S7: Monthly mean of the prior (PRIOR) and posterior (DA\_FLUX) CO<sub>2</sub> flux of the DA\_FLUX experiment, the posterior flux of CT2017, and the CO<sub>2</sub> flux measurements; Table S8: Mean and median CO<sub>2</sub> concentration values of the SIM, DA\_FLUX experimental results, and observations from TCCON, tower, and aircraft.

**Author Contributions:** Q.Z.: conceptualization, methodology, software, validation, formal analysis, investigation, data curation, writing—original draft preparation, writing—review and editing, visualization. M.L.: writing—review and editing, supervision. M.W.: resources, supervision. A.P.M.: software, writing—review and editing. Y.H.: writing—review and editing. C.W.: writing—review and editing. J.J.: writing—review and editing. Q.G.: conceptualization, formal analysis, resources, writing—original draft preparation, writing—review and editing, supervision. All authors have read and agreed to the published version of the manuscript.

**Funding:** This research was funded by the Key Research Program of Chinese Academy of Sciences (ZDRW-ZS-2019-1) and the International Partnership Program of Chinese Academy of Sciences (131211KYSB20180002).

**Institutional Review Board Statement:** Not applicable.

**Informed Consent Statement:** Not applicable.

**Data Availability Statement:** The CO<sub>2</sub> flux and concentration data over the contiguous United States in 2016 generated by this study are openly available in: [https://github.com/aqap2012/remote\\_sensing\\_2021\\_CO2](https://github.com/aqap2012/remote_sensing_2021_CO2), accessed on 18 June 2021.

**Acknowledgments:** The authors acknowledge the free availability of the WRF-Chem model (<https://www2.acom.ucar.edu/wrf-chem>, last access: 18 June 2021), DART system (<https://svn-dares-dart.cgd.ucar.edu/DART>, last access: 18 June 2021), CarbonTracker CT2017 results provided by NOAA ESRL, Boulder, Colorado, USA (<http://carbontracker.noaa.gov>, last access: 18 June 2021), OCO-2 data from the NASA Goddard Earth Science Data and Information Services Center ([https://disc.gsfc.nasa.gov/datasets/OCO2\\_L2\\_Lite\\_FP\\_9r/summary](https://disc.gsfc.nasa.gov/datasets/OCO2_L2_Lite_FP_9r/summary), last access: 18 June 2021), XCO<sub>2</sub> measurements from TCCON (<https://tccodata.org/>, last access: 18 June 2021), the meteorological data from NCEP (<https://rda.ucar.edu/datasets/ds083.2/>, last access: 18 June 2021), CO<sub>2</sub> observations from the ObsPack product (<https://www.esrl.noaa.gov/gmd/ccgg/obspace/data.php>, last access: 18 June 2021), Eddy covariance measurements from AmeriFlux (<https://ameriflux.lbl.gov/>, last access: 18 June 2021) and ONEFlux (<https://ameriflux.lbl.gov/data/download-data-oneflux-beta/>, last access: 18 June 2021), OCO-2 model intercomparison project (MIP) level-4 flux product from <https://www.esrl.noaa.gov/gmd/ccgg/OCO2/> (last access: 18 June 2021), CO<sub>2</sub> flux from the CarbonTracker (<ftp://afp.cmdl.noaa.gov/products/carbontracker/co2/CT2017/fluxes/priors>, last access: 18 June 2021).

**Conflicts of Interest:** The authors declare no conflict of interest.

## References

1. IPCC. *Climate Change 2014: Synthesis Report*; IPCC: Geneva, Switzerland, 2014; p. 151.
2. Friedlingstein, P.; O'Sullivan, M.; Jones, M.W.; Andrew, R.M.; Hauck, J.; Olsen, A.; Peters, G.P.; Peters, W.; Pongratz, J.; Sitch, S.; et al. Global Carbon Budget 2020. *Earth Syst. Sci. Data* **2020**, *12*, 3269–3340. [[CrossRef](#)]
3. UNFCCC. The Paris Agreement. Available online: [http://unfccc.int/files/essential\\_background/convention/application/pdf/english\\_paris\\_agreement.pdf](http://unfccc.int/files/essential_background/convention/application/pdf/english_paris_agreement.pdf) (accessed on 8 June 2021).
4. IPCC. 2019 Refinement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories. Available online: <https://www.ipcc.ch/report/2019-refinement-to-the-2006-ipcc-guidelines-for-national-greenhouse-gas-inventories/> (accessed on 29 July 2021).
5. UNFCCC-SBSTA. Subsidiary Body for Scientific and Technological Advice. Available online: <https://unfccc.int/resource/docs/2017/sbsta/eng/07.pdf> (accessed on 20 July 2021).
6. UNFCCC. INFORMATION PAPER Systematic Observations. Available online: [https://unfccc.int/sites/default/files/resource/Mandates\\_systematic\\_%20observation.pdf](https://unfccc.int/sites/default/files/resource/Mandates_systematic_%20observation.pdf) (accessed on 20 July 2021).
7. Broquet, G.; Chevallier, F.; Rayner, P.; Aulagnier, C.; Pison, I.; Ramonet, M.; Schmidt, M.; Vermeulen, A.T.; Ciais, P. A European summertime CO<sub>2</sub> biogenic flux inversion at mesoscale from continuous in situ mixing ratio measurements. *J. Geophys. Res. Atmos.* **2011**, *116*. [[CrossRef](#)]
8. Broquet, G.; Chevallier, F.; Bréon, F.M.; Kadygrov, N.; Alemanno, M.; Apadula, F.; Hammer, S.; Haszpra, L.; Meinhardt, F.; Morguá, J.A.; et al. Regional inversion of CO<sub>2</sub> ecosystem fluxes from atmospheric measurements: Reliability of the uncertainty estimates. *Atmos. Chem. Phys.* **2013**, *13*, 9039–9056. [[CrossRef](#)]
9. Peters, W.; Miller, J.B.; Whitaker, J.; Denning, A.S.; Hirsch, A.; Krol, M.C.; Zupanski, D.; Bruhwiler, L.; Tans, P.P. An ensemble data assimilation system to estimate CO<sub>2</sub> surface fluxes from atmospheric trace gas observations. *J. Geophys. Res.-Atmos.* **2005**, *110*. [[CrossRef](#)]
10. Peters, W.; Jacobson, A.R.; Sweeney, C.; Andrews, A.E.; Conway, T.J.; Masarie, K.; Miller, J.B.; Bruhwiler, L.M.; Petron, G.; Hirsch, A.I.; et al. An atmospheric perspective on North American carbon dioxide exchange: CarbonTracker. *Proc. Natl. Acad. Sci. USA* **2007**, *104*, 18925–18930. [[CrossRef](#)] [[PubMed](#)]
11. Peylin, P.; Law, R.M.; Gurney, K.R.; Chevallier, F.; Jacobson, A.R.; Maki, T.; Niwa, Y.; Patra, P.K.; Peters, W.; Rayner, P.J.; et al. Global atmospheric carbon budget: Results from an ensemble of atmospheric CO<sub>2</sub> inversions. *Biogeosciences* **2013**, *10*, 6699–6720. [[CrossRef](#)]
12. Thompson, R.L.; Patra, P.K.; Chevallier, F.; Maksyutov, S.; Law, R.M.; Ziehn, T.; van der Laan-Luijckx, I.T.; Peters, W.; Ganshin, A.; Zhuravlev, R.; et al. Top-down assessment of the Asian carbon budget since the mid 1990s. *Nat. Commun.* **2016**, *7*, 10724. [[CrossRef](#)]
13. Kuze, A.; Suto, H.; Nakajima, M.; Hamazaki, T. Thermal and near infrared sensor for carbon observation Fourier-transform spectrometer on the Greenhouse Gases Observing Satellite for greenhouse gases monitoring. *Appl. Opt.* **2009**, *48*, 6716–6733. [[CrossRef](#)]
14. Glumb, R.; Davis, G.; Lietzke, C. The TANSO-FTS-2 instrument for the GOSAT-2 greenhouse gas monitoring mission. In Proceedings of the IGARSS 2014-2014 IEEE International Geoscience and Remote Sensing Symposium, Quebec City, QC, Canada, 13–18 July 2014; pp. 1238–1240.
15. Eldering, A.; Wennberg, P.O.; Crisp, D.; Schimel, D.S.; Gunson, M.R.; Chatterjee, A.; Liu, J.; Schwandner, F.M.; Sun, Y.; O'Dell, C.W.; et al. The Orbiting Carbon Observatory-2 early science investigations of regional carbon dioxide fluxes. *Science* **2017**, *358*. [[CrossRef](#)]
16. Eldering, A.; Taylor, T.E.; O'Dell, C.W.; Pavlick, R. The OCO-3 mission: Measurement objectives and expected performance based on 1 year of simulated data. *Atmos. Meas. Tech.* **2019**, *12*, 2341–2370. [[CrossRef](#)]
17. Taylor, T.E.; Eldering, A.; Merrelli, A.; Kiel, M.; Somkuti, P.; Cheng, C.; Rosenberg, R.; Fisher, B.; Crisp, D.; Basilio, R.; et al. OCO-3 early mission operations and initial (vEarly) XCO<sub>2</sub> and SIF retrievals. *Remote. Sens. Environ.* **2020**, *251*. [[CrossRef](#)]
18. Liu, Y.; Wang, J.; Yao, L.; Chen, X.; Cai, Z.; Yang, D.; Yin, Z.; Gu, S.; Tian, L.; Lu, N.; et al. The TanSat mission: Preliminary global observations. *Sci. Bull.* **2018**, *63*, 1200–1207. [[CrossRef](#)]
19. Karafolas, N.; Sodnik, Z.; Cugny, B.; Buisson, F.; Jouglet, D.; Tauziede, L.; Loesel, J.; Buil, C.; Pascal, V. An improved microcarb dispersive instrumental concept for the measurement of greenhouse gases concentration in the atmosphere. In Proceedings of the International Conference on Space Optics—ICSO 2014, Tenerife, Spain, 17 November 2017.
20. Bertaux, J.-L.; Hauchecorne, A.; Lefèvre, F.; Bréon, F.-M.; Blanot, L.; Jouglet, D.; Lafrique, P.; Akaev, P. The use of the 1.27 μm O<sub>2</sub> absorption band for greenhouse gas monitoring from space and application to MicroCarb. *Atmos. Meas. Tech.* **2020**, *13*, 3329–3374. [[CrossRef](#)]
21. Polonsky, I.N.; O'Brien, D.M.; Kumer, J.B.; O'Dell, C.W. Performance of a geostationary mission, geoCARB, to measure CO<sub>2</sub>, CH<sub>4</sub> and CO column-averaged concentrations. *Atmos. Meas. Tech.* **2014**, *7*, 959–981. [[CrossRef](#)]
22. Moore, B., III; Crowell, S.M.R.; Rayner, P.J.; Kumer, J.; O'Dell, C.W.; O'Brien, D.; Utembe, S.; Polonsky, I.; Schimel, D.; Lemen, J. The Potential of the Geostationary Carbon Cycle Observatory (GeoCarb) to Provide Multi-scale Constraints on the Carbon Cycle in the Americas. *Front. Environ. Sci.* **2018**, *6*. [[CrossRef](#)]
23. Kuhlmann, G.; Broquet, G.; Marshall, J.; Clement, V.; Loscher, A.; Meijer, Y.; Brunner, D. Detectability of CO<sub>2</sub> emission plumes of cities and power plants with the Copernicus Anthropogenic CO<sub>2</sub> Monitoring (CO2M) mission. *Atmos. Meas. Tech.* **2019**, *12*, 6695–6719. [[CrossRef](#)]

24. Janssens-Maenhout, G.; Pinty, B.; Dowell, M.; Zunker, H.; Andersson, E.; Balsamo, G.; Bézy, J.L.; Brunhes, T.; Bösch, H.; Bojkov, B.; et al. Toward an Operational Anthropogenic CO<sub>2</sub> Emissions Monitoring and Verification Support Capacity. *Bull. Am. Meteorol. Soc.* **2020**, *101*, E1439–E1451. [[CrossRef](#)]
25. Basu, S.; Guerlet, S.; Butz, A.; Houweling, S.; Hasekamp, O.; Aben, I.; Krummel, P.; Steele, P.; Langenfelds, R.; Torn, M.; et al. Global CO<sub>2</sub> fluxes estimated from GOSAT retrievals of total column CO<sub>2</sub>. *Atmos. Chem. Phys.* **2013**, *13*, 8695–8717. [[CrossRef](#)]
26. Houweling, S.; Baker, D.; Basu, S.; Boesch, H.; Butz, A.; Chevallier, F.; Deng, F.; Dlugokencky, E.J.; Feng, L.; Ganshin, A.; et al. An intercomparison of inverse models for estimating sources and sinks of CO<sub>2</sub> using GOSAT measurements. *J. Geophys. Res. Atmos.* **2015**, *120*, 5253–5266. [[CrossRef](#)]
27. Maksyutov, S.; Takagi, H.; Valsala, V.K.; Saito, M.; Oda, T.; Saeki, T.; Belikov, D.A.; Saito, R.; Ito, A.; Yoshida, Y.; et al. Regional CO<sub>2</sub> flux estimates for 2009–2010 based on GOSAT and ground-based CO<sub>2</sub> observations. *Atmos. Chem. Phys.* **2013**, *13*, 9351–9373. [[CrossRef](#)]
28. Deng, F.; Jones, D.B.A.; Henze, D.K.; Bousseres, N.; Bowman, K.W.; Fisher, J.B.; Nassar, R.; O'Dell, C.; Wunch, D.; Wennberg, P.O.; et al. Inferring regional sources and sinks of atmospheric CO<sub>2</sub> from GOSAT XCO<sub>2</sub> data. *Atmos. Chem. Phys.* **2014**, *14*, 3703–3727. [[CrossRef](#)]
29. Liu, J.; Bowman, K.W.; Lee, M. Comparison between the Local Ensemble Transform Kalman Filter (LETKF) and 4D-Var in atmospheric CO<sub>2</sub> flux inversion with the Goddard Earth Observing System-Chem model and the observation impact diagnostics from the LETKF. *J. Geophys. Res. Atmos.* **2016**, *121*, 13,066–13,087. [[CrossRef](#)]
30. Chevallier, F.; Palmer, P.I.; Feng, L.; Boesch, H.; O'Dell, C.W.; Bousquet, P. Toward robust and consistent regional CO<sub>2</sub> flux estimates from in situ and spaceborne measurements of atmospheric CO<sub>2</sub>. *Geophys. Res. Lett.* **2014**, *41*, 1065–1070. [[CrossRef](#)]
31. Deng, F.; Jones, D.B.A.; O'Dell, C.W.; Nassar, R.; Parazoo, N.C. Combining GOSAT XCO<sub>2</sub> observations over land and ocean to improve regional CO<sub>2</sub> flux estimates. *J. Geophys. Res. Atmos.* **2016**, *121*, 1896–1913. [[CrossRef](#)]
32. Crowell, S.; Baker, D.; Schuh, A.; Basu, S.; Jacobson, A.R.; Chevallier, F.; Liu, J.; Deng, F.; Feng, L.; McKain, K.; et al. The 2015–2016 carbon cycle as seen from OCO-2 and the global in situ network. *Atmos. Chem. Phys.* **2019**, *19*, 9797–9831. [[CrossRef](#)]
33. Chevallier, F.; Remaud, M.; O'Dell, C.W.; Baker, D.; Peylin, P.; Cozic, A. Objective evaluation of surface- and satellite-driven carbon dioxide atmospheric inversions. *Atmos. Chem. Phys.* **2019**, *19*, 14233–14251. [[CrossRef](#)]
34. Wang, H.; Jiang, F.; Wang, J.; Ju, W.; Chen, J.M. Terrestrial ecosystem carbon flux estimated using GOSAT and OCO-2 XCO<sub>2</sub> retrievals. *Atmos. Chem. Phys.* **2019**, *19*, 12067–12082. [[CrossRef](#)]
35. Villalobos, Y.; Rayner, P.; Thomas, S.; Silver, J. The potential of Orbiting Carbon Observatory-2 data to reduce the uncertainties in CO<sub>2</sub> surface fluxes over Australia using a variational assimilation scheme. *Atmos. Chem. Phys.* **2020**, *20*, 8473–8500. [[CrossRef](#)]
36. Zheng, T.; French, N.H.F.; Baxter, M. Development of the WRF-CO<sub>2</sub> 4D-Var assimilation system v1.0. *Geosci. Model Dev.* **2018**, *11*, 1725–1752. [[CrossRef](#)]
37. Peng, Z.; Zhang, M.; Kou, X.; Tian, X.; Ma, X. A regional carbon data assimilation system and its preliminary evaluation in East Asia. *Atmos. Chem. Phys.* **2015**, *15*, 1087–1104. [[CrossRef](#)]
38. Zhang, Q.; Li, M.; Wei, C.; Mizzi, A.P.; Huang, Y.; Gu, Q. Assimilation of OCO-2 retrievals with WRF-Chem/DART: A case study for the Midwestern United States. *Atmos. Environ.* **2021**, *246*, 118106. [[CrossRef](#)]
39. Grell, G.A.; Peckham, S.E.; Schmitz, R.; McKeen, S.A.; Frost, G.; Skamarock, W.C.; Eder, B. Fully coupled “online” chemistry within the WRF model. *Atmos. Environ.* **2005**, *39*, 6957–6975. [[CrossRef](#)]
40. Anderson, J.; Hoar, T.; Raeder, K.; Liu, H.; Collins, N.; Torn, R.; Avellano, A. The Data Assimilation Research Testbed A Community Facility. *Bull. Am. Meteorol. Soc.* **2009**, *90*, 1283–1296. [[CrossRef](#)]
41. Anderson, J.L. An ensemble adjustment Kalman filter for data assimilation. *Mon. Weather Rev.* **2001**, *129*, 2884–2903. [[CrossRef](#)]
42. Anderson, J.L. A Local Least Squares Framework for Ensemble Filtering. *Mon. Weather Rev.* **2003**, *131*, 634–642. [[CrossRef](#)]
43. NCEP. NCEP FNL Operational Model Global Tropospheric Analyses, continuing from July 1999. 2000. Available online: <https://data.ucar.edu/dataset/ncep-fnl-operational-model-global-tropospheric-analyses-continuing-from-july-19993> (accessed on 29 July 2021).
44. Eldering, A.; O'Dell, C.W.; Wennberg, P.O.; Crisp, D.; Gunson, M.R.; Viatte, C.; Avis, C.; Braverman, A.; Castano, R.; Chang, A.; et al. The Orbiting Carbon Observatory-2: First 18 months of science data products. *Atmos. Meas. Tech.* **2017**, *10*, 549–563. [[CrossRef](#)]
45. Crisp, D.; Pollock, H.R.; Rosenberg, R.; Chapsky, L.; Lee, R.A.M.; Oyafuso, F.A.; Frankenberg, C.; O'Dell, C.W.; Bruegge, C.J.; Doran, G.B.; et al. The on-orbit performance of the Orbiting Carbon Observatory-2 (OCO-2) instrument and its radiometrically calibrated products. *Atmos. Meas. Tech.* **2017**, *10*, 59–81. [[CrossRef](#)]
46. Osterman, G.; Eldering, A.; Avis, C.; Chafin, B.; O'Dell, C.; Frankenberg, C.; Fisher, B.; Mandrake, L.; Wunch, D.; Granat, R.; et al. Orbiting Carbon Observatory-2 (OCO-2) Data Product User's Guide, Operational L1 and L2 Data Versions 8 and Lite File Version 9, Version 1, Revision J., October 10, 2018. Available online: [https://docserver.gesdisc.eosdis.nasa.gov/public/project/OCO/OCO2\\_DUG.V9.pdf](https://docserver.gesdisc.eosdis.nasa.gov/public/project/OCO/OCO2_DUG.V9.pdf) (accessed on 8 June 2021).
47. Wunch, D.; Wennberg, P.O.; Osterman, G.; Fisher, B.; Naylor, B.; Roehl, C.M.; O'Dell, C.; Mandrake, L.; Viatte, C.; Kiel, M.; et al. Comparisons of the Orbiting Carbon Observatory-2 (OCO-2) XCO<sub>2</sub> measurements with TCCON. *Atmos. Meas. Tech.* **2017**, *10*, 2209–2238. [[CrossRef](#)]

48. OCO-2 Science Team/Michael Gunson; Eldering, A. OCO-2 Level 2 Bias-Corrected XCO<sub>2</sub> and Other Select Fields from the Full-Physics Retrieval Aggregated as Daily Files, Retrospective Processing V9r. 2018. Available online: [https://disc.gsfc.nasa.gov/datasets/OCO2\\_L2\\_Lite\\_FP\\_9r/summary](https://disc.gsfc.nasa.gov/datasets/OCO2_L2_Lite_FP_9r/summary) (accessed on 29 July 2021).
49. Kiel, M.; O'Dell, C.W.; Fisher, B.; Eldering, A.; Nassar, R.; MacDonald, C.G.; Wennberg, P.O. How bias correction goes wrong: Measurement of XCO<sub>2</sub> affected by erroneous surface pressure estimates. *Atmos. Meas. Tech.* **2019**, *12*, 2241–2259. [[CrossRef](#)]
50. Connor, B.J.; Boesch, H.; Toon, G.; Sen, B.; Miller, C.; Crisp, D. Orbiting Carbon Observatory: Inverse method and prospective error analysis. *J. Geophys. Res. Atmos.* **2008**, *113*. [[CrossRef](#)]
51. Ma, C.; Wang, T.; Mizzi, A.P.; Anderson, J.L.; Zhuang, B.; Xie, M.; Wu, R. Multiconstituent Data Assimilation with WRF-Chem/DART: Potential for Adjusting Anthropogenic Emissions and Improving Air Quality Forecasts Over Eastern China. *J. Geophys. Res. Atmos.* **2019**, *124*, 7393–7412. [[CrossRef](#)]
52. Mizzi, A.P.; Edwards, D.P.; Anderson, J.L. Assimilating compact phase space retrievals (CPSRs): Comparison with independent observations (MOZAIC in situ and IASI retrievals) and extension to assimilation of truncated retrieval profiles. *Geosci. Model Dev.* **2018**, *11*, 3727–3745. [[CrossRef](#)]
53. Liu, X.; Mizzi, A.P.; Anderson, J.L.; Fung, I.Y.; Cohen, R.C. Assimilation of satellite NO<sub>2</sub> observations at high spatial resolution using OSSEs. *Atmos. Chem. Phys.* **2017**, *17*, 7067–7081. [[CrossRef](#)]
54. Mizzi, A.P.; Arellano, A.F., Jr.; Edwards, D.P.; Anderson, J.L.; Pfister, G.G. Assimilating compact phase space retrievals of atmospheric composition with WRF-Chem/DART: A regional chemical transport/ensemble Kalman filter data assimilation system. *Geosci. Model Dev.* **2016**, *9*, 965–978. [[CrossRef](#)]
55. Kang, J.-S.; Kalnay, E.; Liu, J.; Fung, I.; Miyoshi, T.; Ide, K. “Variable localization” in an ensemble Kalman filter: Application to the carbon cycle data assimilation. *J. Geophys. Res.* **2011**, *116*. [[CrossRef](#)]
56. Kang, J.-S.; Kalnay, E.; Miyoshi, T.; Liu, J.; Fung, I. Estimation of surface carbon fluxes with an advanced data assimilation methodology. *J. Geophys. Res. Atmos.* **2012**, *117*. [[CrossRef](#)]
57. Gaspari, G.; Cohn, S.E. Construction of correlation functions in two and three dimensions. *Q. J. R. Meteorol. Soc.* **1999**, *125*, 723–757. [[CrossRef](#)]
58. Liu, Y.; Kalnay, E.; Zeng, N.; Asrar, G.; Chen, Z.; Jia, B. Estimating surface carbon fluxes based on a local ensemble transform Kalman filter with a short assimilation window and a long observation window: An observing system simulation experiment test in GEOS-Chem 10.1. *Geosci. Model Dev.* **2019**, *12*, 2899–2914. [[CrossRef](#)]
59. Novick, K.A.; Biederman, J.A.; Desai, A.R.; Litvak, M.E.; Moore, D.J.P.; Scott, R.L.; Torn, M.S. The AmeriFlux network: A coalition of the willing. *Agric. For. Meteorol.* **2018**, *249*, 444–456. [[CrossRef](#)]
60. Pastorello, G.; Trotta, C.; Canfora, E.; Chu, H.; Christianson, D.; Cheah, Y.W.; Poindexter, C.; Chen, J.; Elbashandy, A.; Humphrey, M.; et al. The FLUXNET2015 dataset and the ONEFlux processing pipeline for eddy covariance data. *Sci. Data* **2020**, *7*, 225. [[CrossRef](#)]
61. Wunch, D.; Toon, G.C.; Blavier, J.-F.; Washenfelder, R.A.; Notholt, J.; Connor, B.J.; Griffith, D.W.; Sherlock, V.; Wennberg, P.O. The total carbon column observing network. *Philos. Trans. R. Soc. A* **2011**, *369*, 2087–2112. [[CrossRef](#)]
62. Wennberg, P.O.; Roehl, C.M.; Wunch, D.; Toon, G.C.; Blavier, J.-F.; Washenfelder, R.; Keppel-Aleks, G.; Allen, N.T.; Ayers, J. TCCON data from Park Falls (US); Release GGG2014.R1; 2017. Available online: <https://data.caltech.edu/records/295> (accessed on 29 July 2021).
63. Wennberg, P.O.; Wunch, D.; Roehl, C.M.; Blavier, J.-F.; Toon, G.C.; Allen, N.T. TCCON Data from Lamont (US); Release GGG2014.R1; 2016. Available online: <https://data.caltech.edu/records/279> (accessed on 29 July 2021).
64. Iraci, L.T.; Podolske, J.R.; Hillyard, P.W.; Roehl, C.; Wennberg, P.O.; Blavier, J.-F.; Landeros, J.; Allen, N.; Wunch, D.; Zavaleta, J.; et al. TCCON Data from Edwards (US); Release GGG2014.R1; 2016. Available online: <https://data.caltech.edu/records/270> (accessed on 29 July 2021).
65. Wennberg, P.O.; Wunch, D.; Roehl, C.M.; Blavier, J.-F.; Toon, G.C.; Allen, N.T. TCCON Data from Caltech (US); Release GGG2014.R1; 2015. Available online: <https://data.caltech.edu/records/285> (accessed on 29 July 2021).
66. Cooperative Global Atmospheric Data Integration Project. Multi-laboratory compilation of atmospheric carbon dioxide data for the period 1957–2018. 2019. Available online: [https://gml.noaa.gov/ccgg/obspack/providerlist/obspack\\_co2\\_1\\_GLOBALVIEWplus\\_v4.2\\_2019-03-19.html](https://gml.noaa.gov/ccgg/obspack/providerlist/obspack_co2_1_GLOBALVIEWplus_v4.2_2019-03-19.html) (accessed on 29 July 2021).
67. Masarie, K.A.; Peters, W.; Jacobson, A.R.; Tans, P.P. ObsPack: A framework for the preparation, delivery, and attribution of atmospheric greenhouse gas measurements. *Earth Syst. Sci. Data* **2014**, *6*, 375–384. [[CrossRef](#)]
68. Chevallier, F.; Bréon, F.-M.; Rayner, P.J. Contribution of the Orbiting Carbon Observatory to the estimation of CO<sub>2</sub> sources and sinks: Theoretical study in a variational data assimilation framework. *J. Geophys. Res.* **2007**, *112*. [[CrossRef](#)]
69. Lauvaux, T.; Schuh, A.E.; Uliasz, M.; Richardson, S.; Miles, N.; Andrews, A.E.; Sweeney, C.; Diaz, L.I.; Martins, D.; Shepson, P.B.; et al. Constraining the CO<sub>2</sub> budget of the corn belt: Exploring uncertainties from the assumptions in a mesoscale inverse system. *Atmos. Chem. Phys.* **2012**, *12*, 337–354. [[CrossRef](#)]