



Article Transformation of Soil Accumulated Phosphorus and Its Driving Factors across Chinese Cropping Systems

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Abstract: Understanding the transformation of accumulated phosphorus (P) is vital for P management. However, previous studies are limited to a few sites in Chinese agroecosystems. In this study, to investigate the temporal-spatial differences of transformation from accumulated P to available P (determined by the Olsen method), a dataset was assembled based on 91 national long-term experimental sites across China in the recent 31 years (1988-2018). A boosted regression tree (BRT) and a structural equation model (SEM) were used to analyze the factors influencing the transformation. The results showed that the transformation from accumulated P to available P in South China $(1.97 \text{ mg kg}^{-1})$ was significantly higher than that in other regions $(0.69-1.22 \text{ mg kg}^{-1})$. Soil properties were the main driving factors with a relative contribution of 81.8%, while climate and management practices explained 7.8% and 10.4% of the variations, respectively. Furthermore, SEM analysis revealed that the soil organic matter (SOM) could positively and directly affect the transformation, whereas the soil pH, soil silt content, and P fertilizer had negative and direct effects on it. For the first time, this study analyzed the transformation from soil accumulated P to available P at a national scale and at multiple sites and quantified the contribution of the main influencing factors. These results help to predict the soil available P content across different agroecosystems based on the input amount of P fertilizer, contributing to the regional precise management of P fertilizer application.

Keywords: soil Olsen-P content; soil P surplus; soil P transformation efficiency; driving factors; cropping systems

1. Introduction

Phosphorus (P), as an essential macronutrient element for plant growth, has an important nutritional and physiological function in agricultural systems [1]. However, the efficiency of P fertilizer utilization is often less than 25% mainly due to P transformation processes in soils including P adsorption/desorption and precipitation/dissolution [2]. In China, large amounts of fertilizer P has been increasingly applied to agricultural soils to maintain high crop yields [3,4]. However, the excessive P is fixed in soils as surplus P or lost via leaching and drainage runoff, which increases the risk of water eutrophication [5,6]. Based on the statistical data, the soil P surplus increased from 4.6 to 42.1 kg P ha⁻¹ yr⁻¹ between 1980 and 2012 in China's arable land [7–9]. This accumulated soil P has attracted global concerns over soil P management practices [10,11]. Thus, understanding the transformation of accumulated soil P is important for the improvement of regional fertilizer P management in Chinese cropping systems.



Citation: Chen, Y.; Guo, N.; He, W.; Abbasi, N.A.; Ren, Y.; Qu, X.; Zhang, S. Transformation of Soil Accumulated Phosphorus and Its Driving Factors across Chinese Cropping Systems. *Agronomy* **2023**, *13*, 949. https://doi.org/10.3390/ agronomy13040949

Academic Editor: Diego Pizzeghello

Received: 14 February 2023 Revised: 8 March 2023 Accepted: 17 March 2023 Published: 23 March 2023



Copyright: © 2023 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). Soil available P content transformed from accumulated soil P is one of the most important indicators of soil fertility and expected crop yields. Soil Olsen-P is one type of available P and Olsen-P content is commonly used to determine P fertilizer requirements and for environmental pollution control [12,13]. Previous studies indicated a linear correlation between the soil P surplus and soil P content [14–16], in which the slope of the linear regression indicates the transformation of soil Olsen-P from the accumulated P. The soil P transformation efficiency (SPTE, $mg \cdot kg^{-1}$) was often used to characterize the contribution of soil accumulated P to the increase in soil Olsen-P. In other words, the SPTE referred to the change of soil Olsen-P content when 100 kg P ha⁻¹ was surplus in soil. As a consequence, the SPTE value reflected the ability of transformation from accumulated P to Olsen-P [14,15]. The SPTE varied greatly among regions in China, which was mainly attributed to the differences in soil properties [16], climate conditions [17], and agronomic practices [18].

Soil properties (e.g., pH, soil organic matter (SOM)) have been often considered as the primary factors affecting the SPTE [15,18–20]. The effect of pH on the SPTE is mainly through its influence on the P precipitation reaction or dissolution of P associated with calcium (Ca-P), iron (Fe-P), aluminum (Al-P), and other P forms [21]. The SOM is positively correlated with the SPTE because high SOMs can enhance soil Olsen-P content by increasing microbial activity and competing for adsorption to metal (hydr)oxides [14]. Temperature and precipitation are two primary climate indicators [22] that indirectly affect the SPTE by regulating the microbial activity under the P cycle and soil properties [13]. Cao et al. [16] reported that the SPTE values in the middle and lower reaches of the Yangtze River and the southwest regions were high, which was mainly attributed to the high P activation caused by favorable pH (5.5–7.0), high temperature, and precipitation. Agronomic practices (e.g., P fertilizer application, land use, crop type, and tillage) could also affect the SPTE [23,24]. Shen et al. [17] reported that the SPTE under organic manure treatment was higher than that of the chemical fertilizer treatment, which was partially due to the increase in SOC. Extra organic C could decrease the soil fixation of Fe-P and release P into the soil by low-molecular-weight organic acids [25]. Thus, the factors affecting the SPTE are highly dependent on influencing (directly, indirectly, or interactively) the turnover process of the soil P [16]. Therefore, exploring the temporal-spatial differences of the SPTE and its driving factors is crucial for developing appropriate and efficient P nutrient management practices and reducing environmental pollution in China's croplands.

As outlined above, previous studies of SPTE were mainly based on simple qualitative analyses or a limited number of factors on a farm-size scale and over a short time period. The quantification of the contribution and interaction of these factors to the SPTE remains to be elucidated at regional and national scales. The first aim of this study was to identify the SPTE in different regions of China. The second aim was to explore the effects of climate, agricultural practices, and soil properties on the transformation of accumulated soil P. We hypothesized that the SPTE in the south region is higher than the north region. Furthermore, climate, agricultural practices, and soil properties were hypothesized to jointly effect SPTE, but soil properties play the most prominent role. To test these hypotheses, we systematically investigated the relationship between the change of soil Olsen-P and accumulated soil P in the past 31 years based on 91 national long-term monitoring sites in China. This study will provide a framework for the prediction of Olsen-P contents and the formulation of various P fertilization schemes in different regions of China.

2. Materials and Methods

2.1. Experimental Sites and Designs

The data used in this study were based on the long-term monitoring network of the Ministry of Agriculture in rural areas, with a collection over 31 years (1988–2018). A total of 91 experimental sites were selected in this study (Figure 1). Among them, 16 sites were established in 1988, 32 in 1998, and 43 in 2004. To compare the SPTE in different regions, 91 experimental sites covering 23 provinces were divided into 6 regions based on geograph-

ical locations and China's administrative divisions [26], including northeast China (NE, n = 7, containing Jilin and Heilongjiang Provinces), the Huang-Huai-Hai region (HH, n = 18, Hebei, Shandong, and Henan Provinces), the middle and lower reaches of the Yangtze River (YR, n = 38, Jiangsu, Anhui, Fujian, Hubei, Hunan, and Guangxi Provinces), South China (SC, n = 6, Guangdong and Hainan Provinces), southwest China (SW, n = 7, Sichuan, Guizhou, Yunnan, and Chongqing Provinces), and northwest China (NW, n = 15, Xinjiang, Ningxia, Gansu, Shanxi, Shaanxi, and Qinghai Provinces).



Figure 1. Distribution of long-term fertilization experiment sites in this study (n = 91). The regional abbreviations shown are NE: northeast China; NW: northwest China; HH: Huang-Huai-Hai region; YR: middle and lower reaches of the Yangtze River; SC: South China and SW; southwest China.

The basic soil physicochemical properties of each monitoring site are shown in Table 1. The 91 experimental sites are distributed in four climatic zones: the middle temperate zone, the warm temperate zone, the subtropical zone, and the tropical zone. The planting systems include mono-cropping, double-cropping, and triple-cropping. The land-use types included paddy, upland, and paddy-upland. Wheat, corn, and rice were the main crops at most sites, while soybean, rape, cotton, green manure crops, tomato, celery, cowpea, peanut, pepper, and potato were planted at others. The designs of all experimental sites were the same, including the control (without fertilizer) and conventional fertilizer treatment (farmers' fertilization practice). The area of each plot was no less than 334 m². Except for the fertilizer treatment as mentioned above, all other agronomic practices were identical between the two treatments. The amount of fertilizer P applied is shown in Table 2. The drought index (*DI*) was calculated as follows [27]:

$$DI = \frac{Annual \ precipitation}{Annual \ temperature + 10} \tag{1}$$

Region	рН	SOM ¹ (g kg ⁻¹)	Total Nitrogen (g kg ⁻¹)	Soil Olsen-P (mg kg ⁻¹)	Available Potassium (mg kg ⁻¹)	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	DI
NE ²	6.40 ± 0.10 c 3	$33.85\pm1.74b$	$1.86\pm0.05~\mathrm{a}$	25.05 ± 4.17 a	190.46 ± 14.74 a	$3.31\pm0.55~\mathrm{e}$	$542.73 \pm 14.88 \text{ d}$	$41.25\pm5.56~\mathrm{a}$
HH	$7.20\pm0.26\mathrm{b}$	$14.38\pm0.89~\mathrm{d}$	$0.96\pm0.05~\mathrm{b}$	23.50 ± 2.59 a	$108.02\pm7.63~\mathrm{ab}$	$13.97 \pm 0.25 \mathrm{c}$	$722.34 \pm 36.62 \text{ c}$	$30.13\pm6.00~\mathrm{b}$
YR	$6.44\pm0.13~{ m c}$	$26.11 \pm 1.70 \text{ c}$	$1.60\pm0.09~\mathrm{ab}$	19.01 ± 2.22 a	$109.87 \pm 6.32 \text{ ab}$	$16.51\pm0.41\mathrm{b}$	$1165.56 \pm 48.93 \mathrm{b}$	$43.69 \pm 9.03 \text{ a}$
SC	$5.60 \pm 0.25 \text{ d}$	$42.57\pm7.29~\mathrm{a}$	$2.17\pm0.30~\mathrm{a}$	$26.20\pm4.57~\mathrm{a}$	$70.56 \pm 8.05 \mathrm{b}$	$22.23\pm0.77~\mathrm{a}$	1742.28 ± 93.05 a	54.24 ± 7.33 a
SW	$6.44\pm0.40~{ m bc}$	33.64 ± 8.44 ab	1.81 ± 0.31 a	27.73 ± 9.07 a	127.74 ± 16.21 a	$16.73 \pm 0.55 \mathrm{b}$	$1068.41 \pm 58.04 \mathrm{b}$	40.16 ± 6.48 a
NW	$8.24\pm0.07~\mathrm{a}$	$15.37 \pm 1.16 \text{ d}$	$0.94\pm0.06~{ m b}$	22.86 ± 3.26 a	166.79 ± 14.40 a	$8.43 \pm 0.66 \text{ d}$	$313.41 \pm 52.34 \text{ e}$	$16.70 \pm 9.34 \text{ c}$

Table 1. Summary of experimental site information at different regions.

¹ SOM: soil organic matter, $g \cdot kg^{-1}$. ² Abbreviations for regions: NE: northeast China, HH: Huang-Huai-Hai region, YR: middle and lower reaches of the Yangtze River, SC: South China, SW: southwest China, NW: northwest China. ³ Values given represent the mean values \pm standard errors. Different lowercase letters indicate significant differences among regions in the same column (p < 0.05).

Table 2. Fertilizer P information at different regions.

		Annua	l P Fertilizer Input	(kg ha $^{-1}$)	P Removal	Annual P	SPTE
Region	Main Crop	O-P ¹	C-P	T-P	(kg ha ⁻¹)	surplus (kg ha ⁻¹)	(mg kg ⁻¹)
NE ²	Corn, rice, soybean Wheat, corn, rice, soybean, rape,	5.34 ± 2.0 d 3	$79.66 \pm 5.29 \text{ b}$	$85.01\pm5.47~\mathrm{c}$	$59.20 \pm 6.88 \text{ ab}$	$25.81\pm8.37~\mathrm{c}$	$1.22\pm0.34b$
HH	cotton, peanut, green manure crops, pepper	$33.23\pm6.56~ab$	134.27 ± 7.67 a	167.51 ± 8.25 a	67.32 ± 6.59 a	$100.19 \pm 8.50 \text{ ab}$	$0.91\pm0.16~\text{b}$
	Wheat, corn, rice,						
YR	soybean, rape, green manure	$11.49\pm1.91~\mathrm{c}$	$120.62\pm4.20~\text{a}$	$132.12\pm4.66b$	$68.29\pm3.66~\text{a}$	$63.82\pm5.66~b$	$1.04\pm0.13b$
SC	crops, pepper Rice Wheat, corn, rice,	$10.12\pm2.76~\mathrm{c}$	118.23 ± 18.28 a	$128.35\pm20.46~\mathrm{bc}$	$75.29\pm6.47~\mathrm{a}$	$53.06\pm21.87~\mathrm{b}$	$1.97\pm0.57~\mathrm{a}$
SW	green manure crops	$24.41\pm6.57~b$	$85.33\pm13.84b$	$109.74\pm19.57\mathrm{bc}$	$29.10\pm10.06~\mathrm{c}$	$80.64\pm9.81~b$	$1.12\pm0.24~b$
NW	Wheat, corn, rice, rape, green manure crops, tomato, celery, cowpea	51.2 ± 12.40 a	130.77 ± 10.13 a	181.98 ± 20.09 a	$49.20\pm6.03~\text{b}$	132.78 ± 22.49 a	$0.87\pm0.12~\mathrm{b}$

¹ O-P: Organic fertilizer P; C-P: Chemical fertilizer P; T-P: Total fertilizer P, the sum of organic fertilizer P and chemical fertilizer P. ² Abbreviations for regions are in Table 1. ³ Values given represent the mean values \pm standard errors. Different lowercase letters indicate significant differences among regions in the same column (p < 0.05).

The DI, crop type, fertilizer dosage, soil nutrient content, field management information, and yield were recorded in detail at all experimental sites (Tables A1 and A2).

2.2. Soil Sample Collection and Analysis

The soil samples (0–20 cm) of 3 replicates in every experiment site were taken in the autumn every year after harvest and before the fertilizer application. These samples were air-dried and sieved through a 2-mm sieve. The soil pH was measured with a soil/distilled water ratio of 1:2.5. The SOM content was measured by potassium dichromate oxidation (0.167 mol L⁻¹ K₂Cr₂O₇) method. The total nitrogen was extracted and determined by the Kjeldahl method [28]. The concentration of soil Olsen P was measured by the ascorbic acid method with 0.5 M NaHCO₃ (soil/solution ratio of 1:20) [29]. The available potassium was extracted with 1.0 M NH₄OAc, according to Lu [28]. The grain and straw samples of all crops were collected artificially, dried at 105 °C for 2 h, and then at 80 °C for 72 h. The weight of plants was measured. The plant concentration was measured by the molyb-dovanado phosphate method after being digested in concentrated H₂SO₄ and H₂O₂ [28]. The SoilGrids system (https://soilgrids.org, accessed on 20 March 2022) was used to derive the clay, silt, and sand contents in the 0–15 cm surface layer based on the site location information (longitude and latitude), with a spatial resolution of 250 m [30]. The mean annual temperature and mean annual precipitation data were calculated by collecting

meteorological data in China (http://cma.gov.cn/, accessed on 20 March 2022). The soil type was classified as a Eutric Cambisol [31].

In this study, the surplus of soil P was the difference between the P applied as fertilizer (chemical and organic P fertilizer) to the soil surface and the P uptake by crops. The runoff and infiltration of P were relatively small and ignored in the calculations [14]. The equation for the soil P surplus calculation is as follows:

Soil P surplus =
$$PF - PC$$
 (2)

where PF is the P applied through P fertilizer (kg ha^{-1}), and PC is the P removed by crops (kg ha^{-1}). PC is calculated from:

$$PC = YG \times CG + YS \times CS$$
(3)

where YG is the grain yield (kg ha⁻¹), CG is the P content in grain (%), YS is the weight of straw (kg ha⁻¹), and CS is the P content in straw (%).

$$\triangle$$
 Soil P surplus = \sum_{1}^{i} Soil P surplus (4)

where \triangle Soil P surplus (kg ha⁻¹) is the sum of the apparent P surplus and the deficit during the monitoring period.

The change of soil Olsen-P (\triangle AP, mg kg⁻¹) refers to the change of test P from the beginning to the end of the monitoring period.

$$\triangle AP = APi - AP0 \tag{5}$$

where APi is the soil Olsen-P content in the end year of the monitoring period, and AP0 is the soil Olsen-P content in the first year of the monitoring period.

 \triangle AP and \triangle Soil P surplus of each experimental site were fitted, and the slope of the line was the SPTE (mg kg⁻¹), which refers to the change of soil Olsen-P content caused by 100 kg P ha⁻¹ in surplus [15,16].

The Olsen-P content in the future year can be predicted by the current available P content, plus the increase in available P caused by P surplus. Assuming that the apparent P surplus maintains the current value, the accumulated P surplus in any year of future can be calculated. Since the increase in available P caused by the P surplus of 100 kg ha⁻¹ is known as the SPTE, the increase in available P can be obtained according to the accumulative P surplus [15].

2.3. Statistical Analysis

To explore the temporal-spatial variations and avoid the influence of outliers in soil Olsen-P content and soil P surplus, the data were divided into six stages based on every five experimental years. The situ sets were 1988–1992 (5 years), 1993–1997 (10 years), 1998–2002 (15 years), 2003–2007 (20 years), 2008–2012 (25 years), and 2013–2018 (30 years) for 6 stages, respectively. The temporal-spatial patterns of soil Olsen-P and soil P surplus were obtained using linear regression with the least squares method in the software Origin (version 9.0). The spatial distribution of monitored sites was using the GPS coordinates and the distribution of the SPTE were based on the spatial resolution of the SPTE concentration range in the Arc-GIS 10.4.

Based on the theoretical knowledge and the existing database, a total of nine factors (DI, P fertilizer rate, land use, crop rotation, SOM, soil pH, sand, silt, and clay content) were considered to explore their relationships with SPTE. The boosted regression tree model (BRT) was used to explore the effect of climate, agricultural practices, and soil properties on the response variable (SPTE). In the BRT process, the recommended parameters (learning rate (0.01), bag fraction (0.50), cross-validation (10), and tree complexity (5)) were used in this study [32]. Because there were alphabetic variables in environment variables, the

Bernoulli method was adopted for determining BRT. The relative importance of each predictor indicated a percentage of the total variations of the model in the dependent variable. The BRT analyses was executed in the GBM package of R version 3.3.3. Global Moran's I statistic was used to evaluate the spatial structure of residuals in BRT using spdep package version 0.6-8 [33].

Structural equation modeling (SEM) was used to explore the relationship between the external environmental factors and SPTE. In the development of the initial path model, the following paths were assumed. Firstly, the five selected factors (DI, P fertilizer dosage, soil silt content, soil pH, SOM) had a direct impact on the SPTE. Second, the DI and P fertilizer dosage had a direct effect on the soil properties (soil pH, SOM). Third, the soil silt content might indirectly affect the SPTE by affecting the soil pH and SOM. Finally, all significant correlation paths were retained in the framework (Figure A1). SEM analysis was based on the acceptance of the zero hypotheses ($p \le 0.05$). The comparison between model-implied variance and covariance matrix was conducted using the maximum likelihood estimation method with software AMOS 17.0 (IBM SPSS Inc., Chicago, IL, USA). The overall goodness-of-fit of the model was evaluated with the *p*-value (0.05 < *p* < 1.00 indicates no statistical difference between the covariance matrices produced by the model fits and the observed covariance matrices) and the root mean square error of approximation (RMSEA) (<0.05 indicates an ideal fit) [34].

3. Results

3.1. Temporal-Spatial Changes in Soil Olsen-P and P Surplus

The soil Olsen-P content was significantly increased (p < 0.01) in the recent 31 years (Figures 2a and A2a). The average soil Olsen-P content in 2013–2018 reached 28.92 mg kg⁻¹, which was 86.0% higher than that in 1988–1992 (15.56 mg kg⁻¹). The average growth rate was 0.45 mg kg⁻¹ yr⁻¹. The soil Olsen-P content in South China and southwest China was significantly higher than those in other regions in the same period (Figure 3). The increased rate of soil Olsen-P content (mg kg⁻¹ yr⁻¹) in different regions is ordered as: southwest China (0.99) > Huang-Huai-Hai region (0.68) > middle and lower reaches of the Yangtze River (0.56) > South China (0.41) > northeast China (0.29) > northwest China (0.28).



Figure 2. Soil Olsen-P content (**a**) and soil P surplus (**b**) over time under 91 long-term fertilization experiments from 1988 to 2018. The black line and dot, lower and upper edges and bars represent median and mean values, 25th and 75th, and 5th and 95th percentiles of all data, respectively. Probability levels < 0.05 were considered as statistically significant.



Figure 3. The temporal-spatial patterns of soil Olsen-P content and apparent soil P surplus across Chinese cropping systems from 1988 to 2018. Regional abbreviations are as in Figure 1. Probability levels < 0.05 were considered as statistically significant.

The soil apparent P surplus showed a significant decreased trend (p < 0.05) in the recent 31 years at the national scale (Figures 2b and A2b), while the accumulated P surplus significantly increased (Figure A3). The soil apparent P surplus varied greatly among regions. It was higher in northwest China and the Huang-Huai-Hai region compared with other areas (Figure 3, Table 2). Furthermore, the decreased rates in northeast China, northwest China, and Huang-Huai-Hai region were -1.99, -8.44, and -3.02 kg ha⁻¹ yr⁻¹, respectively, whereas the soil apparent P surplus in the middle and lower reaches of the Yangtze River and southwest China regions fist increased and then decreased.

3.2. Temporal-Spatial Variation of SPTE and Its Driving Factors

There was significant spatial heterogeneity in the SPTE (Figure 4a). The SPTE and drought index showed a significant positive linear correlation (p < 0.01, Figure 4b). The SPTE in South China (1.97 mg kg⁻¹) was significantly higher than those in northeast China (1.22 mg kg⁻¹), southwest China (1.12 mg kg⁻¹), the middle and lower reaches of the Yangtze River (1.04 mg kg⁻¹), the Huang-Huai-Hai region (0.99 mg kg⁻¹), and northwest China (0.69 mg kg⁻¹) (Figure 4c).



Figure 4. The spatial pattern of national SPTE (**a**), the relationship between SPTE and drought index (**b**), and SPTE at regional scale (**c**) across Chinese cropping systems. Regional abbreviations are as in Figure 1.

The highest SPTE (1.36 mg kg⁻¹) was observed when the soil pH was in the range of 5.5–6.5 compared to other pH ranges of 4.5–5.5, 6.5–7.5, and 7.5–8.5 (Figure 5a). The SPTE (1.65 mg kg⁻¹) was significantly higher when the SOM was more than 30 g kg⁻¹, compared with the cases when the SOM was between 15 and 30 g kg⁻¹ or lower than 15 g kg⁻¹ (Figure 5b). The SPTE of clay soils (1.44 mg kg⁻¹) was significantly higher compared with sandy and silty soils (Figure 5c). For agricultural practices, the SPTE (1.38 mg kg⁻¹) in the paddy field was substantially higher than the upland and paddy upland rotation (Figure 5d). There was no significant difference in the SPTE among the different cropping systems (mono-cropping, double-cropping, and triple-cropping) (Figure 5e).



Figure 5. The differences of SPTE under various soil pH values (**a**), SOMs (**b**), soil textures (**c**), land use types (**d**), and cropping patterns (**e**). P, U, and P-U are paddy, upland, and paddy-upland, respectively. MC, DC, and TC are mono-cropping, double-cropping, and triple-cropping, respectively. Different lowercase letters indicate significant differences (p < 0.05) under various soil pH, SOM, soil texture, land use types, and cropping patterns.

The BRT results showed that the climate (drought index), agricultural practices (P fertilizer, land use, and cropping patterns), and soil properties (SOM, soil pH, and soil texture) jointly drive the variation of the SPTE (Figure 6a). The soil properties were the main driving factors, with a relative contribution of 81.8%, while the climate and agronomic



practices accounted for 7.8% and 10.4% of the variation, respectively. The BRT model driven by the above variables explained 71% of the variation of the SPTE (Figure 6b).

Figure 6. The relative contributions (%) of climate (drought index), soil properties (soil pH, SOM, and soil texture) and management practices (P fertilizer input, land use types, and cropping patterns) to SPTE, estimated by the boosted regression tree (BRT) model (**a**). The observed SPTE and those predicted by the boosted regression tree model using various predictors are shown in (**b**). The dotted line indicates the 1:1 line.

Through SEM analysis, we constructed different pathways of the aforementioned variables on the variation of the SPTE (Figure 7). The soil pH, soil silt content, and P fertilizer amount had negative and direct effects on the SPTE, with path coefficients of -0.42, -0.27, and -0.25, respectively. SOM could positively and directly affect the SPTE, and the path coefficients were 0.43. The drought index could indirectly affect the SPTE by regulating the SOM, silt content, and pH. The P fertilizer amount also indirectly affected the SPTE by regulating SOM. In general, soil properties, agricultural practices, and climate together a counted for 64% of the variation of SPTE.



Figure 7. Path analysis results on the direct and indirect effects of DI, soil pH, SOM, soil silt content, and P fertilizer input on SPTE in 91 long-term fertilization experimental sites across Chinese cropping systems (chi/df = 1.36, P = 0.43; RMSEA = 0.05). DI is drought index. The number indicates the path coefficients. The solid and dashed lines indicate significant positive and negative effects, respectively.

4. Discussion

4.1. Temporal-Spatial Differences of Soil Olsen-P and P Surplus

Our study showed an increasing trend in the Olsen-P content with a slope of 0.45 mg kg⁻¹ yr⁻¹ (0.29–0.99 mg kg⁻¹ yr⁻¹) based on a linear model from 1988 to 2018 in Chinese croplands, which was mainly related to the high input of P fertilizer in recent years. Our results were similar to the findings in previous studies [5,8,9], which showed an annual increase rate of 0.17–1.09 mg kg⁻¹ in the soil Olsen-P content during the past 30 years. In comparison with other regions, the soil Olsen-P content rapidly increased in southwest China mainly due to the higher fertilizer input and lower crop P removal [9]. There was a low increase in the Olsen-P content over time in northwest region and northeast region, which was partially attributed to the low temperature, high pH, and low SOC content [16] (Table A1).

The decreasing trend in the apparent P surplus in our study was highly relevant to the lower P fertilizer (Figure A4) and the increased crop P uptake due to improved yield (Figure A5). What is more, the fertilizer P utilization efficiency, resulting from the improvements in fertilization technology, an increased use of more efficient mechanized fertilization methods, and the increased application of organic P fertilizer [8,35,36], played important roles. The apparent P surplus in the northeast, northwest, and Huang-Huai-Hai regions decreased over time, whereas it increased first and then decreased in the middle and lower reaches of the Yangtze River and southwest regions. The spatial differences of the soil P surplus were mainly related to the regional differences in fertilizer input, soil and climate conditions, and economic development [7,9] (Tables A1 and A2).

4.2. SPTE and Its Influencing Factors

The relationship of the P surplus and Olsen-P content was defined as the SPTE in our study for evaluating P use efficiency and providing a basis for P management. In the recent 31 years, the SPTE value of South China (1.97 mg kg⁻¹) was significantly higher than in other regions (0.69–1.22 mg kg $^{-1}$). Our results of SPTE were lower than the values $(1.44-16.04 \text{ mg kg}^{-1})$ in previous studies [14,16,17,37]. This is partially due to the fact that the SPTE values were related to P fertilizer management in these earlier studies, where the order of influence was: chemical P fertilizer combined with organic fertilizer > chemical P fertilizer > no P fertilizer [14,16,17]. In our study, the chemical P fertilizer application was the dominant fertilizer type (Table 2), while the previous studies were mostly based on combined organic-inorganic fertilization. In addition, the difference of the SPTE may be related to P test methods. For example, the SPTE value was 0.16 mg L^{-1} in Messiga et al. [38], based on water-soluble P in Canada, which was significantly lower than that obtained from the Olsen-P method. Another possibility is that the P runoff may occur at the slopping sites, especially those with heavy rainfall during the growing season. Thus, the P surplus may be overestimated in these regions, which leads to an underestimated SPTE.

In the present study, soil properties have been identified as the main driving factors on SPTE, which is consistent with previous studies [17,39]. The SOM and pH explained 33.6% and 31.6% of the variation in the SPTE, respectively (Figure 6a). The SOM was positively correlated with the SPTE, which may be due to SOM reductions in P adsorption to competing adsorption sites by its provision of alternative organic anions and/or a SOM contribution to mineral-P dissolution by its content low-molecular-weight organic acids [40]. Thus, the "promoting P with carbon" technique could be adopted for soils with a low SOM content by increasing the input of organic fertilizers, biochar, or similar practices [41,42]. In contrast, the amount of P fertilizer should be reduced in the soils that contains high SOM content to avoid P leaching and runoff [43]. The effect of soil pH on the SPTE mainly depends on the chemical adsorption and precipitation/dissolution process of P [44]. Based on our results, the highest SPTE was observed in the pH of 5.5–6.5, which was found in the soils of South China (Table 1), and then the SPTE decreased when soil pH changed from 5.5–6.5 to 4.5–5.5, which was partially due to the increase in fixation by iron [45,46]. Additionally, the SPTE decreased when the soil pH changed from 6.5 to 8.5. This could be attributed to the increased fixation by calcium [21]. Therefore, the application of acidic fertilizers or conditioners (e.g., low molecular weight organic acids, humic acids) could be feasible practices to reduce the pH and increase the SPTE in the northern alkaline regions of China [44,47], whereas the addition of lime has been recommended as a means to increase the SPTE in southern regions with low soil pH [48]. Through the above measures, the SPTE and P use efficiency were increased, culminating in reductions in the P fertilizer input and P surplus.

Climate had an impact on the SPTE (contributes 7.8% of the variation), which is consistent with the findings in Shen et al. [17] and Hou et al. [22]. The drought index was used to combine the two climate indicators of temperature and rainfall, which is more conducive to the formulation of guidelines for P fertilizer usage in different regional climates [49]. The drought index was seen to indirectly influence the SPTE by regulating SOM, soil texture, and pH [22,50].

The fertilizer P application had a direct negative effect on the SPTE and an indirect pathway via its negative effect on SOM (Figure 7). By contrast, previous studies have reported a positive correlation between organic P fertilizer usage and SOM content [16,43]. This can possibly be attributed to the focus of our study on the impacts of the long-term application of chemical P fertilizers on SPTE. Continuous chemical P application may significantly decrease the activity of microorganisms, the concentrations of microbial residues, and their contributions to SOM accumulation [51].

The highest SPTE (1.97 mg kg⁻¹) in our study was found in South China soils (Figure 4a,c), which may be attributed to the high SOM content, low pH, low silt content, and high drought index (Tables 1, A1 and A2). Although the BRT and SEM models showed that the interpretation rates of the influencing factors on the SPTE were 71% and 64%, respectively, other important factors including soil sampling depth, soil tillage, topography, and oxalate-extractable Fe/Al (Fe/Al oxides) should be considered in future work [52,53]. For example, the soil sampling depth and soil tillage can also affect SPTE. The soil sampling depth in this research was 0–20 cm for soil Olsen-P determination. Messiga et al. [54] reported that the SPTE at 0–5 cm, 0–10 cm, and 0–15 cm was higher than that in 0–20 cm samples. Compared with tilling by the moldboard plough, no-tillage is more conducive to the accumulation of Mehlich-3 P and Olsen-P [52]. Stratification produces high concentrations of P at the soil surface (0–5 cm) but decreases the concentrations in deeper soil layers [55]. Topography influences the soil available P and SPTE because it affects the runoff, drainage, soil temperature, soil erosion, and soil formation [56]. Additionally, oxalate-extractable Fe/Al (Fe/Al oxides) have been proven to generally play a dominant role in controlling the dynamics of phosphate ions in acidic and non-acidic soils [57]. In BRT and SEM analysis, we integrated the influencing factors of the SPTE across regions. Thus, the main driving factors that affected the mechanism of soil P process were concluded at the national level, but they might be different among regions. This needs to be further investigated when more data are available at the regional level.

4.3. Recommendation of P Fertilizer Based on SPTE

The goal of P management is to ensure an optimum crop yield without causing environmental risks by keeping the Olsen-P content in the range between the agronomic threshold and environmental threshold [12]. Several studies have evaluated the dose and rate of P fertilizer usage in China, mainly based on high yield and high P use efficiency [8], changes in soil Olsen-P [16], or a combination of changes in Olsen-P and agronomic threshold [15]. In this study, the SPTE was used to estimate the soil Olsen-P and establish recommendations for P fertilizer usage based on the agronomic and environmental thresholds. Previous studies showed that the agronomic threshold varied in crops ranging from 11.1 to 28.2 mg kg⁻¹, and the environmental threshold was 40–52 mg kg⁻¹ in Chinese agricultural systems [12,13,58]. In this study, the Olsen-P contents in most regions exceeded the agronomic threshold (Table A3). For example, during the period 2013–2018,

the South China region had a high soil Olsen-P content of 37.39 mg kg⁻¹ and a high SPTE of 1.97 mg kg⁻¹ (Figure 3, Table 2), which was close to the environmental threshold [12]. Our results were consistent with Zhou et al. [59], which pointed out that uplands in eastern and southern China had a high risk for elevated soil P loss, particularly, in Guangdong, Fujian, and Zhejiang provinces. The high soil Olsen-P content increased the risk of water eutrophication, which would damage the fishing industry and may pose a serious health hazard to humans [60].

Thus, more attention to the environmental threshold of Olsen-P is required to prevent environmental pollution. The future Olsen-P content in soils can be predicted based on the SPTE, the current P application rate, and the P surplus. In this study, the soil Olsen-P contents in South China and the southwest China regions were relatively high and are predicted to reach the environmental threshold in 4.6–16.1 and 7.5–34.1 years, respectively (Table A3), if the current P fertilizer practices are maintained. Xi et al. [13] also reported that the soil Olsen-P content in northern China would reach the environmental critical P value in 5–41 years, depending on the fertilizer P management procedure utilized. Although the time to reach the critical value is given in this and other studies, such boundaries are hard to predict, and the negative effects may become manifest before the thresholds are crossed. Once the soil Olsen-P content exceeds the environmental threshold, it takes a longer time to return below the environmental threshold [11,61]. Therefore, it is urgent to improve P use efficiency and reduce the fertilizer P amount in the regions which have high potential environmental risks. Overall, the SPTE driving factors based on 91 long-term experimental sites across China, which covers various soil types, climate conditions, and cropping systems, could provide a reference for fertilizer P management in other regions or countries.

5. Conclusions

Based on 91 long-term monitoring sites across Chinese cropping systems, the temporalspatial patterns of the soil P transformation efficiency (SPTE) were investigated. The SPTEs presented significant regional differences and the highest value of SPTE was found in South China. Based on the BRT and SEM models, thDistribution of long-term fertilization e study confirmed that soil properties, agronomic practices, and climate comprehensively affect SPTE, among which soil properties were the main driving factors with a relative contribution of 81.8%. With regard to the soil properties, the SOM and pH were two vital indicators influencing SPTE, which were mainly due to their impacts on P adsorption/desorption and precipitation/dissolution. These results suggest that the highest SPTE of South China may be attributed to the high SOM, low pH, low silt content, and high drought index. Therefore, for different regions of the world, the SOM and soil pH should be considered firstly for P fertilizer recommendation. In this study, the experimental sites' distribution did not cover all provinces, which may lead to the uncertainty in the SPTE predication when scaling up from site to region. Additional influencing factors on the SPTE need to be further investigated, including soil tillage, sampling depth, topography, and oxalate-extractable Fe/Al (Fe/Al oxides) in a long-term period. This study suggests that the SPTE can be used to estimate the time when each region will reach the agronomic threshold or environmental threshold of Olsen-P. The SPTE driving factors in this study included various soil types, climate conditions, and cropping systems, which could provide a useful reference for improving P management practices to other regions or countries.

Author Contributions: Conceptualization, S.Z.; methodology, Y.C.; software, W.H.; validation, N.A.A., Y.R. and X.Q.; formal analysis, Y.C.; investigation, N.G.; resources, Y.R. and X.Q.; data curation, S.Z.; writing—original draft preparation, Y.C.; writing—review and editing, Y.C. and N.G.; visualization, W.H.; supervision, S.Z.; project administration, S.Z.; funding acquisition, S.Z. All authors have read and agreed to the published version of the manuscript.

Funding: This research was funded by the National Key Research and Development Program of China (2021YFD1500205) and the National Natural Science Foundation of China (41977103).

Institutional Review Board Statement: Not applicable.

Informed Consent Statement: Not applicable.

Data Availability Statement: The data sets generated and/or analyzed during the current study are available from the corresponding author upon reasonable request.

Acknowledgments: The authors would like to thank the financial support by the National Key Research and Development Program of China (2021YFD1500205) and the National Natural Science Foundation of China (41977103).

Conflicts of Interest: The authors declare no conflict of interest.

Appendix A

Table A1. Detailed information of field experimental sites.

Site			· · ·		Yield of Main Crops (kg ha ⁻¹)				:	Soil Proj	perties			
Number	Region	Province	Longitude	Latitude	Wheat	Maize	Rice	pН	SOM ¹ (g/kg)	AK ² (mg/kg)	Clay (%)	Silt (%)	Sand (%)	Soil Texture
1	NE ³	Jilin	126.5	44.8		8940		6.57	30.56		25	50	25	Loam
2	NE	Jilin	129.1	42.8			7350	5.93	40.48	157.07	20	41	39	Loam
3	NE	Heilongjiang	126.8	46		9600		6.64	33.75	265.02	31	53	16	Silty clay loam
4	NE	Heilongjiang	126.3	45.3		9780		6.73	27.63	174.84	26	42	32	Loam
5	NE	Heilongjiang	127	47			7725	6.35	36.97	184.69	32	49	19	Silty clay loam
6	NE	Heilongjiang	126.8	46.7			8100	6.29	37.29	165.37	32	49	19	Silty clay loam
7	NE	Heilongjiang	129.7	46.2		8970	7380	6.31	30.29	195.77	29.6	41	29.4	Clay loam
8	HH	Hebei	114.4	37.2	4695	6375		8.06	16.81	98.43	20	43	37	Loam
9	HH	Hebei	114.3	38.1	6075	7095		8.00	20.63	123.80	17	47	36	Loam
10	HH	Shandong	117.1	36.5	6720	9195		7.76	15.58	137.55	22	47	15	Loam
11	HH	Shandong	115.5	35.3	6000	6945		8.04	12.93	127.83	19	50	31	Loam
12	HH	Shandong	120.4	36.8	6510	7950		7.04	14.42	114.64	22.3	33.1	44.6	Loam
13	HH	Shandong	119.8	37.2	2880	4650		5.51	8.72	72.90	23	35.2	41.8	Loam
14	HH	Shandong	120	37.1	2940	5220		5.22	8.67	70.01	23	35.2	41.8	Loam
15	HH	Henan	112.3	32.6	6555	6840		6.68	15.44	126.45	29	44	27	Clay loam
16	HH	Henan	114.1	31.2			6795	5.85	21.32	92.34	30	52	18	Silty clay loam
17	HH	Henan	114.5	32.4	5850			5.48	14.14	70.42	32.3	50.4	17.3	Silty clay loam
18	HH	Henan	112.8	34.6	4965	5475		7.97	15.67	134.30	19.7	48.3	32	Loam
19	HH	Henan	114.6	33.5	7605	6780		6.78	17.37	156.24	31	51.3	17.7	Silty clay loam
20	HH	Henan	115.2	34.7	5235			8.04	8.37	66.39	18.8	47.9	33.3	Loam
21	HH	Henan	115.1	34.6	6360	7080		8.17	12.76	81.71	18.8	47.9	33.3	Loam
22	HH	Henan	114.9	33.8	6285	6120		7.95	17.02	172.29	26.7	49.1	24.2	Loam
23	HH	Henan	112.8	32.7	6285	5445		6.35	14.99	133.31	35.7	29.8	34.5	Clay loam
24	HH	Henan	115.1	35.8	6465	8145		8.21	14.24	86.57	23.5	46.4	30.1	Loam
25	HH	Henan	114.8	34.6	6360	7395		8.45	9.77	79.17	19.6	50.5	29.9	Silty loam

Site	n ·		x % 1	T (* 1	Yield (of Main (kg ha ⁻¹)	Crops				Soil Proj	perties		
Number	Region	Province	Longitude	Latitude	Wheat	Maize	Rice	pН	SOM ¹ (g/kg)	AK ² (mg/kg)	Clay (%)	Silt (%)	Sand (%)	Soil Texture
26	YR	Jiangsu	119.7	31.4	5625		9060	5.76	25.54	92.88	29	53	18	Silty clay loam
27	YR	Jiangsu	119.7	31.9	5265		8670	6.62	21.03	104.15	28	57	15	Silty clay loam
28	YR	Guangxi	110.8	24.8			5910	6.95	56.36	109.81	29	38	33	Loam
29	YR	Jiangsu	120	32.9	5820		9345	6.93	20.08	121.03	27	56	17	Silty loam
30	YR	Jiangsu	120.5	33.2	6075		9705	8.09	17.21	137.45	27	53	20	Silty clay loam
31	YR	Anhui	116.2	33.5	6465	6552		7.10	16.30	139.41	43.1	45.5	11.4	Silty clay
32	YR	Anhui	116.4	33.1	6735	7590		6.46	19.26	124.23	25	47	28	Loam
33	YR	Anhui	117.9	32.7	5415		7650	6.28	20.49	109.17	29	49	22	Clay loam
34	YR	Anhui	118	30.9			7830	7.64	27.70	91.94	35.8	47.3	16.9	Silty clay loam
35	YR	Fujian	117.8	27.3			7785	5.18	40.04	81.58	29.6	27.4	43	Sandy clay loam
36	YR	Fujian	117.1	26.8			7260	5.26	41.71	107.36	32	37	31	Clay loam
37	YR	Hubei	110.1	32.2	4050	5625		6.64	14.88	115.17	21	41	38	Loam
38	YR	Hubei	114	30.1	6210		6900	6.91	28.34	98.54	26	52	22	Silty loam
39	YR	Hubei	112.9	30.7	2850			7.23	21.66	137.34	29.9	38.3	31.8	Clay loam
40	YR	Hubei	112.7	30.3	3510		7815	7.32	19.57	86.90	27	46	27	Loam
41	YR	Jiangsu	119	31.6	4770		8340	6.17	25.84	114.77	35.4	47.6	17	Silty clay loam
42	YR	Jiangsu	118.7	32.4	5310		8565	6.47	22.39	111.28	32.7	49.9	17.4	Silty clay loam
43	YR	Jiangsu	119.2	32.5	5490		7830	6.22	18.44	78.64	31.7	53.7	14.6	Silty clay loam
44	YR	Jiangsu	120	32.5	6300		8640	7.34	20.92	73.16	25.5	54.6	19.9	Silty loam
45	YR	Jiangsu	120.4	32.6	6405		9180	7.58	33.17	105.11	20	56.5	23.5	Silty loam
46	YR	Jiangsu	120.5	32.8	5985	4920	9825	8.00	16.00	74.32	26.1	57.2	16.7	Silty loam
47	YR	Anhui	116.8	33.6	6885	7410		6.84	19.34	219.40	35.3	49.4	15.3	Silty clay loam
48	YR	Anhui	116.3	32.6	6915	6300		6.09	14.05	152.38	35.1	45.9	19	Silty clay loam
49	YR	Anhui	115.2	32.8	7890	8310		6.08	16.95	155.47	35	54	11	Silty clay loam
50	YR	Anhui	117.1	32.3	5715		7590	6.33	13.80	99.67	27.5	56.2	16.3	Silty loam
51	YR	Anhui	117.3	31.2			6825	5.69	20.41	79.09	37.4	49.6	13	Silty clay loam
52	YR	Anhui	118.5	30.2			7725	5.86	37.80	54.93	25.7	28.9	45.4	Sandy clay loam
53	YR	Fujian	117.8	24.4			6870	5.78	39.89	172.07	32.5	33	34.5	Clay loam
54	YR	Hubei	111.5	30.1	3390	7140		7.21	20.47	184.09	25.9	41.4	32.7	Loam
55	YR	Hubei	112.5	31.2	4395		8970	6.83	24.35	114.98	26.6	37	36.4	Loam
56	YR	Hubei	110.7	32.6	3225	5475		6.81	15.83	107.39	35.8	38.6	25.6	Clay loam
57	YR	Hubei	113	30.6			8460	6.43	31.67	84.52	30	37.3	32.7	Clay loam
58	YR	Hunan	114.1	28.3			6540	5.31	33.91	73.70	27.8	35.7	36.5	Loam
59	YR	Hunan	111.8	27.1		5805		4.77	18.29	200.58	36	39.8	24.2	Clay loam
60	YR	Hunan	113.1	29.3			6375	5.72	30.97	52.35	31.4	41.3	27.3	Clay loam
61	YR	Hunan	111.2	25.5			6315	5.93	42.33	63.73	21.9	31.2	46.9	Loam
62	YR	Hunan	110	27.1			7890	5.17	40.49	74.19	29.3	46.4	24.3	Clay loam
63	YR	Guangxi	108.8	23.2			5700	5.63	44.85	72.27	39.7	46	14.3	Silty clay loam
64	SC	Guangdong	110.2	20.5			5910	6.77	75.78	71.71	40	27	33	Clay loam

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Table A1. Cont.
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Site					Yield of Main Crops (kg ha ⁻¹) le				:	Soil Prop	perties			
Number	Region	Province	Longitude	Latitude	Wheat	Maize	Rice	pН	SOM ¹ (g/kg)	AK ² (mg/kg)	Clay (%)	Silt (%)	Sand (%)	Soil Texture
65	SC	Guangdong	112.2	22.6			5655	5.45	42.36	65.20	24	36	40	Loam
66	SC	Guangdong	113.5	24.1			6060	4.98	46.66	67.73	28.8	25.7	45.3	Sandy clay loam
67	SC	Guangdong	112.6	23.1			7500	5.49	33.56	107.91	30.3	48.1	21.6	Clay loam
68	SC	Hainan	110.4	19.2			6570	5.38	28.31	50.41	35.3	30.8	33.9	Clay loam
69	SC	Hainan	110.3	19.4			5835	5.52	28.78	60.39	35.3	30.8	33.9	Clay loam
70	SW	Sichuan	104.6	31	2985	3810		6.58	10.30	61.52	30	47	23	Clay loam
71	SW	Sichuan	104.7	31.1	3315	4065		8.06	15.94	113.35	30	47	23	Clay loam
72	SW	Guizhou	105.8	26.5		5535	5400	7.23	32.03	86.80	38	41	21	Clay loam
73	SW	Yunnan	103.9	25.4	2265	7365		5.65	42.67	152.46	37.2	38.4	24.4	Clay loam
74	SW	Yunnan	100.1	23.8		8235		5.56	78.08	186.29	30	39.5	30.5	Clay loam
75	SW	Chongqing	106.2	29.1			7215	5.14	25.66	146.46	25.6	35.3	39.1	Loam
76	SW	Chongqing	106.4	28.7			7065	6.88	30.79	147.28	25.6	35.3	39.1	Loam
77	NW	Xinjiang	79.9	37	5715	5880		8.21	12.46	159.46	19	43	38	Loam
78	NW	Xinjiang	79.7	37.2	5550	5895		8.27	13.21	162.68	22.8	42.3	34.9	Loam
79	NW	Gansu	101.9	38.2	7470			8.35	19.68	314.06	16	42	42	Sandy loam
80	NW	Gansu	102.9	37.5	6000	6645		8.61	8.13	120.07	15	47	38	Loam
81	NW	Ningxia	106.5	39.4	5040	9945		8.33	17.66	213.49	20	55	25	Silty loam
82	NW	Ningxia	105.8	37.8	5595	8355	9675	8.30	12.97	116.78	19	47	24	Loam
83	NW	Xinjiang	88.6	45.1	4890	8310		7.50	21.21	240.29	28	45	27	Loam
84	NW	Ningxia	106.1	38.6	13,095		8850	8.30	19.38	142.57	26.8	43	30.2	Loam
85	NW	Ningxia	106.4	38.7			8955	8.23	22.09	187.63	26.8	43	30.2	Loam
86	NW	Shanxi	110.98	35.02	4710	6090		8.61	13.38	123.61	20	44	36	Loam
87	NW	Gansu	105.5	34.3	4680	8490		8.25	12.35	113.69	18	47	35	Loam
88	NW	Qinghai	102.4	36.4	2745			8.34	11.88	198.76	18	46	36	Loam
89	NW	Qinghai	102.2	36.8	4095			8.15	20.50	133.45	20	45	35	Loam
90	NW	Shaanxi	109.6	33.4	4020	5745		7.85	16.50	135.87	19.6	47.6	32.8	Loam
91	NW	Shanxi	111.23	37.88		7080		8.25	9.13	139.50	14	43	43	Loam

Table A1. Cont.

¹ SOM: soil organic matter, g·kg⁻¹. ² AK: available potassium, mg·kg⁻¹. ³ NE: northeast China, HH: Huang-Huai-Hai region, YR: middle and lower reaches of the Yangtze River, SC: South China, SW: southwest China, NW: northwest China.

						Cli	mate			Agronomic Practices	i	Topographic Information
Site Number	Region	Province	Longitude	Latitude	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	Drought Index	Climatic Zone	Fertilizer P Rate (kg P ₂ O ₅ ha ⁻¹)	Cropping Pattern	Land Use Type	Slope (°C)
1	NE ¹	Jilin	126.5	44.8	4.0	600	42.86	middle temperate zone	108.19	mono-cropping	upland	2–6
2	NE	Jilin	129.1	42.8	5.5	550	35.48	middle temperate zone	75.22	mono-cropping	paddy	0–2
3	NE	Heilongjiang	126.8	46	3.3	505	37.97	middle temperate zone	95.56	mono-cropping	upland	2–6
4	NE	Heilongjiang	126.3	45.3	4.4	481	33.40	middle temperate zone	85.86	mono-cropping	upland	2–6
5	NE	Heilongjiang	127	47	2.4	547	44.11	middle temperate zone	90.22	mono-cropping	paddy	6–15
6	NE	Heilongjiang	126.8	46.7	2.4	547	44.11	middle temperate zone	65.80	mono-cropping	paddy	2–6
7	NE	Heilongjiang	129.7	46.2	1.2	569.1	50.81	middle temperate zone	74.20	mono-cropping	paddy-upland	2–6
8	HH	Hebei	114.4	37.2	13.9	480	20.08	warm temperate zone	168.44	double-cropping	upland	2–6
9	HH	Hebei	114.3	38.1	13.3	536	23.00	warm temperate zone	219.54	double-cropping	upland	2–6
10	HH	Shandong	117.1	36.5	14.3	665.7	27.40	warm temperate zone	208.34	double-cropping	upland	2–6
11	HH	Shandong	115.5	35.3	14.8	640	25.81	warm temperate zone	167.42	double-cropping	upland	2–6
12	HH	Shandong	120.4	36.8	11.3	732	34.37	warm temperate zone	230.53	double-cropping	upland	2–6
13	HH	Shandong	119.8	37.2	12.5	809	35.96	warm temperate zone	155.96	double-cropping	upland	0–2
14	HH	Shandong	120	37.1	12.5	809	35.96	warm temperate zone	144.94	double-cropping	upland	2–6

Table A2. Detailed information of field experimental sites.

					Climate					Agronomic Practices	;	Topographic Information
Site Number	Region	Province	Longitude	Latitude	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	Drought Index	Climatic Zone	Fertilizer P Rate (kg P ₂ O ₅ ha ⁻¹)	Cropping Pattern	Land Use Type	Slope (°C)
15	HH	Henan	112.3	32.6	15.1	721	28.73	warm temperate zone	149.09	double-cropping	upland	2–6
16	HH	Henan	114.1	31.2	15.2	1100	43.65	warm temperate zone	91.61	double-cropping	upland	2–6
17	HH	Henan	114.5	32.4	15.0	938	37.52	warm temperate zone	125.95	double-cropping	upland	2–6
18	HH	Henan	112.8	34.6	14.2	570	23.55	warm temperate zone	174.03	double-cropping	upland	2–6
19	HH	Henan	114.6	33.5	14.5	785.1	32.04	warm temperate zone	145.44	double-cropping	upland	0–2
20	HH	Henan	115.2	34.7	14.1	674	27.97	warm temperate zone	209.78	double-cropping	upland	2–6
21	HH	Henan	115.1	34.6	14.1	674	27.97	warm temperate zone	193.72	double-cropping	upland	2–6
22	HH	Henan	114.9	33.8	14.0	740	30.83	warm temperate zone	139.30	double-cropping	upland	2–6
23	HH	Henan	112.8	32.7	15.2	910.1	36.12	warm temperate zone	174.43	double-cropping	upland	2–6
24	HH	Henan	115.1	35.8	13.4	540	23.08	warm temperate zone	158.03	double-cropping	upland	2–6
25	HH	Henan	114.8	34.6	14.0	678.2	28.26	warm temperate zone	158.57	double-cropping	upland	2–6
26	YR	Jiangsu	119.7	31.4	16.1	1300	49.81	subtropical zone	120.57	double-cropping	paddy-upland	2–6
27	YR	Jiangsu	119.7	31.9	16.5	1043	39.36	subtropical zone	143.67	double-cropping	paddy-upland	2–6
28	YR	Guangxi	110.8	24.8	18.8	1894	65.76	subtropical zone	157.13	double-cropping	paddy	2–6
29	YR	Jiangsu	120	32.9	15.0	1032.3	41.29	subtropical zone	125.42	double-cropping	paddy-upland	2–6
30	YR	Jiangsu	120.5	33.2	14.1	1042.2	43.24	subtropical zone	174.52	double-cropping	paddy-upland	2–6
31	YR	Anhui	116.2	33.5	15.1	809.8	32.26	subtropical zone	115.72	double-cropping	upland	2–6

Table A2. Cont.

					Climate					Agronomic Practices	5	Topographic Information
Site Number	Region	Province	Longitude	Latitude	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	Drought Index	Climatic Zone	Fertilizer P Rate (kg P ₂ O ₅ ha ⁻¹)	Cropping Pattern	Land Use Type	Slope (°C)
32	YR	Anhui	116.4	33.1	14.9	812	32.61	subtropical zone	161.83	double-cropping	upland	2–6
33	YR	Anhui	117.9	32.7	14.9	1100	44.18	subtropical zone	158.37	double-cropping	paddy-upland	2–6
34	YR	Anhui	118	30.9	15.0	1000	40.00	subtropical zone	130.91	double-cropping	paddy-upland	2–6
35	YR	Fujian	117.8	27.3	18.1	1742	61.99	subtropical zone	86.68	double-cropping	paddy	2–6
36	YR	Fujian	117.1	26.8	19.2	1753	60.03	subtropical zone	109.95	double-cropping	paddy	0–2
37	YR	Hubei	110.1	32.2	14.0	905	37.71	subtropical zone	148.51	triple-cropping	upland	0–2
38	YR	Hubei	114	30.1	17.0	900	33.33	subtropical zone	155.78	triple-cropping	paddy-upland	2–6
39	YR	Hubei	112.9	30.7	16.2	900	34.35	subtropical zone	149.94	double-cropping	upland	2–6
40	YR	Hubei	112.7	30.3	16.1	1200	45.98	subtropical zone	175.74	double-cropping	upland	2–6
41	YR	Jiangsu	119	31.6	15.6	1037.6	40.53	subtropical zone	130.19	double-cropping	paddy	2–6
42	YR	Jiangsu	118.7	32.4	15.6	1100	42.97	subtropical zone	125.40	double-cropping	paddy	2–6
43	YR	Jiangsu	119.2	32.5	15.8	1015	39.34	subtropical zone	100.66	double-cropping	paddy	2–6
44	YR	Jiangsu	120	32.5	14.5	991.7	40.48	subtropical zone	133.74	double-cropping	paddy	2–6
45	YR	Jiangsu	120.4	32.6	15.8	959	37.17	subtropical zone	103.33	double-cropping	paddy	2–6
46	YR	Jiangsu	120.5	32.8	15.0	1061.2	42.45	subtropical zone	131.93	double-cropping	paddy	0–2
47	YR	Anhui	116.8	33.6	15.3	850	33.60	subtropical zone	190.30	double-cropping	upland	2–6
48	YR	Anhui	116.3	32.6	15.4	930	36.61	subtropical zone	127.10	double-cropping	upland	2–6
49	YR	Anhui	115.2	32.8	15.4	890	35.04	subtropical zone	142.65	double-cropping	upland	0–2
50	YR	Anhui	117.1	32.3	15.0	960	38.40	subtropical zone	152.36	double-cropping	paddy	2–6
51	YR	Anhui	117.3	31.2	16.2	1262.9	48.20	subtropical zone	133.75	double-cropping	paddy	2–6
52	YR	Anhui	118.5	30.2	15.6	1430	55.86	subtropical zone	52.48	double-cropping	paddy	0–2
53	YR	Fujian	117.8	24.4	21.5	1563.2	49.63	subtropical zone	119.92	double-cropping	paddy	2–6
54	YR	Hubei	111.5	30.1	16.2	1250	47.71	subtropical zone	136.70	double-cropping	upland	2–6
55	YR	Hubei	112.5	31.2	16.4	987.6	37.41	subtropical zone	153.45	double-cropping	paddy	2–6

					Climate					Agronomic Practices	5	Topographic Information
Site Number	Region	Province	Longitude	Latitude	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	Drought Index	Climatic Zone	Fertilizer P Rate (kg P ₂ O ₅ ha ⁻¹)	Cropping Pattern	Land Use Type	Slope (°C)
56	YR	Hubei	110.7	32.6	15.4	780	30.71	subtropical zone	138.00	double-cropping	upland	0–2
57	YR	Hubei	113	30.6	16.4	960	36.36	subtropical zone	90.20	double-cropping	paddy	2–6
58	YR	Hunan	114.1	28.3	18.2	1395.7	49.49	subtropical zone	85.43	double-cropping	paddy	2–6
59	YR	Hunan	111.8	27.1	16.6	1250	46.99	subtropical zone	177.40	double-cropping	upland	2–6
60	YR	Hunan	113.1	29.3	16.8	1295	48.32	subtropical zone	114.52	double-cropping	paddy	2–6
61	YR	Hunan	111.2	25.5	18.0	1600	57.14	subtropical zone	119.08	double-cropping	paddy	0–2
62	YR	Hunan	110	27.1	17.0	1700	62.96	subtropical zone	144.80	double-cropping	paddy	2–6
63	YR	Guangxi	108.8	23.2	28.8	1589.2	40.96	subtropical zone	102.25	double-cropping	paddy	2–6
64	SC	Guangdong	110.2	20.5	23.3	1364	40.96	subtropical zone	210.46	double-cropping	paddy	6–15
65	SC	Guangdong	112.2	22.6	21.5	1663.7	52.82	subtropical zone	81.09	double-cropping	paddy	2–6
66	SC	Guangdong	113.5	24.1	19.3	1923	65.63	subtropical zone	101.15	double-cropping	paddy	2–6
67	SC	Guangdong	112.6	23.1	21.2	1650	52.88	subtropical zone	87.90	double-cropping	paddy	0–2
68	SC	Hainan	110.4	19.2	24.1	1900	55.72	tropical zone	163.08	double-cropping	paddy	2–6
69	SC	Hainan	110.3	19.4	24.0	1953	57.44	tropical zone	126.40	double-cropping	paddy	6–15
70	SW	Sichuan	104.6	31	16.7	950	35.58	subtropical zone	169.46	triple-cropping	upland	0–2
71	SW	Sichuan	104.7	31.1	16.7	950	35.58	subtropical zone	170.52	triple-cropping	upland	0–2
72	SW	Guizhou	105.8	26.5	15.1	1378.2	54.91	subtropical zone	72.60	double-cropping	paddy	0–2
73	SW	Yunnan	103.9	25.4	14.5	1008	41.14	subtropical zone	97.12	double-cropping	upland	0–2
74	SW	Yunnan	100.1	23.8	17.9	1158	41.51	subtropical zone	147.37	double-cropping	upland	0–2
75	SW	Chongqing	106.2	29.1	18.0	1000	35.71	subtropical zone	57.30	double-cropping	paddy	0–2
76	SW	Chongqing	106.4	28.7	18.2	1034.7	36.69	subtropical zone	53.80	double-cropping	paddy	0–2
77	NW	Xinjiang	79.9	37	8.9	35	1.85	warm temperate zone	368.49	double-cropping	upland	0–2
78	NW	Xinjiang	79.7	37.2	8.9	35	1.85	warm temperate zone	307.58	double-cropping	upland	0–2

					Climate					Agronomic Practices	5	Topographic Information
Site Number	Region	Province	Longitude	Latitude	Mean Annual Temperature (°C)	Mean Annual Precipitation (mm)	Drought Index	Climatic Zone	Fertilizer P Rate (kg P ₂ O ₅ ha ⁻¹)	Cropping Pattern	Land Use Type	Slope (°C)
79	NW	Gansu	101.9	38.2	4.8	185.1	12.51	warm temperate zone	158.38	mono-cropping	upland	0–2
80	NW	Gansu	102.9	37.5	5.6	300	19.23	warm temperate zone	132.91	mono-cropping	upland	0–2
81	NW	Ningxia	106.5	39.4	8.9	173.2	9.16	middle temperate zone	219.27	mono-cropping	upland	0–2
82	NW	Ningxia	105.8	37.8	8.5	260.7	14.09	middle temperate zone	151.54	double-cropping	paddy-upland	0–2
83	NW	Xinjiang	88.6	45.1	5.9	145	9.12	middle temperate zone	140.21	mono-cropping	upland	0–2
84	NW	Ningxia	106.1	38.6	8.3	210	11.48	middle temperate zone	102.64	mono-cropping	paddy	0–2
85	NW	Ningxia	106.4	38.7	8.5	225	12.16	middle temperate zone	95.30	mono-cropping	paddy	0–2
86	NW	Shanxi	110.98	35.02	13.6	559.3	23.70	warm temperate zone	150.26	mono-cropping	upland	0–2
87	NW	Gansu	105.5	34.3	10.7	531	25.63	warm temperate zone	132.49	mono-cropping	upland	0–2
88	NW	Qinghai	102.4	36.4	7.3	335.4	19.39	warm temperate zone	232.38	mono-cropping	upland	0–2
89	NW	Qinghai	102.2	36.8	5.8	477.4	30.22	warm temperate zone	177.34	mono-cropping	upland	0–2
90	NW	Shaanxi	109.6	33.4	13.1	709	30.69	warm temperate zone	241.08	double-cropping	upland	0–2
91	NW	Shanxi	111.23	37.88	7.7	520	29.38	middle temperate zone	119.80	mono-cropping	upland	0–2

Table A2. Cont.

¹ NE: northeast China, HH: Huang-Huai-Hai region, YR: middle and lower reaches of the Yangtze River, SC: South China, SW: southwest China, NW: northwest China.

Region	Current Olsen-P Content (mg kg $^{-1}$)	Current P Surplus (kg ha ⁻¹)	SPTE (mg kg ⁻¹)	Time Required to Reach Environmental Threshold (Year)
NE ¹	26.63	20.44	1.22	53.6-101.7
HH	29.55	86.9	0.91	13.2–28.4
YR	25.12	55.72	1.04	25.7-46.4
SC	35.27	52.6	1.97	4.6–16.1
SW	36.6	40.38	1.12	7.5–34.1
NW	29.21	103.53	0.87	12.0–25.3

Table A3. A forecast of the time different Chinese regions will reach the environmental P threshold based on SPTE.

¹ NE: northeast China, HH: Huang-Huai-Hai region, YR: middle and lower reaches of the Yangtze River, SC: South China, and SW: southwest China, NW: northwest China.



Step1:

We assumed five factors (DI, P fertilizer dosage, soil silt content, soil pH, SOM) had a direct impact on SPTE.

Step2:

We assumed DI and P fertilizer dosage had a direct effect on soil properties.

Step3:

We assumed soil silt content indirectly affected SPTE by affecting soil pH and SOM.

Step4:

We assumed all significant correlation paths were retained in the framework.

Figure A1. Schematic diagram for structural equation modeling (SEM). DI: drought index. SOM: soil organic matter, $g \cdot kg^{-1}$. AK: available potassium, $mg \cdot kg^{-1}$. SPTE: soil P transformation efficiency, $mg \cdot kg^{-1}$.



Figure A2. Soil Olsen-P content (**a**) and apparent P surplus (**b**) over time under 91 long-term fertilization experiments.



Figure A3. Cumulative P surplus over time in 91 experiment sites. The experimental time (year) 5, 10, 15, 20, 25, and 30 refer to the time periods of 1988–1992, 1993–1997, 1998–2002, 2003–2007, 2008–2012, and 2013–2018, respectively.



Figure A4. P fertilizer amount over time in 91 experiment sites.



Figure A5. Cont.



Figure A5. Yield of main crops in 91 experiment sites.

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