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Carbon and Energy Balance in a Primary Amazonian Forest and Its Relationship with Remote Sensing Estimates

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Abstract: With few measurement sites and a great need to validate satellite data to characterize

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the exchange of energy and carbon fluxes in tropical forest areas, quantified by the Net Ecosystem Exchange (NEE) and associated with phenological measurements, there is an increasing need for studies aimed at characterizing the Amazonian environment in its biosphere-atmosphere interaction, considering the accelerated deforestation in recent years. Using data from a flux measurement tower in the Caxiuana-PA forest (2005-2008), climatic data, CO2 exchange estimated by eddy covariance, as well as Gross Primary Productivity (GPP) data and satellite vegetation indices (from MODIS), this work aimed to describe the site's energy, climatic and carbon cycle flux patterns, correlating its gross primary productivity with satellite vegetation indices. The results found were: (1) marked seasonality of climatic variables and energy flows, with evapotranspiration and air temperature on the site following the annual march of solar radiation and precipitation; (2) energy fluxes in phase and dependent on available energy; (3) the site as a carbon sink $(-569.7 \pm 444.9 \text{ gC m}^{-2} \text{ year}^{-1})$, with intensity varying according to the site's annual water availability; (4) low correlation between productivity data and vegetation indices, corroborating data in the literature on these variables in this type of ecosystem. The results show the importance of preserving this type of environment for the mitigation of global warming and the need to improve satellite estimates for this region. NDVI and EVI patterns follow radiative availability, as does LAI, but without direct capture related to GPP data, which correlates better with satellite data only in the months with the highest LAI. The results show the significant difference at a point measurement to a satellite interpolation, presenting how important preserving any type of environment is, even related to its size, for the global climate balance, and also the need to improve satellite estimates for smaller areas.



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1. Introduction

Many studies on carbon and water exchange have been carried out in different biomes in South America, such as the Caatinga [1–7], grazed pasture [8,9], Pantanal [10,11], Cerrado [12,13], and the Amazon, which probably concentrates most of the studies [14–18], in addition to integrated analyses of the biomes [19–21]. Understanding the seasonality of exchanges and the resistance and adaptability of species to increasingly frequent water stresses [22] is of the utmost importance in studies on biosphere–atmosphere interaction with a focus on South America, especially the Amazon biome, in addition to environmental studies [23–25] to characterize the impacts of environmental changes on soil properties.

The Amazon Basin is at the center of increasingly intense discussions on deforestation, land use and global change, such as studies on carbon (C) cycling and storage and its implications for the global climate [26], as well as factors that affect the quality of life of the population [15]. Forests like this harbor around 10 to 15% of the Earth's biodiversity, and have abundant rainfall of around 2200 mm/year, which makes the region an important source of heat and humidity for the atmosphere and generates an estimated 210,000 m³/s to 220,000 m³/s of river flow, which is ~15% of the freshwater inflow into the oceans. In addition, it stores around 150 to 200 billion tons of carbon and presents a mosaic of ethnological and linguistic diversity [27]. An increasing number of studies [28–30] are aimed at characterizing forest environments in terms of their soil, atmosphere and water characteristics and properties and their contributions to greenhouse gas emissions and climate change mitigation, particularly in Brazil's Amazon region.

Given its significance, studies on Amazonia related to water, carbon and nutrient cycles, as well as solar energy balances, as a result of changes in the region's vegetation cover, are frequently carried out by researchers from the Large Scale Biosphere–Atmosphere Program in the Amazon (LBA). Wofsy et al. [31] showed that high CO_2 concentrations in the early morning occur due to ecosystem respiration in a stable nocturnal atmospheric layer, and some previous studies point to the Amazon rainforest as a potential sink for atmospheric CO_2 [32–34] with micrometeorological measurements via flux towers.

Previous studies have indicated seasonal patterns in gross primary production (GPP), ecosystem respiration (RECO), net ecosystem exchange (NEE), energy exchange, phenology and tree growth [35–38], and they usually found that the seasonal course of canopy photosynthesis was largely controlled by phenology and light. GPP refers to the total amount of carbon fixed by plants in an ecosystem through the process of photosynthesis, and it is an important component of the carbon balance that has been analyzed in studies since the 2000s, whereas carbon cycle studies seek to support tower measurements with data from biometric or satellite measurements [35–38].

Saleska et al. [36] compared micrometeorological estimates and biometric measurements in Santarém-PA, and they reported a CO_2 source to the atmosphere of 1.3 Mg C ha⁻¹ year⁻¹ associated with the prevalence of emissions from the decomposition of necromass in previous episodes of high tree mortality in the region. Miller et al. [37] reported slightly positive ecosystem fluxes (considering NEE > 0 = source, NEE < 0 = sink) of 0.4 Mg C ha⁻¹ year⁻¹ corroborated by biometric measurements. Espírito Santo et al. [38] combined satellite estimates and field data to suggest that uptake by living trees exceeds emission by dead trees, reinforcing evidence of terra firme forest acting as a carbon sink in aerial biomass.

There has been an increase in the number of studies using remote sensing data [39–41] to characterize environmental variables in large areas, sometimes with a very low density of in situ data. Validation of this information is required to increase the security of the information generated and to have confidence in the public policies and mitigating actions to be taken based on these data. This information is important for studies on forest

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biomass [42,43] and other applications [44–46], and although it has good accuracy and validation in some regions, it still requires validation in many remote areas, such as the MODIS products for the Amazon.

The Moderate Resolution Imaging Spectroradiometer (MODIS) remote sensing data suggest that photosynthesis in native forests increases as the dry season progresses over a large area of central Amazonia. The MODIS Leaf Area Index (LAI) product shows a pattern of increase during the dry season, peaking at the beginning and in the middle of the dry season [47]. Similarly, the MODIS Enhanced Vegetation Index (EVI) shows that the equatorial forests of Amazonia continue to "green up" throughout the dry season as the availability of sunlight increases [48]. Costa et al. [20] showed that the seasonality of GPP measurements and flux tower data have different associations and correlations throughout the year depending on the type of forest, with data being better correlated in certain months of the year and in certain locations, which would make it easier to extrapolate the analysis to a large area with satellite data.

Considering the above, it is important to conduct more studies aimed at characterizing the CO_2 exchange environment in areas of the Amazon rainforest with the atmosphere and its relationship with vegetation indices estimated by satellite. Despite many studies on the subject in recent years [16,49,50], the Caxiuanã region still lacks research into these estimates for longer periods of its data series and with analyses integrated with satellite data, which are important due to the diverse mosaic of natural environments within the Amazon biome. There are particularities in the exchange of water and CO_2 with the atmosphere in the Amazon basin, and there may be areas that are sinks, neutral, or even sources of CO_2 for the atmosphere, especially in the context of climate change and the influence of severe droughts caused by low-frequency phenomena such as ENOS (El Niño Southern Oscillation).

It is known that GPP data from eddy covariance towers are not representative of the global distribution of ecosystem types [51,52], increasing the need for remote sensing approaches to estimate GPP that can provide information on ecosystem dynamics in response to climate variability at both the regional and global scales [52]. As not all environments within the same biome show the same correlations between surface data and satellite data, it is important that more studies perform this survey to validate this information and help improve or adjust estimates where it is not yet efficient.

This creates the necessity to understand these cycles, their controls and their influences, which is the purpose of this study, which aims to: (1) characterize diurnal and daily variations in energy and CO_2 fluxes in the Caxiuanã forest (Rn, H, LE, NEE); (2) quantify the CO_2 flux (g C m⁻² d⁻¹) using metrics including gross primary production (GPP), ecosystem respiration (Reco) and NEE measurements; and (3) analyze the patterns and correlations of the site's gross primary productivity with vegetation indices (NDVI, EVI and LAI from MODIS satellite) and other satellite estimates.

2. Materials and Methods

2.1. Description of the Experimental Area

The site is located inside the Caxiuanã National Forest (latitude 01°42′30″S and longitude 51°31′45″W), at Melgaço city, state of Para, about 250 km west of the capital, Belém (Figure 1). About 85% of its area is formed by dense land forest, with a median height of 30 m; however, certain trees reach 50 m [53]. The location is part of the Ferreira Pena Scientific Station (ECFPn). According to Koëppen's classification [54], the region's climate is tropical hot and humid, with minimum, average and maximum annual temperatures equal to 22, 26 and 32 °C, respectively. The rainy season in the region is from December to May, and the dry season is from June to November [55]. According to Souza Filho et al. [56], the radiation balance in the forest is lower in the rainy period compared to the dry period, a factor that can be explained by the higher solar irradiance in the dry period, due to less cloudiness in this period.

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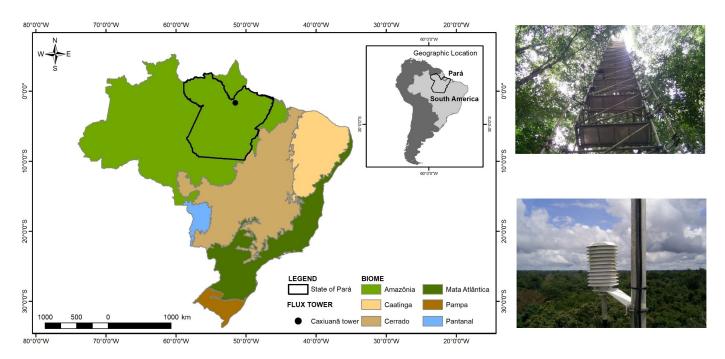


Figure 1. Location of the micrometeorological tower in the Caxiuanã-PA forest, with a view of the base of the installed tower and the temperature and humidity sensor at the top of the tower.

2.2. Database

The database of meteorological, energy and CO_2 fluxes, as well as satellite data, covers the period from 1 January 2005 to 31 December 2008. The data analyzed in this study were collected from an automatic weather station (Table 1) belonging to the Large Scale Biosphere–Atmosphere Program (LBA), installed in a 52 m-high micrometeorological observation tower. Precipitation (PRP) was measured by a Campbell TB3/TB4 (CS700, Campbell Scientific, Leicester, UK) rain gauge. Air temperature (Tar) measurements were obtained using Campbell's HMP45 AC device (Campbell Scientific, Leicester, UK). The GPP was estimated from CO_2 flux data measured in situ, following the methodology of Hutyra et al. [57], with eddy covariance data measured from a Licor 7500 Open Path CO_2/H_2O Analyzer (LICOR, Lincoln, NE, USA), which works in conjunction with a Gill instruments 3D input unic sonic anemometer (Solent, Gill Instruments, Lymington, UK).

Table 1. List of variables, measuring instruments and measuring units involved in the study (Source: LBA project and NASA).

Measurements/Sampling Rate/Averaging Period/Spatial Resolution	Data Source		
Direction (°) and Wind Speed (m/s); Eddy Covariance: Turbulent Fluxes of Latent Heat (W m $^{-2}$), Sensible Heat (W m $^{-2}$) and CO ₂ (μ mol m $^{-2}$ s $^{-1}$)/5 Hz/30 min/55 m.	Three-dimensional sonic anemometer (Solent, Gill Instruments, Lymington, UK); Li-7500 infrared gas analyzer (LICOR, Lincoln, NE, USA)		
Incident and Reflected Photosynthetically Active Radiation $(\mu \text{mol m}^{-2} \text{ s}^{-1})/1 \text{ min}/30 \text{ min}/52 \text{ m}.$	Skye quantum sensor (Skye Instruments, Powys, UK)		
Radiation Balance (W m^{-2})/1 min /30 min /45.5 m .	1 radiometer balance (CNR1, Kipp & Zonen, Delft, The Netherlands)		
Temperature (°C) and Relative Humidity (%)/1 min/30 min/53 m.	HMP45 AC (Campbell Scientific, Leicester, UK)		
Rainfall (mm)/1 min/30 min/53 m.	Bucket rainfall gauge (CS700, Campbell Scientific, Loughborough, UK)		
MODIS GPP (g C m $^{-2}$ day $^{-1}$)/8 days/500 m. (mean in 8 days)	MOD17A2H-MODIS/Terra Gross Primary Productivity		
NDVI and EVI/16 days/250 m. (mean in 16 days)	MOD13Q1-MODIS/Terra Vegetation Indices		
LAI/8 days/500 m. (mean in 8 days)	MCD15A2H–MODIS/Terra + Aqua		

All the inherent corrections of the eddy covariance method (such as coordinate rotation correction using the planar fit method, sonic virtual temperature correction, corrections for

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density fluctuation (WPL-correction), and frequency response correction) for estimating energy and CO_2 fluxes, the complete instrumentation of the tower, and the gap filling method for missing data are described in previous publications [1,49,55,57]. The information generated from the gas analyzer/sonic anemometer set was used to determine the flows and concentration of CO_2 at the ecosystem level, and this information was recorded on a data storage device (iPAQ). The evapotranspiration (*ET*) of the site was estimated using latent heat (Equation (1) [7]):

$$ET = 86,400 \cdot \frac{LE}{L} \tag{1}$$

where L (J kg⁻¹) = $103 \times (2500 - 2.37 \times Ta)$, LE is the latent heat flux, and 86,400 is the daily integration factor. The flux density of an atmospheric variable is estimated on the concept that the turbulent flow of this magnitude is calculated by its covariance with the vertical component of the wind speed. In the case of turbulent CO₂ fluxes (Fc), sensible heat flux (H) and latent heat flux (LE) can be written as [58]

$$Fc = \overline{w'\rho_c'} \tag{2}$$

$$H = \rho c_p \overline{w'T'} \tag{3}$$

$$LE = \rho L \, w' q' \tag{4}$$

where w', $\rho c'$, T' and q' are the fluctuations of the average over a time interval (usually 30 min) of the vertical wind component, CO_2 density, air temperature and air specific humidity, respectively. L is the latent heat of vaporization, ρ is the air density, and cp is the specific heat of air at constant pressure. The NEE is calculated from the total amount of turbulent flux (Fc) and the storage term (St). St is the flux associated with the CO_2 stored in the layer below the level at which the CO_2 flux measurements are made. The value of St is obtained by integrating the variation of the CO_2 concentration in the air in the layer up to the eddy covariance height. Carswell et al. [16] showed that night respiration and daytime photosynthesis eventually compensate for each other in terms of storage for the studied site, so it is possible to assume that $Fc \sim NEE$. GPP is the difference between estimated ecosystem respiration (Reco) and observed NEE:

$$GPP = Reco - NEE \tag{5}$$

Reco are the averages of the nocturnal (19–06 h) $\rm CO_2$ flux data. The vapor pressure deficit data was estimated according to Campos et al. [2] in 30-min averages. A net radiometer measured radiation net (Rn) with incoming and reflected solar and longwave radiation (CNR1, Kipp & Zonen, Delft, The Netherlands).

2.3. Remote Sensing Products

Vegetation indices play an important role in monitoring agricultural crops and vegetated surfaces using multispectral satellite imagery or other technologies. Several sensors can be used to obtain vegetation indices, but the most widely used are the Normalized Difference Vegetation Index (NDVI), which is used to monitor surface growth and vegetative vigor [59], and the Enhanced Vegetation Index (EVI), which uses the red and infrared bands like NDVI, but uses the blue band to remove atmospheric influences [59].

To access the seasonality of phenology on site, the x NDVI was used, which presents the direct relationship with leaf chlorophyll content, i.e., the higher the index, the more new leaves (leaf flush) in the canopy. The time series of surface reflectance from the MODIS sensor (on board the Terra and Aqua satellites) with atmospheric correction from the MAIAC algorithm (multi-angle implementation of atmospheric correction) have 1 km of

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spatial resolution and are temporally aggregated in 16-day compositions. These images are used to calculate the Normalized Vegetation Index (NDVI) described in Equation (6):

$$NDVI = \rho NIR - \rho Red / \rho NIR + \rho Red$$
 (6)

where NIR is near-infrared (841–876 nm) and Red is red band (620–670 nm). These wavelength values refer to the MODIS product. For the estimation of the leaf area index (LAI), the product MCD15A3H (v006)—Terra + Aqua LAI was used, which consists of the composition of 4 days in the selection of best data values in the selected period by two satellites, according to Myneni et al. [60]. However, the EVI decouples soil and atmosphere influences from the vegetation response by correcting the ρ NIR/ ρ red ratio using visible blue reflectance (ρ blue) and correction coefficients [61]:

$$EVI = 2.5 \left(\frac{\rho NIR - \rho red}{\rho NIR + 6\rho red - 7.5\rho blue + 1} \right)$$
 (7)

2.4. MOD15 and MOD17 Products

The Leaf Area Index (LAI), one of the six layers of the MOD15 product, was used to evaluate the relationship between satellite data and surface phenological patterns. The MOD15 algorithm uses daily surface reflectance values (MOD09) and data from a radiative transfer model, which are stored in a two-dimensional Lookup Table (LUT) [62]. MOD17 contains three layers, corresponding to gross primary production (GPP), net photosynthesis (PSnet) and PSnet pixel quality, with a 500 m spatial resolution. GPP is calculated on a daily scale, but the products are available with the sum for an 8-day period. The GPP derived from the MOD17 algorithm (GPPmod) is based on the linear relationship between absorbed photosynthetically active radiation and light use efficiency (LUE) [63] and is obtained according to Equation (8).

$$GPP = \varepsilon X APAR \tag{8}$$

where ε is the light use efficiency (LUE) and APAR is the absorbed photosynthetically active radiation. Although there are a large number of satellite environmental data measurement platforms [64,65], the MODIS data were chosen because of the ease of NASA's specific data site tool, which provides various satellite information for these and other land measurement points, with varied environmental information available at: https://modis.ornl.gov/sites/?id=br_para_caxiuana_forest_almeirim#product_table, accessed on 25 January 2023.

The NDVI and EVI data are averaged every 16 days, while LAI values are estimated every 8 days. In order to compare them with other variables, they were averaged over the same period as the NDVI, EVI and LAI values.

2.5. Data Quality Control and Outlier Detection

Data post-processing was conducted in three stages: (i) data quality assessment, rejecting low-quality data, data associated with sensor malfunction and visibly inconsistent data; (ii) the data were subjected to a robust outlier detection algorithm, as proposed in the literature [15]; and (iii) due to the low turbulence conditions at night, all night flow data were rejected if the friction velocity (u^*) was below a critical threshold (m/s); (iv) due to the low turbulence conditions at night, all night flow data were rejected if the friction velocity (u^*) was below a critical threshold (0.22 m/s) according to the method suggested in the literature [17,66].

The estimation of the u^* -threshold was based on the moving point test (MPT) algorithm on nighttime data, as described by the literature [67]. To eliminate spurious oscillations in the CO_2 and energy fluxes values, an algorithm based on median analysis (moving median) was used to identify spurious oscillations. The method consists of separating the series of values into a smooth part and a residual part, after which the discrepant values are manually removed.

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The gaps arising from the exclusion of spurious data during the screening process were filled using the marginal distribution sampling (MDS) gap-filling algorithm, which takes into account the covariation of fluxes with meteorological variables and also the temporal auto-correlation of fluxes [68]. In this algorithm, actions are taken under the following conditions: (i) if flow data are missing but meteorological data are available (air temperature and vapor pressure deficit—VPD), the gap is filled with the average value considering similar meteorological conditions in a 7-day window; (ii) if only incident solar radiation data are available, the gap is filled with the average value considering similar meteorological conditions in a 7-day window; (iii) if no meteorological data are available, the gap is filled with the average value of the last hour, thus considering the diurnal variation of each variable. If there are still data gaps after applying the algorithm, the same procedures are performed but consider longer time windows. The gap-filling method was performed using an online tool from the Max Planck Institute for Biogeochemistry (http://www.bgc-jena.mpg.de/~MDIwork/eddyproc/, accessed on 27 January 2023).

The hyperbolic relationship pattern (Michaelis–Menten model) between NEE and photosynthetically active radiation (PAR, or photosynthetic photon flux density (PPFD)) involves coefficients that represent the maximum photosynthetic efficiency of the canopy (a2), its yield (a3) and the average nocturnal respiration of the ecosystem (a1) as the intercept of the curve [17], and this allows for the correction of NEE daytime data failures, such that:

$$NEE = a1 + (a^2 \times PAR/a^3 + PAR)$$
(9)

3. Results

3.1. Climate Variables and Energy Fluxes

In carbon cycle characterization work, a complete and in-depth description of local atmospheric patterns is required, which are usually the main controls on site productivity variability. The study site showed marked seasonality in terms of rainfall and air temperature (Figure 2), with more intense precipitation events (rainy days) during the region's rainy season (December to May) and less intense events during the dry season (June to November). The large number of intense rainfall events causes the highest density of daily air temperature averages to be below 25 °C in the months of February and March, and the greatest variability of air temperature occurs at the end of the rainy season, with temperatures between 18 and 27.5 °C, in the month of May.

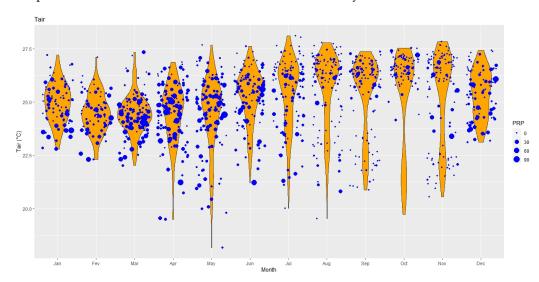


Figure 2. Data density distribution of monthly air temperature (°C) by precipitation intensity (mm/day) in Caxiuanã-PA.

The wind circulation patterns are shown in Figure 3. Throughout the year, the highest frequency of wind direction is from the NE–E quadrants, with more than 40% of the local

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wind coming from the NE direction in October. The highest average speed occurred in December (1.9 m/s), and the lowest average speed occurred in April (1.4 m/s). Wind speed tended to increase (highest averages and highest occurrences in the 2–4 and 4–6 m/s ranges) between June and December, which marks the beginning of the dry season and the start of the rainy season, with a tendency for speed to decrease from January onwards.

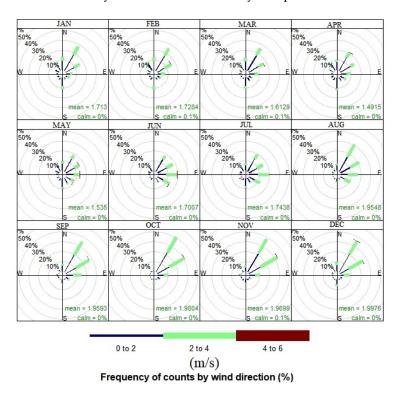


Figure 3. Annual wind patterns at the Caxiuanã site.

Monthly temperature averages tend to decrease continuously from December until reaching their minimum value in May (Figure 4a) and from this month onwards, with an increasing trend until August, coinciding with the maximum photosynthetically active radiation (Figure 4c). The maximum evapotranspiration occurs in September (Figure 4b), the same period in which the minimum rainfall occurs (Figure 4d), with the highest rainfall in April (around 400 mm). The statistical summary of the data by season is presented in Table 2.

Table 2. Seasonal midday average of meteorological variables, carbon variables and energy balance components. GPP is from the eddy covariance method. Rainfall and ET are accumulated for each period; EVI and NDVI are 16-day averages.

Variable	Dry Season	Wet Season				
variable	Mean	Sd	Max	Mean	Sd	Max
T _{air} (°C)	27.3	±3.0	31.9	23.3	±1.6	35.31
PPFD (μ mol m ⁻² s ⁻¹)	772.5	± 391.0	1667.5	9.1	± 70.6	1391.9
Rainfall (mm)	3.3	± 8.2	77.7	10.2	± 14.7	113.4
ET (mm)	1.5	± 0.46	3.0	1.2	± 0.56	2.5
$R_n (W m^{-2})$	328.1	± 208.6	802.0	12.0	± 24.1	151.2
LE (W m^{-2})	202.0	± 132.9	745.0	5.1	± 11.7	160.1
$H(W m^{-2})$	67.7	± 58.7	417.9	12.4	± 11.6	267.9
NDVI	0.83	± 0.02	0.88	0.80	± 0.04	0.86
EVI	0.51	± 0.07	0.64	0.46	± 0.07	0.70
GPP (μ mol m ⁻² s ⁻¹)	12.4	± 3.1	19.4	13.3	± 5.11	1.12
NEE (μ mol m ⁻² s ⁻¹)	-1.14	± 1.8	4.1	-2.4	± 2.9	6.7
RECO (μ mol m ⁻² s ⁻¹)	10.13	± 3.4	18.8	8.5	± 3.0	15.2

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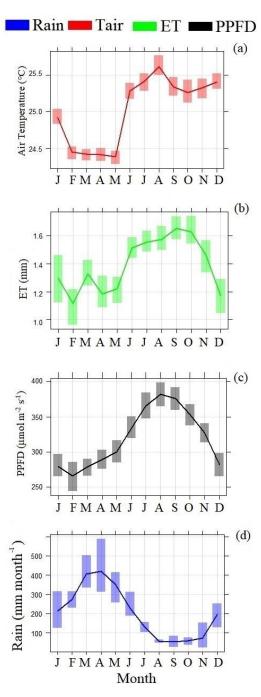


Figure 4. Monthly averages of (a) air temperature (°C), (b) Evapotranspiration (mm), (c) photosynthetic photon flux density (μ mol m⁻² s⁻¹) and (d) precipitation (mm month⁻¹) at Caxiuanã-PA site. The hatched area indicates the 95% confidence interval.

The three main components of the ecosystem's energy flux are shown in Figure 5. Net radiation (Rn) was the main energy input to the ecosystem studied and drove other processes, such as evapotranspiration, the ecosystem's gross productivity, and temperature changes. The average daily Rn peak showed seasonal variations with higher values in the dry season (above 300 W m $^{-2}$) than in the rainy season. Sensible heat flow (H) did not show such an intense pattern (below 100 W m^{-2}) compared to latent heat flow (LE), which dominates the partitioning of available energy throughout the year and showed daily averages of up to 200 W m^{-2} at the height of the dry season.

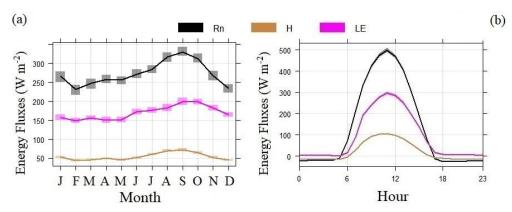


Figure 5. Monthly (**a**) and hourly (**b**) averages of the energy balance fluxes: radiation balance (Rn), sensible heat flux (H) and latent heat flux (LE), all in W m $^{-2}$. The hatched area indicates the 95% confidence interval.

3.2. Light Dependence Curves, Carbon Exchange and Vegetation Indices

Equation (6) describes the diurnal NEE in response to PAR, and this relationship for the Caxiuanã site is shown in Figure 6. It can be seen that the net CO_2 absorption balance concomitantly increases with PAR. In addition, the CO_2 flux decreases linearly with low solar radiation (PAR < 500 μ mol m⁻² s⁻¹) in both the dry season (Figure 6a) and the rainy season (Figure 6b). The light saturation point varies according to the time of day, but it is generally above 1000 μ mol m⁻² s⁻¹. The assimilation rate is higher in the afternoon for PAR values, which are up to 1000 μ mol m⁻² s⁻¹ in the dry season and up to around 700 μ mol m⁻² s⁻¹ in the rainy season.

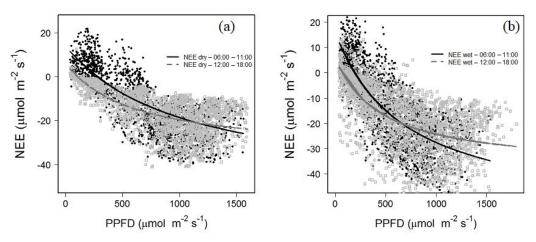


Figure 6. NEE response to PPFD in different seasons ((a) = dry; (b) = wet) and day periods.

There were no significant differences in the CO_2 exchanges detected between the dry and rainy seasons (Figure 7). Both seasons showed a peak CO_2 flux of around $-20~\mu mol~CO_2~m^{-2}~s^{-1}$. The maximum gross productivity of the site occurred in February (Figure 8a), coinciding with the maximum absorption (most negative NEE) and minimum respiration of the ecosystem. The site behaves like a CO_2 sink (negative NEE) for much of the year (from January to May and from September to December). On the other hand, during the dry season (from June to August) the site remained neutral. The seasonal dynamics of vegetation indices (EVI and NDVI) illustrate the pattern of land surface phenology in terms of vegetative canopy structure and vegetation cover development (Figure 8b), which clearly shows that vegetation seasonality is highly responsive to the environmental conditions of the location (Figure 2).

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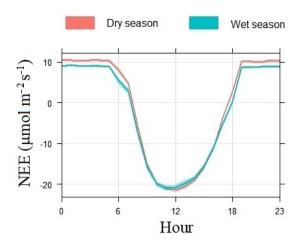
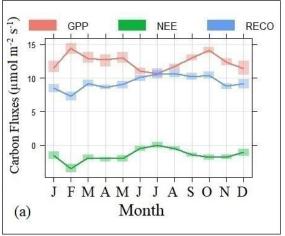


Figure 7. Average diurnal cycle of NEE at Caxiuanã, representing the CO₂ flux above the canopy, for the dry and rainy seasons. The hatched area indicates the 95% confidence interval.



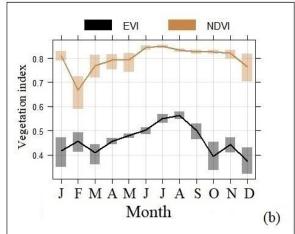


Figure 8. Monthly variation of NEE, GPP and RECO fluxes (**a**) and vegetation indices (**b**) at Caxiuanã. The hatched area indicates the 95% confidence interval.

At the site under study, NDVI and EVI are highest in August (NDVI > 0.8; EVI > 0.5), following the course of the site's greater radiative supply, decreasing in intensity with decreases in solar radiation and air temperature, showing levels of NDVI < 0.8 and EVI < 0.5, indicating the onset of senescence. The linear correlations between NDVI and EVI with measured GPP (Figure 9) are irrelevant because they showed values close to 0.

Figure 9 shows the correlation between the GPP measured at the micrometeorological tower and the vegetation indices (NDVI and EVI) during the study years, with an insignificant correlation between GPP and EVI ($R^2 = 0.01$) as well as with NDVI ($R^2 = 0.0046$). As for the tower GPP X MODIS GPP data (Figure 10), the correlation between the data is insignificant (Figure 10a), and the greatest similarity between the measurements occurs in the months of July and August (Figure 10b), both due to the proximity of the medians and the magnitude of the monthly variability.

In the other months, there is a greater underestimation of MODIS GPP and a seasonal signal that does not accompany the measured data. The best association of the data occurs in the months with the highest LAI (Figure 11), which shows values of order $> 6.0 \, \text{m}^{-2} \, \text{m}^{-2}$ and which are also the months with the lowest variability of the index, which shows a seasonal pattern that is highly related to the radiative availability of the site.

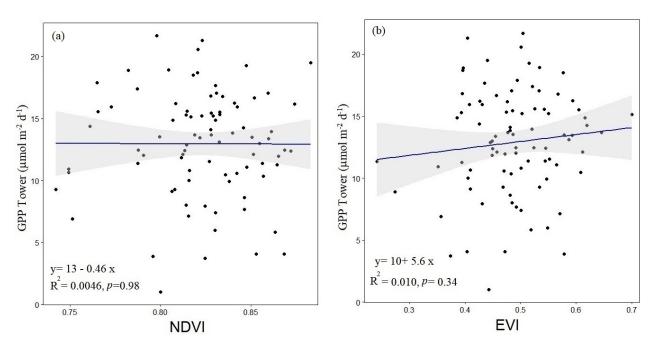


Figure 9. Pearson's linear correlation between the GPP data measured in the tower and the vegetation indices: NDVI (a) and EVI (b). The hatched area indicates the 95% confidence interval.

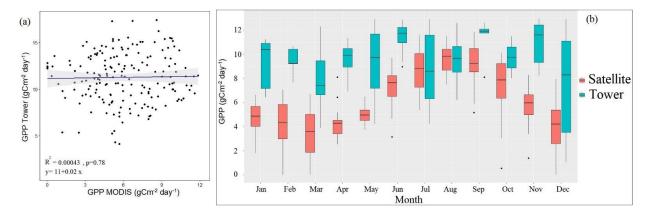


Figure 10. Comparison between measured gross primary production (Tower) and MODIS gross primary production (MODIS) in (**a**,**b**) Monthly GPP boxplot (gC m⁻² day⁻¹) for eddy covariance observed data (Tower) and MODIS—derived data (Satellite). The points refer to the outliers of the data, while the central lines refer to the median of each month.

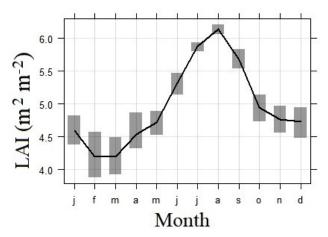


Figure 11. Monthly variation of LAI in the Caxiuanã site. The hatched area indicates the 95% confidence interval.

The carbon balance on the site (Figure 12a) is associated with local water availability (Figure 12b), where the greatest absorption (most negative NEE) occurred in 2006 ($-979~g~C~m^{-2}~year^{-1}$), which is later than the wettest year on the site (2005), which had an accumulated annual rainfall of 2880 mm year $^{-1}$. The year with the lowest absorption (least negative NEE) was 2008 ($-27.6~g~C~m^{-2}~year^{-1}$), which is after the year with the lowest annual rainfall (2105 mm year $^{-1}$) in 2007, indicating that the site's water availability can control the annual carbon balance in subsequent years.

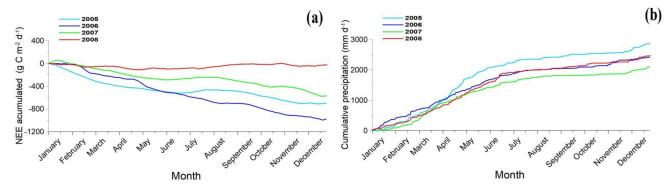


Figure 12. Annual carbon balance (**a**) and accumulated annual precipitation (**b**) at the Caxiuanã site for the years 2005 (aquamarine line), 2006 (blue line), 2007 (green line) and 2008 (red line).

4. Discussion

The meteorological patterns corroborate those in the same study region for other periods [16,49,55] and show the particularity of the site compared to others in the Amazon, such as in the region of Santarém-PA, which has greater distributions of air temperature density above 25 °C throughout the year [21] and has greater rainfall variability in the month of December. Seasonal changes in cloud cover are the main factors that determine the variation in solar radiation, even greater than the solar angle [69], which explains the increase in PPFD to the same extent as the decrease in precipitation.

Analyzing ET and precipitation, there is an increase in ET even in the less rainy period, which can be explained by the fact that primary forest is still a very humid environment and by the local floristic composition, with trees with deep root systems that seek water in the deeper layers of the soil, thus continuing the physiological activity of transpiration and maintaining the flow of water in the biosphere-atmosphere system. The seasonal variability that occurs in energy flows during the rainy and dry periods is noticeable, and this variability is explained by the high incidence of cloudiness that occurs in the Amazon basin, which is partly responsible for the increase in diffuse radiation and the decreases in radiation balance values [70], while during the dry period, it is characterized by less cloud presence, which results in an increase in the amount of net energy available to carry out the different processes that occur at the surface-atmosphere interface, reaching its maximum. The pattern of energy fluxes accompanying energy supply corroborates other studies [16,19,57] for this type of forest, unlike other areas where energy supply does not correspond to increases in other flows, such as LE, due to plant stomatal control mechanisms that cause the lowest ET values to occur at the end of the dry season [71]. For the Amazon rainforest, [72] also found distinct patterns of light dependence for the period before and after 11:30 a.m., but for other types of forest, even in the Amazon, the greatest assimilation occurs strictly in the daytime period [17].

Martínez et al. [73] showed that the coefficients of determination between NDVI and GPP are extremely low ($R^2 < 0.2$) for grasslands and evergreen broad-leaved forests, a pattern similar to the Caxiuanã forest. For many forest sites, NDVI is more sensitive to vegetation change because its mathematical formulation is based on the ratio of the difference in infrared and the sum of red reflectance and the sum of red and infrared reflectance, which is not the case in this environment even with EVI, which is an enhancement of NDVI

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because its mathematical formulation uses not only infrared and red reflectance but also surface cover and ground effects [61].

Zhang et al. [74] evaluated the spatial and temporal patterns of MODIS GPP at 40 sites located in the United States, comparing them with results from an ecosystem process model and with measurements of turbulent vortices (eddy covariance). It was found that there was an underestimation of MODIS GPP in highly productive months and an overestimation in low productivity months. This type of analysis is hampered in the study region and in almost the entire Amazon by the location's frequent cloud cover, which directly affects the images from optical sensors [75]. Although satellite remote sensing provides a viable and comprehensive means of observing environmental and phenological variations in the Amazon in a spatially wide-ranging and temporally frequent basis, the optical assessment of tropical areas is compromised by the high concentration of aerosols (biomass burning) and clouds [76–78], requiring the development of techniques for detecting clouds in images and improving estimates.

The authors suggest that the failure of the MODIS algorithm to represent the monthly and annual variability of GPP may be associated with the inadequacy of light use efficiency (LUE) models in capturing changes in canopy physiology over time. The tendencies in the MODIS GPP in relation to the observed data stem from a lack of sensitivity to canopy density, which is spatially and temporally variable. Other studies [79] evaluated the performance of GPP-MODIS in areas of the Brazilian Amazon, and they found that the MOD17 GPP product presented limitations in estimating the seasonality, magnitude and spatial variations of GPP in the tropical sites studied. When comparing MODIS GPP products with tower data in the Amazon, Nagai et al. [80] argue that the relationship between tower and image data is influenced by the satellite observation conditions and differences in scale, and despite this, it is possible to observe indices such as NDVI and EVI following the pattern of GPP increases from the beginning to the end of the dry season, suggesting changes in the structure and amount of water in the canopy seen by the sensor during the period of water stress and greater availability of radiation.

MOD17 overestimated GPP in low-productivity sites and underestimated it in high-productivity sites. Such results were also obtained in the literature [81,82]. The average annual carbon balance, $-569.7 \pm 444.9 \text{ g C m}^{-2} \text{ year}^{-1}$, is similar in magnitude to other tropical forest sites with high annual assimilation, as pointed out for the same site [17] in this study in a different period ($-560 \text{ g C m}^{-2} \text{ year}^{-1}$), sites like the Cuieiras [34] reserve ($-590 \text{ g C m}^{-2} \text{ year}^{-1}$) and a site [83] in Rondônia ($-400 \text{ g C m}^{-2} \text{ year}^{-1}$). There are studies that show strong correlations between MODIS data and flux tower data ($R^2 = 0.70$ in Temperate grassland, Zhu et al. [84]), while there are also sites where there is no or very weak correlation ($R^2 = 0.01$ in Floodplain forest, Costa et al. [7]; $R^2 = 0.16$ in tropical peatland, Wang et al., [85]; $R^2 = 0.17$ in Alpine Grassland, Zhu et al. [84]).

These differences between satellite and tower data correlation need to be further studied for each type of environment, then it will be possible to identify places where we have poor estimates so the scientific community can propose other tools and methods to improve these estimates to compensate for the lack of information for remote areas with low densities of measured data, such as the site under study. Viable options for solving this problem today range from changing the parameters used to calculate satellite GPP, which is often stationary throughout the year, to applying machine learning techniques to improve estimates.

Although other sites in Amazonia have been shown to be moderate sources of CO_2 to the atmosphere [36,57], the Caxiuanã site has been shown to be a considerable sink, with circulation patterns consistent with the literature [86], although it appears that precipitation can bring it to a state of neutrality in drier years. It is important to emphasize that there are many uncertainties associated with nocturnal respiration with gap filling by u^* at times of weak flows, which is true for practically all studies using the eddy covariance technique.

It is well known that the choice of u* threshold for gap-filling corrections can lead to indirect measurements due to data substitution during periods of weak turbulence, and they

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can even change the final results of a site from source to sink or sink to source [87]. However, gap-filling is accepted by the scientific community as necessary, and the automation of these gaps with valid data under similar environmental conditions as the filled data [88] provide confidence that the uncertainty of the data does not exceed acceptable percentage limits for a gap-fill, hence its widespread use by the scientific community [89].

5. Conclusions

The data show a forest area with well-defined seasonal patterns, with energy and water flows being controlled by radiative availability, and CO₂ seasonality patterns being controlled by the water available the previous year, probably due to the speed of the forest's physiological responses to environmental changes in a given year. In a climate-change context with an increase in the frequency of extreme weather events, this information is important to inform public policies on environmental conservation that can be adopted by the Brazilian government to mitigate these effects on this type of environment. We conclude that the site behaves as a CO₂ sink, like other parts of the Amazon rainforest reported in the literature, with latent heat flux dominating energy partitioning throughout the year and marked seasonality in meteorological variables. The satellite estimates show low correlation with the in situ data, as found in other similar sites in the literature, indicating the need for adjustments to the satellite parameters to better capture the seasonality of carbon exchange at the site. NDVI and EVI patterns follow radiative availability, as does LAI, but without direct capture related to GPP data, which correlates better with satellite data only in the months with the highest LAI. Carbon accumulations are shown to be directly related to the previous year's rainfall at the site, with the intensity of capture varying according to the increase or decrease in precipitation accumulations. In a global context where there is a great call to value the standing forest due to its ecosystem services and the great profits that the bioeconomy can provide to the communities that depend on the forest, the results of this work serve as a basis for showing the importance of conserving the environment studied here as a tool for mitigating climate change.

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