

Article

Nocturnal Ozone Enhancement Induced by Sea-Land Breezes During Summertime in Northern Coastal City Qingdao, China

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Abstract: This study investigated the influence of sea–land breezes on nocturnal spatial and temporal distribution of ozone (O₃) and its potential effects on particulate nitrate formation in Qingdao, a coastal city in northern China. Observation campaigns were conducted to measure surface air pollutants and meteorological factors during a typical sea–land breezes event from 22 to 23 July 2022. A coherent Doppler lidar (CDL) system was employed to continuously detect three-dimensional wind fields. The results revealed that nocturnal ozone levels were enhanced by a conversion of sea–land breezes. Initially, the prevailing northerly land breeze transported high concentrations of O₃ and other air pollutants from downtown to the Yellow Sea. As the sea breeze developed in the afternoon, the sea breeze front advanced northward, resulting in a flow of high O₃ concentrations back into inland areas. This penetration of the sea breeze front led to a notable spike in O₃ concentrations between 16:00 on 22 July and 02:00 on 23 July across downtown areas, with an average increase of over 70 µg/m³ within 10 min. Notably, a time lag in peak O₃ concentration was observed with southern downtown areas peaking before northern rural areas. During this period, combined pollution of O₃ and PM_{2.5} was also observed. These findings indicated that the nighttime increase in O₃ concentrations, coupled with enhanced atmospheric oxidation, would likely promote the secondary conversion of gaseous precursors into PM_{2.5}.

Keywords: nocturnal ozone enhancement; spatial–temporal distribution; sea–land breezes; wind profile; transport



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1. Introduction

Surface ozone (O₃) is a secondary pollutant generated by the photochemical reaction of nitrogen oxides (NO_x) and volatile organic compounds (VOCs), playing a crucial role in atmospheric chemistry [1,2]. O₃ pollution not only impacts air quality and climate change but also poses significant risks to human health, crop yields, and ecosystems [3–5]. Although the Clean Air Action Plan (CAAP, 2013–2017) and the Blue Sky Protection Campaign (BSPC, 2018–2020) successfully reduced PM_{2.5} concentration in China, O₃ concentrations showed a fluctuating but upward trend [6,7], particularly in the central and eastern regions [8].

The atmospheric chemical and physical processes of surface O₃ exhibited distinct differences between day and night [9], typically showing a unimodal diurnal variation that peaks in the afternoon due to photochemical formation from NO_x (NO and NO₂) and volatile organic compounds (VOCs) in the presence of sunlight [10] and reaches a minimum before sunrise because of chemical destruction (such as NO titration, NO + O₃ → NO₂ + O₂) or dry deposition [11]. However, in recent years, significant nocturnal O₃ enhancement

(NOE) events, referring to the phenomenon of nighttime O_3 concentrations rising by 5 to 30 ppbv [12,13], have been observed at various sites around the world, such as urban and rural sites in northern Portugal [14], multiple urban sites in UK [15], and southeast United States [16]. These increases were attributed to the horizontal transport of O_3 from upwind-polluted regions or dynamic atmospheric processes, such as turbulent mixing and vertical mixing triggered by convective storms or low-level jets [17–19]. Elevated nocturnal O_3 levels could enhance the nighttime atmospheric oxidation capacity, promoting the formation of secondary pollutants such as particulate nitrate and secondary organic aerosols [13,20].

With half of the world's population residing within 100 km of coastlines, air quality in coastal areas present a significant health risk to residents [21]. Sea–land breezes are one of the most common mesoscale circulation pattern in these regions, resulting from uneven heating induced by differences in heat capacity between land and sea [22]. Studies have highlighted that air quality is typically worse on sea–land breezes days than on non-sea–land breezes days [23–26]. Moreover, O_3 pollution events during sea–land breezes days remain particularly complex due to spatial and temporal variations in air temperature (T), relative humidity (RH), and wind fields [27–29]. These variations not only impact photochemical reaction conditions but also atmospheric diffusion conditions, leading to inhomogeneous distributions of O_3 or secondary aerosol [27,30,31]. Therefore, understanding the influence of wind flow and pollutant dispersion on O_3 distribution is essential for effectively addressing air quality issues in coastal regions.

While many studies conducted in northern China have focused on O_3 characteristics, meteorological causes, or formation sensitivities [32–35], less attention has been paid to nocturnal O_3 concentrations in coastal regions [36]. Qingdao, an important coastal city located in east of Shandong province, presents an ideal location for examining O_3 pollution processes influenced by sea–land breezes [37]. The objectives of this study were to (a) analyze the impact of sea–land breezes on the spatial and temporal distribution of O_3 during this pollution episode and (b) explore the potential influence of O_3 on nighttime particulate nitrate formation. Thus, ground-based Lidar measurements were employed to investigate three-dimensional (3D) wind fields in Qingdao from 22 to 23 July 2022. Variations in ground-level air pollutant concentrations were comprehensively analyzed in conjunction with meteorological data and synoptic weather patterns. Our results provided insights into nocturnal ozone enhancement (NOE) influenced by sea–land breezes in the northern coastal region of China.

2. Materials and Methods

2.1. Observation Materials

Hourly surface concentration of O_3 , $PM_{2.5}$, NO_2 , NO , and CO from 22 to 23 July 2022 were observed at a supersite in Laoshan (LS), Qingdao (36.11° N, 120.47° E) [38], as depicted in Figure 1. Concurrently, hourly meteorological data, including wind direction, wind speed, T, RH, and solar radiation (SR), were observed. Additionally, surface O_3 concentrations were obtained from an ambient-air-quality-monitoring station network in Qingdao to analyze the spatial and temporal distribution of O_3 . All data were reported in local time (LT, LT = UTC + 8). Synoptic weather charts depicting 500 hPa and 850 hPa geopotential height, as well as sea level pressure, were sourced from the Korea Meteorological Administration (<https://www.kma.go.kr/nchn/image/chart/analysis-chart.do>, accessed on 24 July 2022).

2.2. Wind Fields

Ground-level winds were observed at LS, Xihai (XHA), Shibe (SB), Jiaozhou (JZ), Jimo (JM), Laixi (LX), and Pingdu (PD) to analyze the evolution of sea–land breezes. Three-dimensional wind profiles were collected at SB, yard of Qingdao Eco-Environmental Monitoring Center of Shandong Province (36.07° N, 120.34° E), by using a coherent Doppler lidar (CDL) system manufactured by Qingdao Leice Transient Technology Co., Ltd, Qing-

dao, China. [39]. The monitoring site was located approximately 1.2 km from the coastline. The CDL system was deployed on the roof of a container, approximately 3 m above ground level, and operated in Doppler-Beam-Scanning mode with an angle of 60° to the vertical direction and a laser wavelength of $1.5 \mu\text{m}$. It provided continuous 3D wind profiles with a spatial resolution of 30 m across altitudes ranging from 77 to 2000 m. Wind speed measurement accuracy was less than 0.1 m/s, and wind direction measurement accuracy was within 3° . Horizontal wind direction and speed, as well as vertical wind speed with a time resolution of 10 min, were used for analysis.

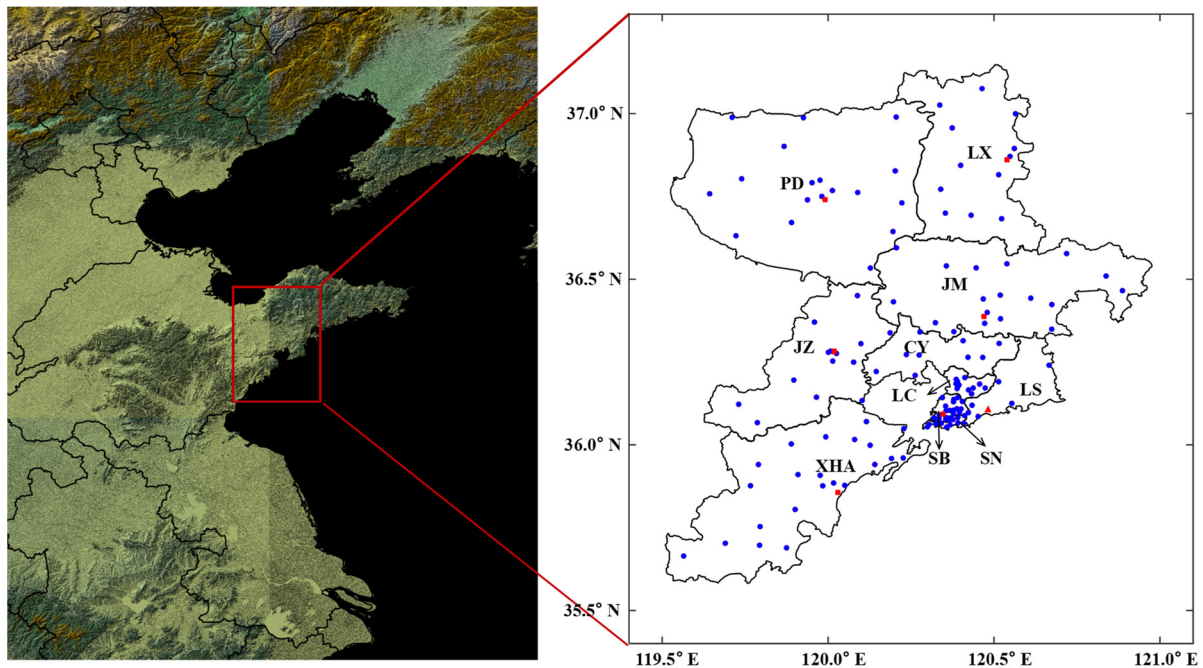


Figure 1. Locations of air-quality-monitoring stations (blue circles), wind field sites (red squares), and supersite (red triangle).

2.3. Criteria to Identify Sea–Land Breezes

As established in previous studies [37], days meeting specific criteria, including wind direction, wind speed, occurrence time, and duration, were classified as sea–land breeze days. Considering the direction of the coastline in Qingdao, a sea breeze was defined by wind directions between 50° and 210° , while a land breeze was defined by wind directions between 0° and 30° or 230° to 360° .

The following criteria were applied [40–42]:

1. Following the sea and land breeze diurnal variation pattern, a period from 01:00 to 08:00 was considered as the land breeze period; 09:00–12:00 was marked as the transition from the land to sea breeze; 13:00–20:00 was identified as the sea breeze period; and 21:00–00:00 was noted as the transition from the sea breeze to the land breeze.
2. Sea breeze occurrences were required to last at least 4 h during the sea breeze period, while land breeze occurrences were limited to a maximum of 2 h.
3. Land breeze occurrences were required to last at least 4 h, whereas sea breeze occurrences should not exceed 2 h.
4. Wind speeds exceeding 10 m/s were excluded from sea–land breezes analysis to eliminate interference from large-scale weather systems.

Additionally, the absolute difference in wind direction between 02:00 and 14:00 was set between 90° and 270° .

2.4. HYSPLIT Model

Air mass back-trajectories were calculated using the National Oceanic and Atmospheric Administration (NOAA) Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model, utilizing CDC1 Meteorological data. To trace the local or regional transport sources of O_3 at midnight (00:00 on 23 July), 24 h backward trajectories were designated at SB site (120.34° E, 36.07° N) at altitudes of 100, 500, and 1000 m above the ground level.

3. Results and Discussion

3.1. Variation in Air Pollutant Concentrations and Meteorological Data

Hourly variations in wind direction, wind speed, T, RH, precipitation, SR, and concentrations of O_3 , NO_2 , NO, CO, and $PM_{2.5}$ observed at LS from 22 to 23 July 2022, were depicted in Figure 2. O_3 pollution was evident in Qingdao over the two consecutive days. The daily maximum 8 h average O_3 concentration (MDA8) on 22 July and 23 July were $215 \mu\text{g}/\text{m}^3$ and $181 \mu\text{g}/\text{m}^3$, respectively, exceeding the national secondary standard limit of $160 \mu\text{g}/\text{m}^3$ for 1.34 and 1.13 times.

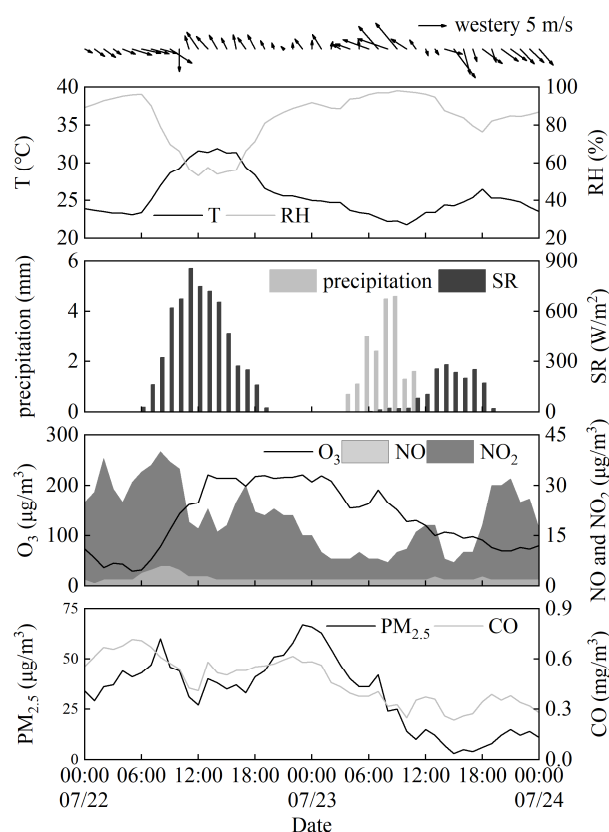


Figure 2. Hourly variations of wind direction, wind speed, air temperature (T), relative humidity (RH), precipitation, solar radiation (SR), as well as concentrations of O_3 , NO, NO_2 , $PM_{2.5}$, and CO.

Typically, O_3 concentrations exhibited a unimodal distribution, peaking between 15:00 and 16:00 [43]. However, on both days, O_3 concentration remained unusually high for extended periods during nighttime. A rapid increase in O_3 was observed from 06:00 to 13:00 on 22 July, reaching a peak of $221 \mu\text{g}/\text{m}^3$ at 13:00, with an increase rate of $26 \mu\text{g}/(\text{m}^3 \cdot \text{h})$. Although SR decreases rapidly after sunset, resulting in weakened photochemical reactions, the O_3 concentration remained high and comparable to midday peak levels. High levels of O_3 , averaging $(215 \pm 6) \mu\text{g}/\text{m}^3$, persisted until early morning on 23 July at 02:00. Afterward, the O_3 concentration gradually declined, but with a slight increase observed from 05:00 to 08:00 on 23 July.

Between 01:00 and 02:00 on 23 July, the change rate in the hourly O₃ concentration ($\Delta\text{O}_3/\Delta t$) was 12 $\mu\text{g}/(\text{m}^3\cdot\text{h})$. According to previous studies [12,13], an hourly nocturnal (between 20:00 and 06:00) increase in O₃ concentration exceeding 10 $\mu\text{g}/\text{m}^3$ from the previous hour can be identified as an NOE event. O₃ formation in the nocturnal boundary layer is impossible without SR, as nocturnal O₃ depletion is primarily driven by dry deposition and chemical reactions with NO_x [2,6,10,11]. Therefore, the increase in O₃ concentration in the nighttime required further investigation.

On 22 July, higher concentrations of PM_{2.5}, NO_x, and CO were observed alongside a northerly land breeze. High T, peaking of 31.9 °C at 14:00, a total daily SR of 5968 W/m², and low average RH of 66%, were conducive to local formation of O₃ produced from photochemistry of NO_x and VOCs, leading to a rapid increase in O₃ concentration. In contrast, on 23 July, the highest temperature dropped to 26.5 °C at 18:00, with lower daily cumulative SR of only 1684 W/m² and higher average RH of 90%. Additionally, 9.3 mm of precipitation fell from 04:00 to 11:00, facilitating the wet removal of air pollutants. These atmospheric conditions contributed to a reduction in air pollutant levels and inhibited local O₃ formation.

MDA8s were observed from 16:00 to 23:00 on 22 July and from 01:00 to 08:00 on 23 July, rather than during the period of intense photochemical activities from noon to dusk, suggesting that the high O₃ levels during the NOE event may be influenced by the transport of O₃ pollution [44]. Notable, at 13:00 on 22 July, the wind direction near the ground shifted from land breeze to sea breeze, persisting for approximately 25 h until noon on 23 July. This wind shift indicated that the NOE pollution process may be related to the conversion from land breeze to sea breeze.

It is also worth noting that the daily variation in the PM_{2.5} concentration exhibited a bimodal distribution on 22 July. The first peak of 60 $\mu\text{g}/\text{m}^3$ occurred around 08:00. The PM_{2.5} concentration remained relatively low from 11:00 to 17:00 but rose again after 18:00, reaching a second peak of 67 $\mu\text{g}/\text{m}^3$ at 23:00.

3.2. Synoptic Weather Patterns and Backward Trajectories

O₃ pollution is generally associated with dominant weather patterns [23]. On 22 July, Qingdao and its surrounding regions were predominantly under the influence of a high-pressure ridge (584–588 dagpm), resulting in clear weather conditions that intensified SR (Figure 3). These atmospheric conditions were conducive for photochemical reactions of O₃ generation from O₃ precursors (NO_x and VOCs) [32], leading to a rapid increase in hourly concentration of O₃ up to 221 $\mu\text{g}/\text{m}^3$ at 13:00. The circulation pattern exhibited a shift in wind currents from northwest to southwest, with relatively sparse isobars at 850 hPa geopotential height, facilitating the development of local circulations [36], such as sea–land breezes. These breezes can transport O₃ from inland areas to coastal areas and then back to inland areas [24]. Additionally, a ground temperature inversion occurred at night, extending a height of 311 m and exhibiting a strength of 0.7 °C/100 m. This inversion prevents vertical mixing, contributing to the accumulation of O₃ formed during the day and trapping the air pollutants in the stable nighttime boundary layer. Consequently, nocturnal high levels of O₃ concentration could be attributed to both the intense formation of O₃ in daytime and/or potential O₃ transport. In contrast, on 23 July, Qingdao and its adjacent regions were positioned ahead of a trough, as discerned by 500 hPa geopotential height field. A low-pressure vortex moved eastward at the 850 hPa geopotential height, coupled with the passage of a cold front at the surface. The cold front brought precipitation, which inhibited local O₃ formation.

To better understand the transport of ground-level O₃ at midnight, 24 h backward trajectories of air masses were calculated at 00:00 on 23 July. The trajectories at 100, 500, and 1000 m above Qingdao were depicted in Figure 4 (red: 100 m; blue: 500 m; green: 1000 m). Analysis of these trajectories revealed no crossover among different heights, indicating limited vertical mixing. The air masses predominantly originated from local sources rather than long-range transport, as the trajectories remained within a distance of less than 150 km.

The air masses initially came from local inland areas, then moved towards the Yellow sea before returning inland. It can also be inferred that the NOE event was influenced by the transition between land breeze and sea breeze.

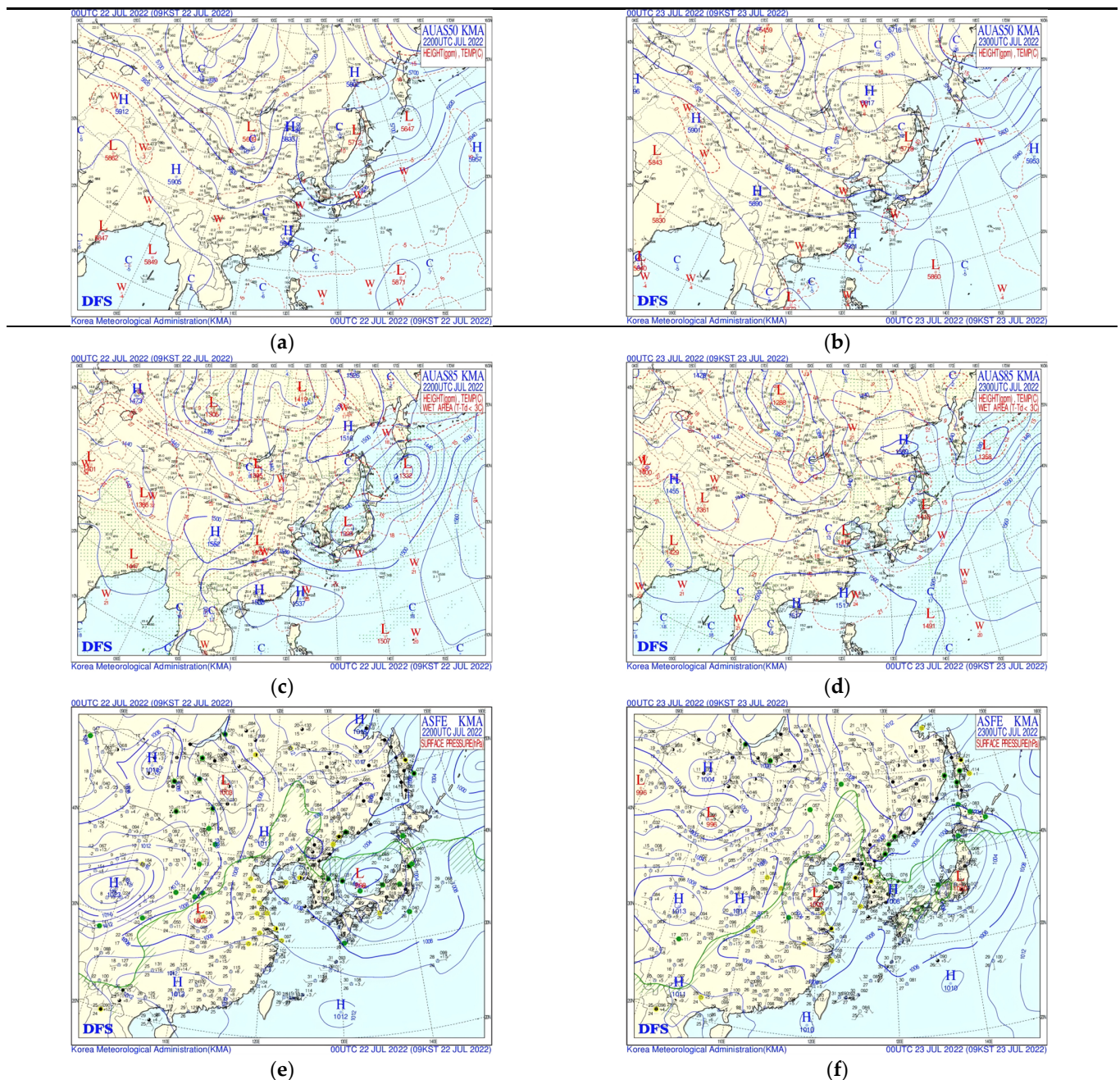


Figure 3. Synoptic weather charts of 500 hPa geopotential height (a,b), 850 hPa geopotential height (c,d), and sea-level pressure (e,f) at 08:00 on 22–23 July 2022. (<https://www.kma.go.kr/nchn/image/chart/analysis-chart.do>, accessed on 24 July 2022).

3.3. Effect of Sea/Land Breezes on Spatial and Temporal Distribution of Ozone

Analysis of surface-wind flows from 22 to 23 July revealed that offshore winds were predominantly northerly (Figure 5). Around noon on 22 July, winds in southern downtown shifted to an onshore southeasterly flow, creating a ‘head-to-head wind’ zone along the coastline due to the convergence of land and sea breezes. By the afternoon, the sea breeze strengthened, altering the wind direction to an onshore southerly or southeasterly flow from southern coastal areas to northern inland regions. At 11:00, wind at LS had fully

transitioned to a sea breeze, followed by XHA at 13:00, SB at 16:00, JZ at 18:00, and JM at 20:00, while PD and LX in the northern rural areas remained under land breeze dominance. The sea breeze persisted until early morning on 23 July as surface cyclone weakened, after which it quickly shifted to a northerly wind, eliminating the sea–land breezes convergence zones.

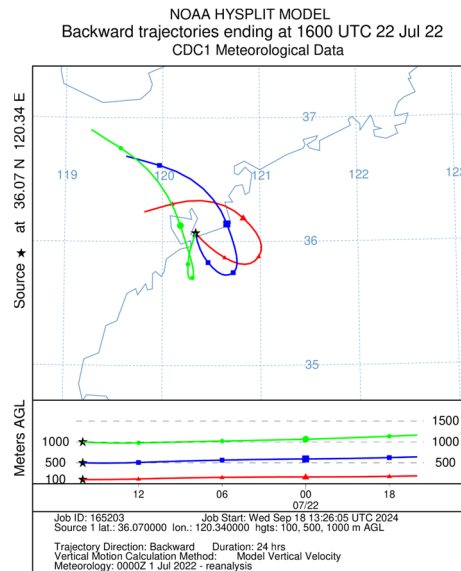


Figure 4. Backward trajectories ending at 00:00 23 July 2022.

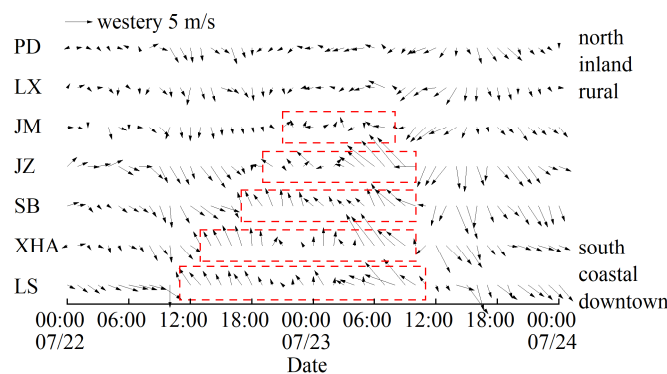


Figure 5. Surface winds observed in LS, XHA, SB, JZ, JM, LX and PD on 22–23 July 2022. (The red rectangles indicate the duration of the sea breeze).

Based on 3D wind profiles at SB (Figure 6), before the occurrence of sea breeze on 22 July, a northerly land breeze prevailed up to 2 km near the ground. At 16:00, the wind gradually shifted to sea breeze at lower altitudes, with height expanding from 300 m to 700 m, ultimately reaching approximately 1.2 km at midnight. Influenced by the surface cyclone, the sea breeze continued until noon on 23 July, before shifting to northerly winds after 13:00. During this period, downdrafts dominated, with weak updrafts occurring near the surface up to 500 m.

It can be inferred that the sea breeze lasted for up to 25 h (LS site) and penetrated inland up to 50 km from the coastline, reaching a maximum height of 1.2 km (SB site). Under conditions of strong SR, high T, and weak northerly wind on 22 July, precursors like NO_x and VOCs underwent photochemical reactions downtown, leading to enhanced generation of O₃. The high levels of O₃ were gradually transported from downtown to the Yellow Sea by northerly winds, causing high O₃ accumulations over the ocean. The ‘head-to-head wind’ along the coastline led to rapid increases in O₃ concentration within the sea–land breezes convergence zone (Figure 7). As the sea breeze developed in the

afternoon, it carried O₃-rich air back to land, resulting in the average O₃ concentration pollution in this convergence zone being approximately 100 µg/m³ higher than that in non-convergence zones. These results were consistent with findings found in Ningbo [24] and Tianjin [35].

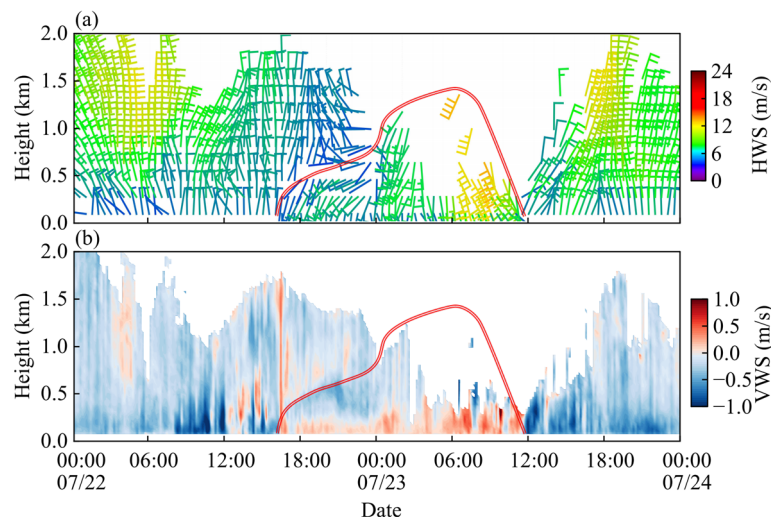


Figure 6. Vertical distributions of the (a) horizontal wind direction and wind speed and (b) vertical wind speed (negative presents downdraught; positive presents updraught) at SB on 22–23 July 2022. (The double red line represents the height of the sea breeze.).

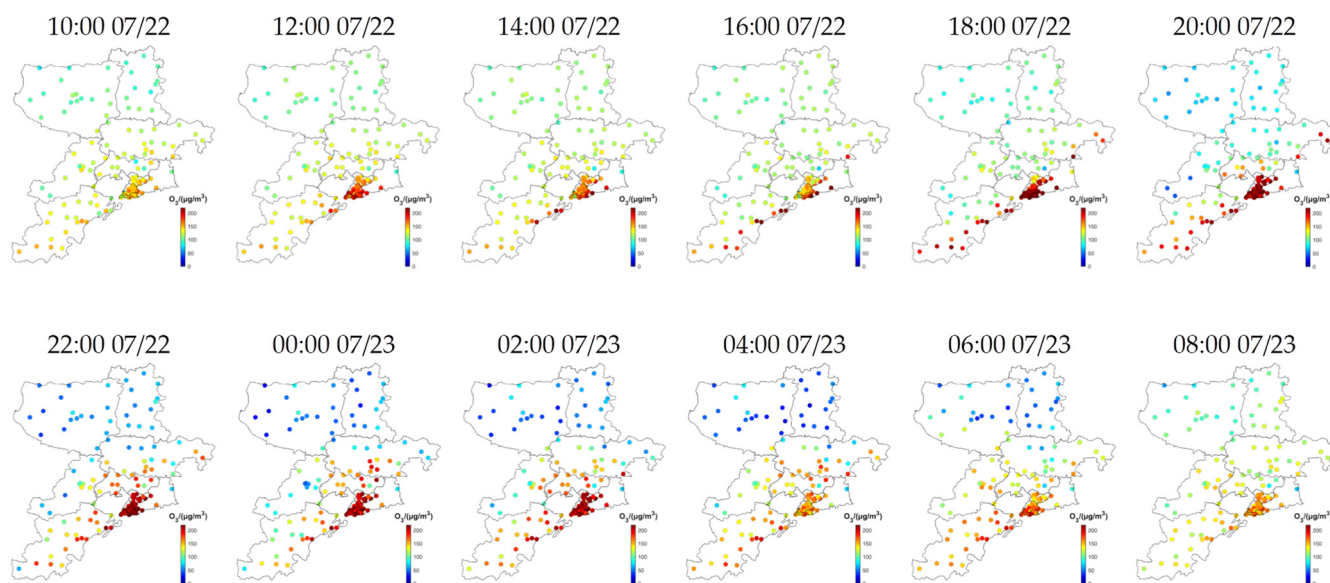


Figure 7. Spatial and temporal distribution of O₃ concentration in Qingdao from 10:00 on 22 July to 08:00 on 23 July 2022.

The sea breeze from the Yellow sea blows horizontally towards inland, with its landward edge defined as the sea breeze front [45]. As the sea breeze strengthened, the sea breeze front advanced northward, leading to a flow of high O₃ concentrations back towards northward inland areas. This caused a distinct spike in O₃ concentrations in downtown areas from south to north between 16:00 and 21:00 (Figure 8), with an average increase of more than 70 µg/m³ within 10 min at multiple sites. Eventually, the sea breeze front moved past downtown, penetrating further inland toward rural areas and resulting in high-concentration O₃ zones migrating from southern downtown to northern rural areas (JZ and JM), demonstrating a time lag in peak O₃ concentration. Similar high-concentration

O₃ zones behind the moving sea breeze fronts were also observed in Hangzhou [25] and New York [46]. Despite NO_x titration, lower boundary layer height, ground temperature inversion, and strong sea breeze induced by ground cyclone, the O₃ concentration remained elevated at night, likely due to backflow of O₃ from offshore areas.

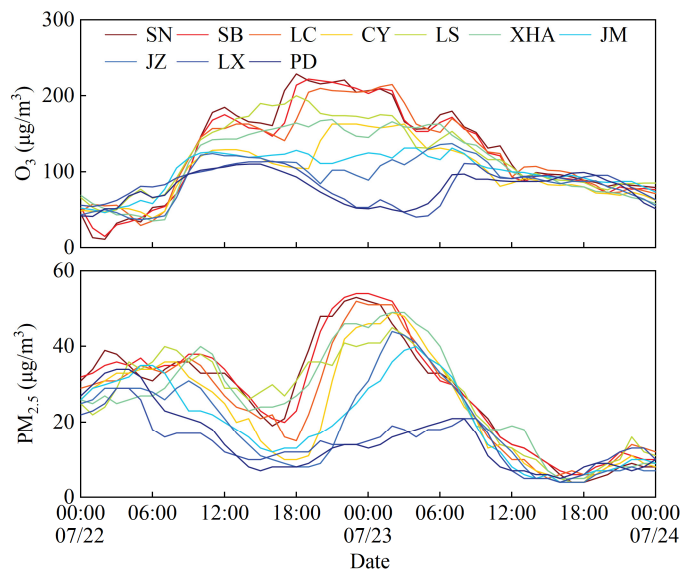


Figure 8. Hourly variation in O₃ concentration for 10 districts in Qingdao on 22–23 July 2022.

By early morning on 23 July, the O₃ concentration started to decline due to wet scavenging. However, a slight increase of around 30 µg/m³ in O₃ was observed around 07:00 in southern downtown, attributed to the retreat of sea breeze and a rapid transition of land breeze.

3.4. Potential Influence of NOE on Particulate Nitrate Formation

From the evening of 22 July to the morning of 23 July, significant positive correlations were observed between O₃ and PM_{2.5} concentrations in the regions of SN, SB, LC, CY, LS, XHA, JZ, and JM (Figure 9), with R values ranging from 0.5 to 0.9. These areas were both influenced by the sea breeze. In contrast, the R values in LX and PD were −0.1 and −0.2, respectively, indicating insignificant negative correlations between the O₃ and PM_{2.5} concentrations. In the convergence zone of the sea and land breezes, the PM_{2.5} concentration increased rapidly, resulting in a combined pollution of PM_{2.5} and O₃ occurring chronologically from south to north as the sea breeze developed.

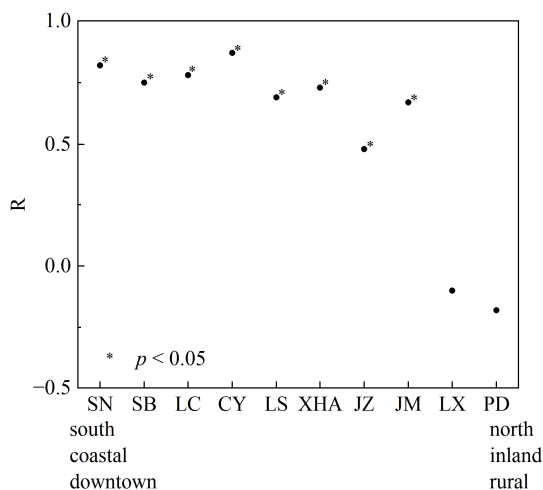


Figure 9. R values between O₃ and PM_{2.5} from the evening of 22 July to the morning of 23 July 2022.

The NOE, coupled with enhanced atmospheric oxidation, would promote the secondary conversion of gaseous precursors into PM_{2.5} [47]. Secondary inorganic ions (SNAs), including nitrate, ammonium, and sulfate, were major components of PM_{2.5}. During the first PM_{2.5} peak around 8:00 on 22 July, the O₃ concentration was 78 µg/m³, and SNA concentrations accounted for 38% of PM_{2.5} (Figure 10). Specifically, nitrate, sulfate, and ammonium accounted for 13.6%, 8.5%, and 13.8% of PM_{2.5}, respectively. In contrast, by the second PM_{2.5} peak around 23:00 on 22 July, the O₃ concentration had increased to 221 µg/m³, and the SNA concentrations rose to 65% of PM_{2.5}. Nitrate, sulfate, and ammonium accounted for 32.3%, 14.9%, and 15.6% of PM_{2.5}, respectively. The proportion of SNA, particularly nitrates, increased notably during nocturnal O₃ pollution period, indicating a potential secondary conversion process.

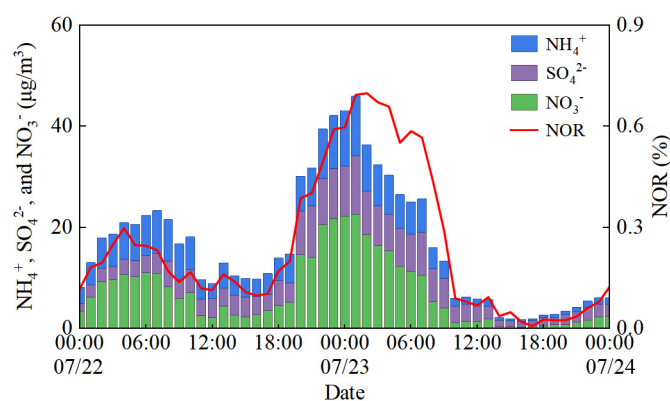


Figure 10. Hourly variation in NH₄⁺, SO₄²⁻, and NO₃⁻ concentrations and NOR in LS on 22–23 July 2022.

Nitrogen oxidation rate (NOR, $\text{NOR} = [\text{NO}_3^-] / ([\text{NO}_3^-] + [\text{NO}_2])$) has been widely used to evaluate the efficiency of the secondary transformation of NO₂ to NO₃⁻ [48]. From 20:00 on 22 July to 02:00 on 23 July, the average O_x (O₃ + NO₂) concentration was (227 ± 16) µg/m³, with NOR values ranging from 0.39 to 0.70, averaging 0.55. This suggested that the nocturnal increase in O₃ concentration likely promoted the generation of secondary particles [13]. Previous studies have indicated that nocturnal O₃ can react with NO_x to produce NO₃ radicals and N₂O₅, which subsequently form HNO₃ and contribute to particulate nitrate formation [47,49]. Additionally, the increase in RH during the nighttime can also promote the reactions of NO₂, O₃, and NO₃ to generate N₂O₅ [50]. Further investigation is required to fully understand the role of NOE in promoting particulate nitrate formation through N₂O₅ heterogeneous hydrolysis.

4. Conclusions

The effects of sea–land breezes on NOE and its potential influence on particulate nitrate formation were investigated using surface air pollutants and meteorological observations, along with continuous 3D wind profiles in Qingdao from 22 to 23 June 2022. On 22 July, O₃ was generated by accelerated local photochemical reactions under a weak, high-pressure synoptic pattern characterized by high T and intense SR. This weak synoptic pattern was favorable for the development of local circulations, such as sea–land breezes. The land breeze initially dominated before noon on 22 July, shifting to the sea breeze in the afternoon, lasting until the next morning, before transitioning back to land breeze again between 09:00 and 12:00 on 23 July. The sea breeze lasted for up to 25 h, penetrating inland as far as 50 km from the coastline and reaching a maximum height of 1.2 km. A “head-to-head wind” convergence zone was created due to the conversion between land breeze and sea breeze, which had a significant influence on the spatial–temporal O₃ distribution and made it more complicated in coastal areas.

The prevailing northerly land breeze first transported high levels of O₃ and other air pollutants from downtown to the Yellow Sea. In the afternoon, the developing sea breeze

brought O₃-rich air mass back inland. However, as the sea breeze strengthened and a strong ground temperature inversion formed at night, the diffusion of air pollutants was inhibited, resulting in elevated O₃ and PM_{2.5} concentrations at midnight. The penetration of the sea breeze front led to further inland transport of O₃ and PM_{2.5}, with a distinct spike in O₃ and PM_{2.5} concentrations observed from southern downtown to northern rural areas between 16:00 on 22 July and 02:00 on 23 July. These findings showed that the nighttime increase in O₃ concentrations and enhanced atmospheric oxidation would promote the secondary conversion of gaseous precursors to PM_{2.5}. It is necessary for further comprehensive research to fully understand the underlying causes of this combined O₃ and PM_{2.5} pollution.

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