



# Article Climate Variability and Floods—A Global Review

#### Zbigniew W. Kundzewicz, Małgorzata Szwed and Iwona Pińskwar \*

Institute for Agricultural and Forest Environment, Polish Academy of Sciences, 60-809 Poznań, Poland

\* Correspondence: iwonp1@wp.pl

Received: 30 May 2019; Accepted: 4 July 2019; Published: 7 July 2019



Abstract: There is a strong inter-annual and inter-decadal variability in time series of flood-related variables, such as intense precipitation, high river discharge, flood magnitude, and flood loss at a range of spatial scales. Perhaps part of this variability is random or chaotic, but it is quite natural to seek driving factors, in a statistical sense. It is likely that climate variability (atmosphere–ocean oscillation) track plays an important role in the interpretation of the variability of flood-related characteristics, globally and, even more so, in several regions. The aim of this review paper is to create an inventory of information on spatially and temporally organized links of various climate-variability drivers with variability of characteristics of water abundance reported in scientific literature for a range of scales, from global to local. The climate variability indices examined in this paper are: El Niño-Southern Oscillations (ENSO), North Atlantic Oscillations (NAO), Atlantic Multi-decadal Oscillation (AMO), and Pacific Decadal Oscillations (PDO). A meta-analysis of results from many studies reported in scientific literature was carried out. The published results were collected and classified into categories after regions, climate variability modes, as well as flood-related variables: precipitation, river flow, and flood losses.

**Keywords:** floods; climate variability; ocean–atmosphere system; precipitation extremes; river discharge; flood loss

#### 1. Introduction

Floods are major weather-related events that continue to cause high economic and human losses all over the Globe, reaching, on average, tens of billions of \$US and thousands of fatalities per year. Hence, an understanding of mechanisms of temporal change and variability of floods and flood-related variables (such as precipitation, river discharge, flood loss) is of vast theoretical and practical interest and importance.

The search for ubiquitous, spatially coherent trends in long time-series of hydro-climatic observation records has been very successful for temperature, less successful for precipitation (including intense precipitation), and even less so for river flow. Actually, no persuading, spatially-organized long-term trends could be detected in records of maximum river discharge.

Analyses of the time series of annual maximum (or peak-over-threshold) values of river discharge at a given station (cross-section) and of flood indices do not show a convincing ubiquitous upward trend. In the absence of a gradual trend, the structure of the time series is dominated by strong, and seemingly rather irregular, inter-annual and inter-decadal variability. An intriguing question emerges—what is the mechanism of variability of flood-related variables? Perhaps there exist some links with recognized climate variability and its indices. Therefore, it is quite natural to examine whether the climate variability expressed via atmosphere–ocean oscillation indices can drive the variability of flood-related indices.

Better understanding of spatially- and temporally-organized links between climate variability patterns and flood hazard and flood risk at a range of spatial and temporal scales would be a considerable

asset. Anticipating regional predisposition to river floods in a particular phase of oscillation in the atmosphere–ocean system, measured by a range of climate variability indices, could help in improving

preparedness to emergency.

The aim of this review paper is to assemble a selective summary of the existing information on links of various climate-variability drivers with variability of characteristics of destructive water abundance at a range of scales. We discuss the modes of climatic variability that may have links to flood-related variables. We review publications reporting on links between variability of intense precipitation, high river discharge and flood indices and natural oscillations in the ocean–atmosphere system, starting from global scale to continental, river-basin, national, regional, to local scale.

#### 2. Modes of Climate Variability

Many indices of inter-annual and inter-decadal climate variability have been proposed in the scientific literature in order to quantify oscillation in the atmosphere–ocean system. Some of them have been studied in the context of links with hydrological variables. Time series of characteristics of climate variability indices of concern may reflect volcanic eruptions and indices of solar irradiance, illustrated by sunspot numbers, as well as various classes of ocean–atmosphere oscillations. The first two indices from the above list are likely the drivers of climate variability. In contrast, ocean–atmosphere oscillations, being manifestations of climate variability, are likely the drivers of climate impact variability, that are the focus of this paper.

Accounting for the heat and mass exchange between the Earth's atmosphere and the oceans is absolutely necessary to understand large-scale meteorological processes and to interpret the climate variability. The latter consists of quasi-periodic oscillations in the climate system, i.e., the alternating occurrence of positive and negative anomalies of meteorological variables such as temperature or air pressure in a specific reference location, or of differences between values of these variables in two specific reference locations. These spatially defined oscillations in the ocean–atmosphere system have been found to affect the climate and various climate impact variables over large areas, including remote regions, via teleconnections.

There are four principal modes of climate variability that, according to a plethora of references, are of particular relevance for links with abundance of water—i.e., with characteristics of intense precipitation, high river discharge, and floods. They are oscillations of quasi-periodicity of several years: ENSO (El Niño-Southern Oscillation) and NAO (North Atlantic Oscillation), as well as of several decades: AMO (Atlantic Multi-decadal Oscillation) and PDO (Pacific Decadal Oscillation). Brief information about these four modes of climate variability is given below.

The ENSO is the sea surface temperature oscillation over the tropical Pacific Ocean. During the warm phase, El Niño, water off the South American coast is warmer than average (in the absence of cold upwelling), while during the cold phase, La Niña, with increased cold upwelling, warm water is further west off the South American coast than usual. Between the warm phase, El Niño, and the cold phase, La Niña, there is a neutral phase. The Southern Oscillation is the accompanying atmospheric (air pressure) component, coupled with the sea surface temperature (SST) change. Many relevant indices have been proposed and some of them were used in references reviewed in this paper. Among the indices are: ONI (Oceanic Niño Index, 3-month running mean of SST anomalies in the Niño 3.4 region, stretching from the 120° W to 170° W astride the equator, five degrees of latitude on either side; for declaration of El Niño or La Niña, the anomalies must exceed ±0.5 °C for at least five consecutive months), Niño 3.4 (5-month running mean, and SSTs must exceed  $\pm 0.4$  °C for a period of six months at least), Niño 1 + 2 (0–10° S, 90° W-80° W), Niño 3 (5° N-5° S, 150° W-90° W), Niño 4 (5° N-5° S, 160° E-150° W), and the Trans-Niño Index (TNI), which is defined as the difference in normalized SST anomalies between the Niño 1 + 2 and Niño 4 regions and the event is classified as a "central Pacific El Niño" or "El Niño Modoki". Also, the Southern Oscillation Index (SOI, https://www.ncdc.noaa.gov/teleconnections/enso/indicators/soi/) has been used, expressing fluctuations in air pressure between the western and the eastern tropical Pacific, calculated based on the differences in air pressure anomaly between Tahiti and Darwin, Australia, [1,2]. The term NAO refers to cyclic air pressure changes in the Icelandic Low and Azores High as well as cyclic displacements of the center of the Azores High. If the (seasonal or monthly) value of the NAO index is positive, there is a stronger air pressure contrast between the Icelandic Low and the Azores High, i.e., the pressure value in the Azores Region is higher than the long-term average, while the corresponding pressure in the Icelandic Low is lower than the long-term average. Several indices were proposed to measure the strength of NAO, related to air pressure in various stations, such as: the Hurrell index—between Lisbon (Portugal) and Stykkishólmur (Iceland), the Rogers index—between Ponta Delgada (Azores) and Akureyri (Iceland), and the Jones index—between Gibraltar and Stykkishólmur. More detailed information on the concept and studies of NAO can be found in Wanner et al. [3].

The PDO is the decadal-scale temperature oscillation of the Pacific Ocean. Sometimes it is called a long-lived ENSO-like Pacific climate oscillation pattern. Two quasi-periodicities have been determined: of 15–25 and 50–70 years. Temperature anomalies (departures from average) in the East Pacific (off the west coast of North America) have an opposite sign to those in the West Pacific, therefore there is some compensation effect and the impact on the global temperature is small [4,5].

The term AMO refers to a coherent pattern of oscillatory changes in the North Atlantic sea-surface temperature (SST) with a period of about 60–90 years. Traditionally, indices of the AMO have been based on average annual SST anomalies in the North Atlantic. Temperature anomalies in the North Atlantic have an opposite sign to those in the South Atlantic, therefore there is a kind of compensation effect and the impact on the global temperature is small. The nature and origin of the AMO is uncertain, and it remains unknown whether it represents a persistent, quasi-periodic driver in the climate system, or merely a transient feature. A deterministic mechanism relying on atmosphere–ocean–sea ice interactions for the AMO was proposed by Dima and Lohmann [6].

There are several data sets of climate variability indices that have been collected at various sources and made publicly available (e.g., Table 1).

Table 1. Data sets of climate variability indices (monthly values) available at the US NOAA (National Oceanic and Atmospheric Administration) ESRL (Earth System Research Laboratory), under the GCOS (Global Climate Observing System) umbrella (https://www.esrl.noaa.gov/psd/gcos\_wgsp/Timeseries/SOI/ or https://www.esrl.noaa.gov/psd/gcos\_wgsp/Timeseries/Nino34/ or https://www.esrl.noaa.gov/psd/gcos\_wgsp/Timeseries/Nino34/ or https://www.esrl.noaa.gov/psd/gcos\_wgsp/Timeseries/PDO/ or https://www.esrl.noaa.gov/psd/gcos\_wgsp/Timeseries/PDO/ or https://www.esrl.noaa.gov/psd/gcos\_wgsp/Timeseries/AMO/).

| <b>Oscillation Pattern</b> | Comment   | Available from |
|----------------------------|---|----------------|
| ENSO                       | SOI, normalized pressure difference between Tahiti and Darwin       | 1866           |
|                            | Niño 3.4 SST Index, area averaged SST from 5° S–5° N and 170–120° W | 1870           |
| NAO                        | Normalized pressure difference between Gibraltar and SW Iceland     | 1821           |
| PDO                        |   | 1900           |
| AMO                        |   | 1871           |

#### 3. In Search for Links between Variability of Flood-Related Variables and Climate Oscillations

The structure of the time series of annual maximum river discharge and flood damage worldwide, does not show a persuading ubiquitous upward trend [7], even if the Clausius–Clapeyron law demonstrates that more water vapor can be stored in the warmer atmosphere, suggesting the increase of the potential for more intense precipitation (that may lead to floods) in the warmer climate. What the time series of annual maximum river discharge shows is a dominating inter-annual and inter-decadal variability that may overshadow gradual (if any) trends. High river discharges tend to be clustered in time [8].

There is a strong multi-annual and multi-decadal variability in flood-related variables. Perhaps part of this variability is random or chaotic, but there are serious arguments supporting the thesis that climate variability plays, generally, an important role in driving flood hazard over many areas of the world. Prudhomme et al. [9] examined whether the same few circulation types are systematically associated with floods at the European scale. Their results showed that, indeed, at the river basin scale and at the continental, European scale, some circulation types have significant positive frequency anomalies with flood occurrence, i.e., they are associated with a predisposition to floods. However, the results were not robust—they were shown to depend on various technicalities. Burn and Arnell [10] investigated the spatial and temporal pattern of high river flow, globally, using annual maximum discharge data for 200 gauging stations. The results of the investigation revealed that diverse regions of the world exhibit synchronous flood responses which can be related to various global climatic conditions, such as atmosphere–ocean oscillations. They also examined global temporal patterns in flooding and identified distinct periods corresponding to increased global flooding activity.

Many research works have been undertaken and reported in literature, aimed at identification of links between the climate variability and various variables related to floods. Climate variability indices can quantify some climate phenomena which can be considered as drivers of variability of such flood-related characteristics, as intense precipitation, high river discharge, flood severity and magnitude (in the sense of Brakenridge, see explanation in [11,12]), as well as flood loss (material damage, population, number of fatalities, and a combined index). Clearly, there are several climatic and non-climatic factors influencing flood risk (understood as probability times consequences, that can be disaggregated into three driving factors: hazard, exposure, and vulnerability). Non-climatic drivers are of terrestrial/hydrological (e.g., land-use and land-cover change, river regulation, and structural flood defenses), as well as social/economic nature. For instance, changes in rivers and their drainage basins, and changes in flood damage potential depend on the socio-economy (demography, wealth, development, GDP). However, the climate variability track is likely to play an important role in the interpretation of the variability in the time series of flood-related characteristics globally and, even more so, in many regions. Hence, it seems opportune to examine variability in recorded time series of flood-related characteristics and climate variability indices and seek spatially and temporally organized links between them. Improved understanding of climate-variability impacts on variability of flood-related characteristics would be of considerable value.

The expectation of discovery of absolute deterministic rules would be naïve, yet existence of probabilistic, frequency-type rules of spatially-organized validity is likely. The global spatial scale can be used, but particular focus can be put on areas where links have been found to be strong and where robust regional laws can be formulated. In many papers, authors claim to have unveiled new teleconnections (i.e., co-variability of variables at distant locations), possibly in asynchronous (e.g., lagged) mode, and/or to have found an important original interpretation.

This paper aims at building an inventory of selected information on links between climate variability indices and variability of flood-related characteristics. It is virtually impossible to embrace all the many thousands of related references. Search for the logical sum of "ENSO" and "flood" gives 32,500 hits in Google Scholar, and 733 hits in Web of Science. Search for "NAO" and "flood" gives 24,200 hits in Google Scholar, and 147 hits in Web of Science, while search for "PDO" and "flood" gives 8,020 hits in Google Scholar, and 94 hits in Web of Science. Search for "AMO" and "flood" gives 19,900 hits in Google Scholar, and 36 hits in Web of Science (updated on 13 May 2019). The gross orientation numbers given above are clearly overestimating the count, for several reasons, e.g., in some entries, Flood is an author's name, while Nao and Amo refer to geographic names and NAO-to National Audit Office. Nevertheless, the number of relevant source items is very high. In brief, the audience is inundated by information on the links between climate variability and variables related to floods, of varying strength and trustworthiness. Hence, some transparent rules of selection of entries to be included in this paper are needed. We give priority to reviewing papers that have gathered many citations, according to bibliometric indices (related to citations count) reported in Web of Science. However, we also include several papers that have not gathered numerous citations yet (some of them being quite recent), but deserve being considered. In result, bibliography contains 120 source items. In this review paper, there is usually reference to one or two papers authored by a single scientist or

co-authored by a group of scientists, even if the same authors may have possibly published many other papers in the field. An attempt was made to select the appropriate paper(s) out of a larger set.

Meta-analysis (analysis of analyses) of results from selected studies reported in scientific literature is offered. The bibliography is collected and classified into categories (after regions, climate variability modes, and flood-related variables: precipitation, river flow, and flood losses). Reported results are assessed and a synthesis of findings is presented. It could improve understanding of the links of four variability modes with abundance of water and inform further studies in the area. Information is given on who studied what and whether approaches turned out to be successes or failures (though, definitely, the latter are less likely to be well represented in the scientific literature).

From the viewpoint of flood variability, the dominating mode of inter-annual climate oscillation is undoubtedly the El Niño-Southern Oscillation (ENSO), with impacts felt most strongly in the Pacific region but also in much of the world via teleconnections. Colloquially speaking, once the occurrence of a strong El Niño is declared, a range of extreme weather events worldwide, including floods, are blamed on this phenomenon in the media. Other, possibly relevant, climate variability indices examined in many references are: AMO (Atlantic Multi-decadal Oscillation), PDO (Pacific Decadal Oscillation), and NAO (North Atlantic Oscillation). Some papers tackle other variability indices, such as Northern Hemispheric Teleconnection Indices (Arctic Oscillation—AO, Polar/Eurasia Pattern, Siberian High-Intensity, Scandinavia Pattern, East Atlantic/Western Russia Pattern, East Atlantic Pattern, Pacific/North American Pattern—PNA, and Northern Hemispheric Circum-global Teleconnection—CGT), as well as Monsoon Indices (Webster–Yang Monsoon Index, South Asian Monsoon Index, East Asian–Western North Pacific Monsoon Index), see [13], Indian Ocean Dipole, see [14], and AAO (Antarctic Oscillation), see [15].

Some studies show that the overall impact of climate variability modes (e.g., ENSO and PDO) depends on the combination of ENSO and PDO. For instance, Wang et al. [16] investigated the inter-decadal modulation of the PDO upon the impact of ENSO on the EAWM (East Asian Winter Monsoon). No robust relationship was found between ENSO and EAWM on the inter-annual timescale, when the PDO is in its high phase. However, when the PDO is in its low phase, ENSO was found to exert a strong impact on the EAWM. Chan and Zhou [17] found that inter-decadal variations in the early (May–June) summer monsoon rainfall over South China are linked to the ENSO and the PDO (Pacific Decadal Oscillation), with wetter monsoon years during the periods of low PDO index.

Statistical relationships have been established between some quasi-periodic ocean–atmosphere oscillations (e.g., ENSO, NAO) and general wetness (e.g., river runoff) indices in various regions of the world (cf. [18,19]). Some research studies [19,20] reviewed connections between high river flow (in winter and spring) and the NAO index in winter over many areas. This could be particularly important e.g., in regions with snowmelt as the dominating flood generation mechanism [21].

#### 4. Review of Links between Climate Oscillations and Flood-Related Variables—Global Studies

The El Niño-Southern Oscillation (ENSO) is perhaps the most commonly studied mode of inter-annual climate variability that was found to have an impact on river flow and flooding. Indeed, many dramatic floods have occurred in some areas during strong warm phases of ENSO (El Niño) and in other areas during strong cold phases of ENSO (La Niña). This climate variability mode largely influences the Pacific region, but its impacts can also be felt in much of the world via teleconnections, possibly of asynchronous (lagged) nature. Chiew and McMahon [18] examined global teleconnections between ENSO and river flow. They found strong and regionally consistent ENSO-streamflow teleconnections in Australia and New Zealand, South and Central America, and weaker signals in some parts of Africa and North America. They also proposed a hypothesis that the streamflow–ENSO relationship can be stronger than the rainfall–ENSO relationship (so that an amplification effect takes place), due to integrative properties of the streamflow process.

It can be stated that the occurrence of a specific stage of climate variability (e.g., El Niño) can open the door to a visit of abundant water, yet floods may or may not come. In fact, floods can happen any time, regardless of the climate variability phase. However, in the area of interest, they can be much more likely during one phase (e.g., El Niño) and perhaps more unlikely during another phase (e.g., La Niña). Sun et al. [22] postulated that the effect of ENSO on extreme precipitation is asymmetric, in the sense that a strong, significant effect is only for a single ENSO stage, and much weaker (or none) for another.

There is no doubt that the random component of climate variability as well as of variability of flood-related variables is strong. Nevertheless, Ward et al. [23,24] examined statistical links between ENSO and flooding and arrived at interesting and promising frequency-type results globally. They demonstrated that the ENSO quasi-cycle of natural variability likely has serious impacts on flood hazard in many regions of the world and called for further research to reduce uncertainty. In contrast, more recently, Emerton et al. [25] showed that the interpretation of likelihood of ENSO-driven flood hazard is more complex than perceived and reported in literature.

Ward et al. [23,24] arrived at an updated global map of prevalence of wetness (intense precipitation, flooding) for particular phases of the ENSO oscillation. It is clear that various flood-related variables may react differently to climate variability indices in different areas. The behavior of the time series of river discharge may largely differ from the time series of intense precipitation, as per suggestion of integrative properties due to Chiew and McMahon [18]. Large rivers have basins covered by multiple grid cells and the transformation of rainfall into river runoff modulates the signal and introduces a transport lag, not to forget human influence (anthropogenic effects in drainage basins and river beds) that can be substantial. However, it is not only the scale problem. In many river basins, the flow regime of the main stream differs from the regime of tributaries (e.g., pluvial versus glacial, snowmelt, or rain on snow).

There is a rich body of research aimed at connecting floods with climatic variability at the continental scale. Many relevant studies have been undertaken in Asia, North America, South America, Australia, and Europe, using ENSO, NAO, AMO, and PDO indices. Hodgkins et al. [26] examined climate-driven variability in the occurrence of major floods across two continents (North America and Europe), only in minimally altered catchments, to focus on climate-driven changes rather than changes due to catchment alterations. Temporal changes in the occurrence of major floods were shown to be dominated by multi-decadal variability (AMO) rather than by long-term trends.

The question whether one individual flood event (or a train of floods in a flood season) can be meaningfully associated with the occurrence of a particular phase of climate variability is not properly posed. A more appropriate question is whether the occurrence of a particular phase of climate variability can cause the risk of experiencing floods (predisposition to abundance of wetness) to be significantly higher or lower than average. The reported strength of correlation may vary with time (strengthening or weakening over various periods), with possible asymmetry [22]. Even the sign of correlations may change with time. This is the meaning of the so called "epochal behavior" recognized for the Mekong by Räsänen and Kummu [27]. Ward et al. [23,24] speculated that climate change may possibly be modifying the strengths of teleconnections.

Table 2 presents the essential information extracted from selected references, about the links between climate variability and floods, resulting from global studies.

| Climate<br>Oscillation | Reference              | Principal MESSAGE  | Aspect |
|------------------------|------------------------|--|--------|
| ENSO                   | Emerton et al.<br>[25] | Likelihood of increased or decreased flood hazard during ENSO<br>events is much more complex than is often perceived and<br>reported. Probabilities vary greatly across the Globe, with large<br>uncertainties inherent in the data. | R      |

**Table 2.** Review of links between climate oscillations and floods, reported in global studies. Aspects of reference in the rightmost column read: P—precipitation, R—runoff (or river discharge), F—flood.

| Climate<br>Oscillation | Reference                   | Principal MESSAGE  | Aspect |
|------------------------|-----------------------------|--|--------|
|                        | Ward et al. [28]            | The duration of flooding appears to be more sensitive to the<br>ENSO index than is the case for flood frequency. At the globally<br>aggregated scale, floods are significantly longer during both El<br>Niño and La Niña years, compared to neutral years. At the scale<br>of individual river basins, strong correlations were found<br>between ENSO and both flood frequency and duration for a large<br>number of basins, with generally stronger correlations for flood<br>duration than for flood frequency.          | F      |
| ENSO                   | Ward et al. [24]            | The ENSO exerts strong and widespread influences on both flood<br>hazard and flood risk. Reliable anomalies of flood risk exist<br>during El Niño or La Niña years, or both, in basins spanning<br>almost half (44%) of Earth's land surface.  | F      |
| ENSO                   | Ward et al. [23]            | Over 1958–2000, ENSO exerted a significant influence on annual<br>floods in river basins covering over a third of the world's land<br>surface, and its influence on annual floods has been much greater<br>than its influence on average flows. There are more areas in<br>which annual floods intensify with La Niña and decline with El<br>Niño than vice versa. However, in many regions the strength of<br>the relationships between ENSO and annual floods was found to<br>be non-stationary during the study period. | R      |
|                        | Chiew and<br>McMahon [18]   | Strong and regionally consistent ENSO-streamflow<br>teleconnections were found in Australia and New Zealand,<br>South and Central America, and weaker signals in some parts of<br>Africa and North America.  | P, R   |
| AMO, ENSO              | Najibi and<br>Devineni [29] | The flood frequency across the Globe and the tropics can be<br>largely linked to the decadal and multi-decadal climate<br>variability, but the flood duration could not be associated with<br>any of these climate factors. Only extreme flood duration can be<br>partially associated with AMO for the Globe and the tropics and<br>ENSO for the southern subtropics and mid-latitudes.   | F      |
|                        | Dong and Dai<br>[30]        | Significant correlations between the PDO and local P are<br>observed over Eastern Australia, Southern Africa, and the<br>Southwest U.S. Over Northeastern Australia, P versus ENSO<br>correlations are stronger during the PDO cold phases.  | Р      |
| 120                    | Mantua and<br>Hare [4]      | During the warm PDO phase—anomalously wet conditions were<br>found to occur in: Southeast Brazil, South-central America, West<br>Australia, Coastal Gulf of Alaska, Southwest USA, and Mexico  | Р      |

Table 2. Cont.

# 5. Review of Links between Climate Oscillations and Flood-Related Variables—Continental, Regional, and Local Studies

Tables 3–8 present the most essential information extracted about the links between climate variability and floods, resulting from continental (and bi-continental) as well as regional or local studies, as reported in scientific literature. The tables are organized according to continental coverage of studies. Table 3 refers to Asia, Table 4—to North America, Table 5—to South America, Table 6—to Australia, Table 7—to Europe, and Table 8—to Africa. Large regions within continents are placed first, then large river basins.

A few studies examine more than one mode of climate variability. This is explicitly noted in the tables. If a study refers to more than one continent, but does not cover the Globe, it is mentioned in several tables.

Figure 1 illustrates regional predisposition to abundance of water for various phases of four modes of climatic variability. The map in Figure 1 summarizes the findings of continental, regional, to local studies reported in Tables 3–8 of this section.

| Climate<br>Oscillation | Range                                   | Reference                      | Principal Message  | Aspect |
|------------------------|---|--------------------------------|--|--------|
|                        | East Asia                               | Lau and Weng<br>[31]           | Analysis of ENSO indices and extreme rainfall statistics<br>between regions in East Asia and North America<br>suggests the occurrence of two teleconnection patterns.<br>The first pattern is associated with enhanced rainfall<br>over the Yangtze River region. The second pattern<br>shows enhanced rainfall anomalies over the River<br>Huaihe, northeastern, and southern China. It is<br>hypothesized that the first pattern may be influenced by<br>El Niño in the preceding spring but becomes<br>increasingly decoupled from tropical SST during the<br>summer and fall. However, the second pattern has no<br>significant relationship with El Niño. | Р      |
|                        | South-east<br>Asia                      | Kripalani and<br>Kulkarni [32] | Extreme flood situations tend to occur when the internal<br>epochal behavior (epochs of the above- and<br>below-normal rainfall near the Equator tend to last for<br>about a decade, over the tropical regions, away from the<br>Equator—for about three decades) and the external<br>forcings of La Niña events are phase-locked.   | Р      |
|                        | China                                   | Lv et al. [33]                 | During Eastern Pacific El Niño and Central Pacific El<br>Niño decaying years and in La Niña developing years,<br>"wet" indices (the number of very high 24 h<br>precipitation, the maximum 24 h rainfall and the number<br>of consecutive wet days) all increased, suggesting an<br>increased risk of flooding.  | Р      |
| ENSO                   | Southwest<br>China                      | Cheng et al. [34]              | Extreme wet events occur during El Niño and La Niña<br>years, with similar frequency. The frequencies of extreme<br>wet events are lower than extreme dry events since 2001<br>during El Niño and La Niña years. The highest<br>percentage levels of the frequency of the occurrence of<br>wet events mainly occur during El Niño years.   | Р      |
|                        | River Yangtze                           | Wang and Yuan<br>[35]          | Extreme pluvial floods across the Yangtze River basin in<br>the summer of 2016 were strongly connected with<br>intense atmospheric moisture transport after a strong El<br>Niño winter. The predictability of precipitation and<br>moisture flux is higher in post-El Niño summers than in<br>post-La Niña ones, especially for flooding events. As<br>compared with extreme precipitation, the potential<br>detectability of extreme moisture flux increases by 20%<br>in post-El Niño summers.   | Р      |
|                        | River Yangtze                           | Zhang et al. [36]              | Different phase relations were found between annual<br>maximum streamflow of the River Yangtze and ENSO in<br>the lower, the middle and the upper Yangtze basin.<br>In-phase relations were detected between annual<br>maximum streamflow of the Lower Yangtze Basin and<br>anti-phase relations for the Upper Yangtze Basin.  | R      |
|                        | River Yangtze                           | Tong et al. [37]               | In the basin of the middle and lower River Yangtze, El<br>Niño episodes show a close relation with flood events;<br>for the Upper Yangtze Basin, La Niña episodes are in<br>closer relation with floods. The shorter the interval of<br>ENSO events, the sooner the following flood responds.<br>The longer-duration ENSO events in the Pacific result in<br>a longer flood period.  | R      |
|                        | Upper and<br>mid–lower<br>Yangtze River | Ye and Wu [38]                 | The upstream floods in the Yangtze River Basin<br>correspond to the development phase of La Niña; the<br>mid-lower reaches tend to be linked with the decaying<br>El Niño.   | R      |
|                        | River Mekong                            | Räsänen and<br>Kummu [27]      | Precipitation and discharge decreased (increased)<br>during El Niño (La Niña) and the annual flood period<br>was shorter (longer) during El Niño (La Niña). Also,<br>decadal variations in the ENSO-discharge relationship<br>were found.  | P, R   |

**Table 3.** Review of links between climate oscillations and floods for Asia, as reported in literature.Aspects of reference in the rightmost column—as in Table 2.

| Climate<br>Oscillation | Range                               | Reference                    | Principal Message   | Aspect |
|------------------------|-------------------------------------|------------------------------|---|--------|
|                        | Pakistan                            | Iqbal and<br>Hassan [39]     | After 1970, in Pakistan, the highest percentage of floods<br>per year occurred during El Niño, non-ENSO, and<br>positive Indian Ocean Dipole years.   | F      |
|                        | Japan                               | Higashino and<br>Stefan [40] | Annual maximum precipitation and maximum flood<br>discharges have become substantially higher in phase<br>with the ENSO.  | P, R   |
| ENSO                   | Korea                               | Lee et al. [14]              | The tropical ENSO forcing has a coherent teleconnection<br>with September and November–December precipitation<br>patterns, while the Indian Ocean Dipole is identified as a<br>driver for precipitation variability in September and<br>November.   | Р      |
|                        | Bangladesh                          | Chowdhury [41]               | While the relation between SOI and rainfall is strong in<br>the upstream Ganges-Brahmaputra-Meghna basins in<br>India, it offers limited applicability in the context of the<br>Bangladesh climate.   | Р      |
|                        | River Huai,<br>China                | Zhang et al. [42]            | Canonical ENSO regimes tend to increase precipitation<br>in spring and autumn, and especially in winter. Also, in<br>summer, an increase is observed during the canonical<br>ENSO periods. Water vapor flux is more abundant<br>during warm episodes of canonical ENSO events in<br>spring and summer than during cold episodes.  | Р      |
|                        | Poyang Lake<br>basin, China         | Liu et al. [43]              | The ENSO (and IOD) are the major climate indices that significantly correlated with the magnitude and frequency of floods.  | R      |
|                        | China                               | Cao et al. [44]              | Influence of five El Niño–Southern Oscillation (ENSO)<br>types (central Pacific warming—CPW, eastern Pacific<br>cooling—EPC, eastern Pacific warming—EPW,<br>conventional ENSO, and ENSO Modoki) on rainy-season<br>precipitation in China was studied. Precipitation<br>anomaly was shown to reach up to 30% above average<br>precipitation during decaying CPW and EPW phases.  | Р      |
|                        | Rivers Dez<br>and Karun,<br>SW Iran | Saghafian et al.<br>[45]     | The El Niño phenomenon intensifies March–April floods<br>compared with neutral conditions. The opposite is true<br>in La Niña conditions. The degree of the effect is more<br>intense in the El Niño period.  | R      |
|                        | Southern<br>Mekong<br>region        | Delgado et al.<br>[46]       | A connection was found between the Western North<br>Pacific Monsoon (WNPM) and the discharge in the<br>Southern Mekong region. In the frequency domain, the<br>inter-annual to decadal variance of the Mekong<br>discharge closely follows that of WNPM. It was found<br>that PDO plays a role in enhancing the inter-annual<br>variability of WNPM and that WNPM is more variable<br>during warm phases of PDO.                                  | R      |
| PDO                    | East River<br>Basin, China          | Zhang et al. [47]            | The PDO influences the variance of annual maximum<br>streamflow. However, hydrological regulation by<br>storage reservoir in the basin overshadows the<br>influences of climate indices on flood processes.   | F      |
|                        | China                               | Deng et al. [48]             | Large-scale climate indices were found to have more<br>influence on frequency rather than intensity of extreme<br>precipitation over China.   | Р      |
|                        | Yangtze River<br>Basin              | Xiao et al. [49]             | Influences of climate variability indices on the seasonal<br>occurrence and intensity of precipitation events are<br>complex. The negative PDO event tends to increase the<br>spring occurrence of precipitation events in the<br>southwestern part of the Yangtze River Basin, while the<br>positive ENSO event a year earlier tends to increase the<br>spring intensity of precipitation events in the east part of<br>the Yangtze River Basin. | Р      |

### Table 3. Cont.

| Table 5. Com. | Tab | le 3. | Cont. |
|---------------|-----|-------|-------|
|---------------|-----|-------|-------|

| Climate<br>Oscillation | Range       | Reference          | Principal Message   | Aspect |
|------------------------|-------------|--------------------|---|--------|
| NAO                    | Middle East | Cullen et al. [50] | Flooding period on Middle East rivers is rainfall-driven<br>in December-March, regulated on inter-annual to<br>decadal timescales by NAO. | R      |

**Table 4.** Review of links between climate oscillation and floods for North America, as reported in theliterature. Aspects of reference in the rightmost column—as in Table 2.

| Climate<br>Oscillation | Range                     | Reference                    | Principal Message   | Aspect |
|------------------------|---------------------------|------------------------------|---|--------|
|                        | US                        | Cayan et al. [51]            | During SOI-(El Niño) years, the occurrence of two-<br>and three-day above-threshold precipitation events<br>has increased in the Southwest US (by twice as<br>frequent as during neutral and SOI+ period) and<br>decreased in the Northwest (to nearly half of the<br>occurrence during SOI+ years). In El Niño years, days<br>with high precipitation and streamflow are more<br>frequent than average over the Southwest US and less<br>frequent over the Northwest. Nearly the opposite<br>pattern is associated with SOI+ (La Niña) episodes.   | P, R   |
|                        | Western US                | Corringham and<br>Cayan [52] | Links between ENSO and flood losses across the<br>western United States illustrated by US National Flood<br>Insurance Program (NFIP) daily claims and losses<br>were found. In coastal southern California and across<br>the southwest, El Niño conditions were found to be<br>associated with more frequent and higher magnitudes<br>of insured losses (less frequent and lower for La Niña).<br>In the Pacific Northwest, the opposite pattern appears,<br>though the effect is weaker, and less spatially coherent.  | F      |
| ENSO                   | US Midwest                | Lau and Weng<br>[31]         | There is a pattern associated with above-normal<br>rainfall over the US Midwest (as well as in a part of<br>Asia, over the Yangtze River region) and reduced<br>rainfall over the US Mid-Atlantic coast. It may be<br>influenced by El Niño in the preceding spring but<br>becomes increasingly decoupled from tropical SST<br>during the summer and fall (see Lau and Weng, 2002<br>in Table 3).   | Р      |
|                        | South Great<br>Plains, US | Wang et al. [53]             | A developing El Niño tends to increase late-spring<br>precipitation in the South Great Plains and this effect<br>has intensified since 1980.  | Р      |
|                        | Lower<br>Mississippi      | Munoz and Dee<br>[54]        | In the 6–12 months preceding a flood, El Niño<br>generates a positive precipitation anomaly over the<br>lower Mississippi basin   | R      |
|                        | River Ohio                | Nakamura et al.<br>[55]      | The March–May El Niño 3.4 index is weakly correlated with the years of the 20 historical spring floods events.  | R      |
|                        | California<br>Coast       | Andrews et al.<br>[56]       | South of 35° N along the California coast, floods are<br>significantly larger during an El Niño phase than<br>during a non–El Niño phase. North of 41° N, the<br>annual peak floods during an El Niño phase were less<br>than 70% of the geometric mean of annual peak floods<br>during a non-El Niño phase. For exceedance<br>probabilities from 0.50 to 0.02, the ratio of the El Niño<br>to non-El Niño floods varies from greater than 10 near<br>32° N to less than 0.7 near 42° N. Latitude explains<br>much of the observed variation in the relative<br>magnitude of El Niño versus non-El Niño floods over<br>the range of exceedance probabilities. | R      |
| NAO                    | Eastern North<br>America  | Whan and<br>Zwiers [57]      | The observed influence of the NAO on extreme winter<br>precipitation is largest in eastern North America, with<br>the likelihood of an extreme rainfall event increased in<br>the south under the positive phase of the NAO.  | Р      |

| Climate<br>Oscillation | Range  | Reference                 | Principal Message  | Aspect |
|------------------------|--|---------------------------|--|--------|
|                        | River Missouri                                 | Wang et al. [58]          | The NAO climate forcing made a contribution to the magnitude of the dramatic 2011 flood event.   | F      |
| NAO                    | River Missouri                                 | Nasser et al. [59]        | Joint relationships of lagged Pacific/North American<br>(PNA) pattern and NAO with flood durations and<br>volumes were investigated. At a 15-day lag time, a<br>clear nonlinear dependence was found. During the<br>long duration floods (D > 30 days), NAO and PNA are<br>anomalous and anti-correlated, whereas the NAO and<br>PNA anomalies were found to be mostly under neutral<br>conditions for the short duration floods (D < 30 days).  | R      |
|                        | USA  | Villarini et al.<br>[60]  | Over the central United States, the largest flood peaks<br>tend to occur during the neutral and negative ENSO<br>phases, with limited activity during El Niño years  | R      |
|                        | Western and<br>Central US                      | Dai [61]                  | Wet periods over much of the West and Central US,<br>especially the Southwest were found to be associated<br>with the warm phases of PDO.  | Р      |
|                        | Western<br>Canada                              | Gurrapu et al.<br>[62]    | Higher magnitude floods are more likely during the negative phase of PDO than during the positive phase.   | R      |
| PDO                    | Northern<br>Canada                             | Burn et al. [63]          | Negative correlations of PDO index with annual<br>maximum flood date and spring freshet date: during<br>the warm phase of PDO, snowmelt shifts towards<br>earlier spring and during the cool phase—later<br>towards summer. Also, PDO was found to be<br>positively correlated with winter flows and negatively<br>with summer flows.  | R      |
|                        | Missouri River<br>Basin                        | Wang et al. [58]          | The PDO was an important precursor toward<br>anticipating major flood events in the Missouri River<br>Basin in 2011.   | R      |
|                        | Oregon, US                                     | Beebee and<br>Manga [64]  | The DO is correlated with spring snowmelt timing and<br>magnitude, spring and summer runoff, and timing of<br>annual floods  | R      |
|                        | Nyack<br>floodplain,<br>Montana, US            | Whited et al.<br>[65]     | Large magnitude floods and greater frequency of moderate floods were associated with the cooling phases of PDO.  | R      |
|                        | Florida  | Enfield et al.<br>[66]    | During the positive AMO phase, management<br>priorities shift towards flood protection for the region<br>surrounding the Lake Okeechobee (Florida).  | R      |
| АМО                    | River<br>Matawin,<br>Canada                    | Fortier et al. [67]       | Flood flows (magnitude and frequency) are correlated<br>negatively to AMO upstream from the Matawin dam<br>where the winter floods dominate, but positively<br>downstream from the dam (spring floods). The AMO<br>is the only climatic index that influences the<br>inter-decadal variability of heavy floods produced by<br>spring snowmelt.   | R      |
|                        | Tabasco and<br>Chiapas,<br>Southeast<br>Mexico | Valdes-Manzanilla<br>[68] | Many floods coincided with the warm (positive) phase<br>of AMO. The odds ratio showed that floods were 1.9<br>times more likely to occur where the AMO index was<br>positive.  | R      |
| ENSO, AMO              | River<br>Mississippi                           | Munoz et al.<br>[69]      | Multi-decadal trends of flood hazard on the lower<br>Mississippi River are strongly modulated by<br>dynamical modes of climate variability, particularly<br>ENSO and AMO, but the artificial channelization has<br>greatly amplified flood magnitudes over the past<br>century. A multi-proxy reconstruction of flood<br>frequency and magnitude spanning the past 500 years,<br>reveals that the magnitude of the 100-year flood has<br>increased by 20%, with about 75% of this increase<br>attributed to river engineering. | R      |

| Climate<br>Oscillation | Range                           | Reference                        | Principal Message   | Aspect |
|------------------------|---------------------------------|----------------------------------|---|--------|
| ENSO, AMO              | St. Lawrence<br>River, Canada   | Assani et al. [70]               | On the north shore of the St. Lawrence River Basin,<br>magnitude of annual flow and of flooding are correlated<br>to AMO. On the south shore they are correlated to ENSO.   | R, P   |
|                        | US                              | Hamlet and<br>Lettenmaier [71]   | Climate variability associated with PDO and ENSO<br>affects flood risk. The largest changes are associated<br>with years when PDO and ENSO are in phase,<br>particularly in the southwest US.   | F      |
| ENSO, PDO              | Western US                      | McCabe and<br>Dettinger [72]     | In the western US, ENSO teleconnections with winter<br>precipitation are strong when SOI and NINO3 are<br>out-of-phase and PDO is negative. ENSO<br>teleconnections with precipitation in the western US are<br>weak when SOI and NINO3 are weakly correlated and<br>PDO is positive.   | Р      |
|                        | Southwest<br>Canada             | Gobena and Gan<br>[73]           | Basins in Southwest Canada exhibit a strong response to<br>ENSO forcing. Three of the five regions exhibited wet<br>conditions with a consistency better than 75% during the<br>spring and/or summer period following the onset of La<br>Niña events. Also, ENSO-related responses were found<br>to be modulated by PDO: the interaction between ENSO<br>and PDO was constructive when the two are in phase<br>(i.e., during El Niño and warm PDO, as well as La Niña<br>and cool PDO).                               | R      |
|                        | River<br>Similkameen,<br>Canada | Jain and Lall [74]               | Statistically significant relationship between annual<br>maximum floods and both winter NINO3 and<br>PDO indices  | R      |
|                        | River<br>Blacksmith<br>Fork, US | Jain and Lall [75]               | The ENSO and PDO climate indices explain 40% of the variance of the annual maximum flood series   | R      |
|                        | Multiple<br>gauges              | Hodgkins et al.<br>[26]          | Significant negative relations between floods and AMO at large (>1000 km <sup>2</sup> ) North American catchments.  | R      |
|                        | Southern<br>Mexico              | Méndez and<br>Magaña [76]        | Precipitation was greater when the PDO index was in its negative phase  | Р      |
| AMO, PDO               | The Upper Rio<br>Grande, US     | Pascolini-Campbel<br>et al. [77] | laThe highest ranked streamflow years clustered during positive phases of PDO.  | R      |
|                        | River<br>Papaloapan,<br>Mexico  | Valdes-Manzanilla<br>[78]        | Flooding had an in-phase relationship with AMO and PDO, especially when AMO and PDO were in their positive or warm phases.  | R      |
| AMO, NAO               | St. Lawrence<br>River, Canada   | Mazouz et al.<br>[79]            | The timing of spring floods is correlated with NAO<br>(positive correlation) and with AMO<br>(negative correlation).  | R      |
| NAO, AMO,<br>ENSO, PDO | Conterminous<br>USA             | Archfield et al.<br>[80]         | Time series of frequency, peak magnitude, duration, and<br>volume of flood events at a gauge were correlated to the<br>climate variability indices. Little geographic cohesion<br>was noted, apart from the ENSO that has statistically<br>significant correlations to flood duration and volume for<br>approximately 25% of the 345 streamflow stations<br>analyzed across the USA. Other climate indices (NAO,<br>AMO, PDO) do not show widespread and strong<br>correlations with floods (aside from New England). | R      |
|                        | Central US                      | Mallakpour and<br>Villarini [81] | Climate variability can play a significant role in<br>explaining the variations in the frequency of flooding<br>over the central United States. Different climate-modes<br>are related to the frequency of flood events over different<br>parts of the domain and for different seasons.  | R, P   |

# Table 4. Cont.

| Climate<br>Oscillation | Range                            | Reference                    | Principal Message   | Aspect |
|------------------------|----------------------------------|------------------------------|---|--------|
| ENSO                   | South America                    | Isla [82]                    | South American rivers can be classified into<br>ENSO-affected and ENSO-dominated for those within<br>the Arid Diagonal that are exclusively subject to high<br>discharges during those years.   | R      |
|                        | Uruguay,<br>Argentina,<br>Brazil | Isla and Toldo jr.<br>[83]   | The ENSO-triggered floods caused significant impacts<br>on the economies of Argentina, Brazil, and Uruguay.<br>Floods occur approximately simultaneously in<br>subtropical watersheds of Brazil and high-latitude<br>watersheds of Northern Patagonia.  | F      |
|                        | Central<br>America,<br>Guatemala | Guevara-Murua<br>et al. [84] | For the annual rainy season (May to October), only<br>one-quarter of the wet years are coincident with La Niña.<br>The percentage of wet events occurring during a La Niña<br>year rises to 40% for early part of the rainy season (May<br>to July)   | P, R   |
|                        | River Parana                     | Depetris [85]                | The River Parana exhibits an active teleconnection<br>between extreme flooding and the occurrence of ENSO<br>events.  | R      |
|                        | River Parana                     | Prieto [86]                  | The ENSO is an important factor in determining<br>inter-annual rainfall variability in South America.<br>In 1904–2000, 11 out of 16 floods occurred during El<br>Niño events.   | Р      |
|                        | Ecuador, Peru                    | Arteaga et al.<br>[87]       | Strong El Niño events bring high rainfall and floods in<br>Ecuador and northern Peru.   | P, R   |
|                        | Atacama<br>Desert, Chile         | Houston [88]                 | Floods in the central and northern Atacama Desert in<br>summer are associated with La Niña, but they are<br>caused by the summer melt of the previous winter's<br>snow in the southern Atacama Desert which are<br>associated with El Niño. Floods along the coastal<br>Atacama Desert in winter are associated with El Niño. | Р      |
|                        | Distrito<br>Federal, Brazil      | Borges et al. [89]           | Extreme rainfall indices are associated, in a complex way,<br>with La Niña and Madden-Julian oscillation.   | Р      |

**Table 5.** Review of links between climate oscillation and floods for South America, as reported in the literature. Aspects of reference in the rightmost column—as in Table 2.

**Table 6.** Review of links between climate oscillation and floods for Australia, as reported in the literature. Aspects of reference in the rightmost column—as in Table 2.

| Climate<br>Oscillation | Range  | Reference                  | Principal MESSAGE  | Aspect |
|------------------------|--|----------------------------|--|--------|
|                        | New South<br>Wales and<br>Southern<br>Queensland | Micevski et al.<br>[90]    | The PDO modulated the flood risk in New South Wales<br>and Southern Queensland, with flood quartiles<br>increased by a factor of approximately 1.7 during PDO<br>negative periods.   | R      |
| PDO                    | New South<br>Wales                               | Franks [91]                | Identified change in flood frequency corresponds directly to an observed shift in SST and mean circulation   | R      |
|                        | New South<br>Wales                               | Franks and<br>Kuczera [92] | Marked multi-decadal variability is also present in flood extremes   | R      |
| enso, pdo              | Eastern<br>Australia                             | Verdon et al.<br>[93]      | Rainfall and streamflow for Queensland and New South<br>Wales are significantly enhanced during La Niña. On a<br>multi-decadal time scale the negative phase of PDO is<br>associated with "wetter" conditions: La Niña rainfall<br>and streamflow are further magnified by occurring in<br>the PDO negative phase. | P, R   |
|                        | Queensland                                       | Cai and van<br>Rensch [94] | The devastating southeast Queensland flood and the<br>associated extreme rainfall in January 2011 were<br>accompanied by an extraordinarily strong La Niña. The<br>regional summer rainfall is affected by the ENSO cycle,<br>but modulated by PDO.  | Р      |

| Climate<br>Oscillation | Range              | Reference        | Principal MESSAGE  | Aspect |
|------------------------|--------------------|------------------|--|--------|
| ENSO, PDO              | New South<br>Wales | Kiem et al. [95] | The PDO modulation of ENSO events leads to<br>multi-decadal epochs of elevated flood risk. The flood<br>risk in Eastern Australia is higher when the PDO is in a<br>negative phase. PDO modulation of La Niña is related to<br>rainfall flooding in Eastern Australia and affects both the<br>flood magnitude and frequency. Cool-phase PDO<br>increases flood risk. | R, F   |

Table 6. Cont.

**Table 7.** Review of links between climate oscillation and floods for Europe, as reported in the literature. Aspects of reference in the rightmost column—as in Table 2.

| Climate<br>Oscillation | Range   | Reference                               | Principal Message   | Aspect |
|------------------------|---|---|---|--------|
|                        | Europe  | Wrzesiński and<br>Paluszkiewicz<br>[20] | The winter floods typical of the rivers of Western and<br>Southern Europe are markedly higher in a negative<br>NAO phase. In the North and Northeast regions, a<br>positive correlation can be noted between flows and<br>the NAO indices in the winter season.   | R      |
|                        | Northern<br>Europe,<br>Atlantic<br>countries of<br>Europe | Zanardo et al.<br>[96]                  | Links between NAO and catastrophic floods in Europe<br>were demonstrated. The majority of winter (summer)<br>floods in Northern Europe were found to occur during<br>a positive (negative) NAO phase. Similar patterns<br>were observed in individual Atlantic countries. The<br>average flood loss during opposite NAO states were<br>found to differ by up to 50%.  | R, F   |
|                        | Southern<br>Europe  | Benito, et al. [97]                     | See Table 8   | R      |
|                        | Western<br>Iberian<br>Peninsula                           | Benito, et al. [98]                     | Periods with more frequent floods in the western<br>Iberian region coincide with transitions to cool and<br>wetter conditions, and persistent negative NAO mode.  | R      |
| NAO                    | North-eastern<br>Iberian<br>Peninsula                     | Moreno et al.<br>[99]                   | The NAO variability at a decadal scale, modulated by<br>fluctuation in solar radiation at a centennial-scale, are<br>the two forcing mechanisms proposed to generate the<br>rainfall in the Taravilla Lake catchment. The Taravilla<br>Lake and Tagus River records (with the paleo-flood<br>history) suggest a relationship between the occurrence<br>of extreme hydrological events and the North Atlantic<br>Oscillation for the Northeastern Iberian Peninsula. | P, R   |
|                        | UK  | Huntingford et<br>al. [100]             | Winter runoff and flood indicators have been linked to<br>the NAO index. In the 1990s, it was strongly positive<br>and associated with a cluster of winter floods events,<br>contrasting with the early 1960s for which it was<br>negative and which were correspondingly flood poor.   | R, F   |
|                        | UK  | Hannaford and<br>Marsh [101]            | In western UK, high-flow indicators are correlated<br>with NAO. The increasing high-flow trends since the<br>1960s have parallels with observed changes in extreme<br>rainfall and they may reflect an influence of<br>multi-decadal variability related to NAO.  | P, R   |
|                        | Lower Rhine   | Toonen et al.<br>[102]                  | Variability of flood activity correlates with NAO.<br>Winter NAO+ phases align with increased flood<br>frequency of the Lower Rhine, and the largest events<br>of the last six centuries were all generated in NAO+<br>periods.   | R      |
|                        | Swiss Alps  | Schulte et al.<br>[103]                 | Episodes of major floods in the central Swiss Alps<br>were triggered by negative SNAO (Summer NAO)<br>phases during the last pulses of the Little Ice Age (AD<br>1817–1851 and 1881–1927 flood clusters) and by<br>positive SNAO phases during the warmer climate<br>since 1977.  | R      |

| Climate<br>Oscillation | Range              | Reference               | Principal Message   | Aspect |
|------------------------|--------------------|-------------------------|---|--------|
| NAO                    | Swiss Alps         | Wirth et al. [104]      | Increased floods in the southern Alps, reconstructed by<br>flood layers in lake deposits, were found to coincide<br>with annual and decadal NAO- phases during the<br>Little Ice Age. Enhanced flood occurrence in the South<br>compared to the North took place during a NAO-state.  | R      |
| PDO, AMO               | Multiple<br>gauges | Hodgkins et al.<br>[26] | No significant relationships, for any group of<br>catchments, between major flood occurrence and PDO.<br>Significant positive relations between floods and AMO<br>at medium-sized (100–1000 km <sup>2</sup> ) European<br>catchments.   | R      |
|                        | Pan-European       | Nobre et al.<br>[105]   | The link between ENSO and the occurrence of extreme<br>rainfall and intensity of extreme rainfall in Europe is<br>much weaker than the relationship with NAO or EA,<br>but still significant in some regions. Also flood damage<br>and flood occurrence have strong links with climate<br>variability, especially in Southern and Eastern Europe.                                 | Р      |
| ENSO, NAO              | River Danube       | Rimbu et al.<br>[106]   | The impact of NAO and ENSO on variability of the<br>Danube flow shows decadal variations. Winter SST<br>from tropical Pacific and some regions from the North<br>Atlantic are found to be significantly correlated with<br>the streamflow variations in spring and summer.  | R      |
|                        | River Rhine        | Ionita et al. [107]     | Variability of the Rhine flow, during the transition<br>seasons, was found to be strongly related to large-scale<br>atmospheric circulation patterns, being driven by<br>climate interactions at a nearly global scale in spring<br>and more regional in autumn.  | R      |
| AO                     | River Elbe         | Ionita et al. [108]     | High anomalies of the annual Elbe flow are associated<br>with a tripole-like pattern in the North Atlantic and<br>with negative SST anomalies in the Central-North<br>Pacific and positive anomalies in the Eastern and<br>Central tropical Pacific. The pattern identified in the<br>sea level pressure (SLP) resembles the negative phase<br>of the Arctic Oscillation pattern. | R      |

# Table 7. Cont.

**Table 8.** Review of links between climate oscillation and floods for Africa, as reported in the literature.Aspects of reference in the rightmost column—as in Table 2.

| Climate<br>Oscillation | Range  | Reference                   | Principal Message  | Aspect |
|------------------------|--------|-----------------------------|--|--------|
| ENSO                   | Africa | Nicholson, and<br>Kim [109] | ENSO is not a major factor in rainfall variability.<br>El Niño rather reduces rainfall over Africa than<br>enhances it. For Sahel, it may increase rainfall<br>early in the season but decrease it later in the<br>season, so may not influence on seasonal totals.<br>The signal over the rainfall is more intense,<br>consistent, spatially and temporally coherent<br>during the warm phase. Above-average rainfall<br>over most of Africa, particularly during<br>January–March and April–June may occur<br>during the first half of the ENSO cycle. For some<br>regions, rainfall may be enhanced during some<br>months of the rainy season but reduced during<br>others. For equatorial East Africa rainfall is<br>enhanced in the 'short rains' months<br>(October–December) of an El Niño year, but<br>reduced in the main rainy season a few<br>months later. | Р      |

| Climate<br>Oscillation | Range                                      | Reference                                | Principal Message   | Aspec     |
|------------------------|--|--|---|-----------|
|                        | Africa<br>Eastern Africa                   | Li et al. [110]<br>Siderius et al. [111] | Analysis of flood disaster events of 55 countries<br>in Africa from 1990 to 2014 recorded in the<br>International Disaster Database (EM-DAT) was<br>undertaken. Annual flood frequencies showed a<br>fluctuating upward trend and were in good<br>agreement with ENSO (SOI) years: at the 0.01<br>level in Africa as a whole, Southern Africa, and<br>West Africa, and at the 0.05 level in Central<br>Africa. The larger the SOI (La Niña), the greater<br>probability of the annual precipitation leading to<br>flood disasters.<br>El Niño events are generally associated with high<br>precipitation and floods in Eastern Africa. Kenya<br>experienced widespread flooding during the El | P, R<br>P |
|                        | South Africa                               | Quagraine et al.<br>[15]                 | Niño of 1997–1998.<br>Broadly wet conditions have occurred in the SRR<br>(Summer Rainfall Region) and parts of the ARR<br>(Annual Rainfall Region) when La Niña episodes<br>co-exist with a strongly negative AAO (Antarctic<br>Oscillation) during summer. A shift in AAO to<br>strongly positive drives wet conditions in Central<br>and Northern parts of the SRR although peak<br>conditions are centered to the east.  | Р         |
| ENSO                   | South Africa                               | Weldon and<br>Reason [112]               | Nearly all the anomalously wet years over the<br>South Coast of South Africa correspond to<br>mature phase of La Niña years.  | Р         |
|                        | South Africa                               | Alemaw and<br>Chaoka [113]               | Strong and nearly-strong positive linear<br>correlation between annual discharges (decline)<br>and the warm seasonal ENSO. The strong<br>positive correlation occurs in parts of Zambia,<br>Namibia, Mozambique, and in South Africa.   | R         |
|                        | Nile River<br>Basin                        | Siam and Eltahir<br>[114]                | There is a significant teleconnection between the<br>Nile flow and the SST over Eastern Pacific Ocean<br>(ENSO explains about 25% of the inter-annual<br>variability) and a region in the southern Indian<br>Ocean (50°–80° E and 25°–35° S, which explains<br>another 28% of the inter-annual variability in the<br>flow of the Nile). Cold conditions in both oceans<br>mean 83% probability of high flow in the Nile,<br>and insignificant probability of low flow.  | R         |
|                        | Zambia                                     | Brigadier et al.<br>[115]                | Most parts of Zambia experienced normal to<br>above-normal rainfall (floods) during La Niña<br>developing from autumn of the year 2010.   | Р         |
|                        | Coastal region<br>of Kenya and<br>Tanzania | Gamoyo et al. [116]                      | There is a strong relationship between ENSO (El<br>Niño), IOD and the October–December short<br>rains.  | Р         |
|                        | Zimbabwe                                   | Mamombe et al.<br>[117]                  | ENSO has a significant influence on Zimbabwe rainfall. La Niña years correspond to extreme wet summers in most parts of the country.  | Р         |
| NAO                    | Northern<br>Africa                         | Benito et al. [98]                       | Western Mediterranean regions (Southern<br>Europe and Northern Africa) show synchronicity<br>of flood episodes associated with NAO—that are<br>out-of-phase with those evident within the<br>eastern Mediterranean. Regional changes in<br>flood extremes are associated with switches in<br>predominant NAO phase, indicating their high<br>potential to model (as covariates) climate-related<br>impacts on flood risk.   | R         |

### Table 8. Cont.



Figure 1. Map of regional predisposition to abundance of water.

# 6. Concluding Remarks

No persuading, and spatially-coherent, trend could be detected in long time series of flood-related variables, such as high river discharge, flood severity and magnitude, as well as flood loss. The structure of time series of river discharge is dominated by strong, and seemingly rather irregular, inter-annual and inter-decadal variability. Perhaps part of this variability can be random or chaotic, but

there are serious arguments supporting the thesis that climate variability (quasi-periodic oscillations in the atmosphere–ocean system) plays, generally, an important role in driving variables related to floods over many areas of the world. Therefore, it is quite natural to examine whether there are links between the climate variability expressed via atmosphere–ocean oscillation indices and the variability of flood-related indices.

Spatially-defined oscillations in the ocean–atmosphere system have been found to affect the climate and various climate impacts over large areas, covering the oceans and adjacent continents, but also more remote (sometimes—very remote) regions, via so-called teleconnections. The strength of the link can vary from region to region and one can identify zones of influence where a climate variability mode (e.g., ENSO, PDO, NAO, or AMO) affects variables related to floods.

Many research works have been undertaken, with results reported in the literature, aimed at identification of links between the climate variability and various variables related to floods. Studying variability of properties of high river flows (e.g., timing, intensity, volume, duration, frequency) and their links with climate variability is a cutting-edge research activity, attracting broad attention due to scientific interests and practical importance. Anticipating regional predisposition to river floods in a particular phase of oscillation in the atmosphere–ocean system, measured by a range of climate variability indices, could help in improving the system of early preparedness to emergency.

The present paper reviewed published results of studies indicating links between climate variability indices and characteristics of abundant wetness (precipitation, river runoff, and flood indices) in different areas. Many studies have been found that report on the existence of such links, at a range of scales—from global (for ENSO and PDO) to local. Some references indicate the necessity to consider a combination of oscillation modes, e.g., one oscillation mode can modulate another one. Some studies combine analysis of change and variability, or seek for change points in system behavior. The list of references considered in this paper consists of 120 source items. The continents ordered after the number of source items, from highest to lowest, are: North America, Asia, Europe, Africa, South America, and Australia. Some reports on the existence of links are of a rather speculative nature, as a signal is sought among strong noise. Some findings are too complex and do not fit into the simple story of this paper.

As shown in Figure 1, the reviewed references found predisposition to abundance of water during a particular phase of one or more modes of climate variability. More consistent results are reported for South America and El Niño signals wetness there. In Africa, also ENSO is found important, while predisposition to abundance of water is related to La Niña for most areas and El Niño for some. In Europe, it is mostly NAO+ in the Northern and Western parts and NAO- in the Southern part. In Asia, links with ENSO (El Niño for some areas and La Niña for other areas) and PDO were reported. In Australia, it is La Niña and PDO-.

The situation in North and Central America looks most complex (Figure 1), because all modes: ENSO, NAO, AMO, and PDO may play a role. However, Jiang et al. [118] showed that no single index of those four, derived from the Pacific and Atlantic oceans, can explain the multi-scale temporal variability and spatial distribution of heavy precipitation in the western USA. Yet, characterization of their integrated effect, including modulation and reflecting the combined physical oceanic-atmospheric processes, allows useful links to be deciphered.

Reported links have spatial, temporal, and seasonal validity. As expected, different source items may give differing representation of links between climate variability and variables related to floods, for instance due to different definitions of indices, or due to different data windows (start-year and end-year) being considered. In brief, there are still considerable uncertainties in the understanding of the climate variability modes and their impact mechanisms on flood hazard and flood risk. This adds to the recognized general uncertainty of climate impacts on water resources [119].

Reviewed references follow various methodologies to investigate the relationship between climate oscillations and flood-related variables. Correlation and cross-correlation techniques (including time

lags) are commonly used. Many papers use principal component or empirical orthogonal function (EOF) analysis (e.g., [107,118]), as well as wavelet power spectrum analysis (e.g., [107]).

Some publications report on convincing physical explanations of mechanisms behind the relationships. One of examples is in the study for North America, by Gershunov and Barnett [120], who found that the North Pacific oscillation (NPO) modulates the consistency and strength of El Niño anomalies. They hypothesized that a deeper Aleutian low (manifestation of the NPO) shifts the storm track southward while El Niño provides an enhanced eastern tropical Pacific moisture source for the storms to tap. They also demonstrated that seasonal climate anomalies over North America exhibit rather large variability between years characterized by the same ENSO phase. Interdecadal climatic anomalies in the North Pacific (the North Pacific oscillation, NPO) were shown to exert a modulating effect on ENSO teleconnections. Typical El Niño patterns (e.g., low pressure over the northeastern Pacific, dry northwest, and wet southwest, etc.) are strong and consistent only during the high phase of the NPO, which is associated with an anomalously cold northwestern Pacific.

It is hoped that the present review paper could play an important and useful role, if only because of the systematization of scattered knowledge. It could improve understanding of the links of four climate variability modes with abundance of water and inform further studies in the area.

**Author Contributions:** The work was conceived by Z.W.K., who drafted the initial outline and preliminary and incomplete text. Z.W.K., M.S. and I.P. analyzed the rich pool of references and provided input to the tables and offered several rounds of amendments to the text. Z.W.K. provided final refinements to the manuscript.

**Funding:** All authors wish to acknowledge financial support from the project: "Interpretation of Change in Flood-Related Indices based on Climate Variability" (FloVar) funded by the National Science Centre of Poland (project number 2017/27/B/ST10/00924).

Conflicts of Interest: The authors declare no conflict of interest.

#### References

- 1. Trenberth, K.E.; Stepaniak, D.P. Indices of El Niño evolution. J. Clim. 2001, 14, 1697–1701. [CrossRef]
- Trenberth, K. (Ed.) The Climate Data Guide: Nino SST Indices (Nino 1+2, 3, 3.4, 4; ONI and TNI); National Center for Atmospheric Research Staff: Boulder, CO, USA, 2019. Available online: https://climatedataguide. ucar.edu/climate-data/nino-sst-indices-nino-12-3-34-4-oni-and-tni (accessed on 13 May 2019). [CrossRef]
- 3. Wanner, H.; Brönnimann, S.; Casty, C.; Gyalistras, D.; Luterbacher, J.; Schmutz, C.; Stephenson, D.B.; Xoplaki, E. North Atlantic Oscillation–concepts and studies. *Surv. Geophys.* **2001**, *22*, 321–382. [CrossRef]
- 4. Mantua, N.J.; Hare, S.R. The Pacific Decadal Oscillation. J. Oceanogr. 2002, 58, 35–44. [CrossRef]
- Fei-Fei, J. A theory of interdecadal climate variability of the North Pacific ocean–Atmosphere system. *J. Clim.* 1997, 10, 1821–1835.
- Dima, M.; Lohmann, G. A hemispheric mechanism for the Atlantic Multidecadal Oscillation. J. Clim. 2007, 20, 2706–2719. [CrossRef]
- Kundzewicz, Z.W.; Graczyk, D.; Maurer, T.; Pińskwar, I.; Radziejewski, M.; Svensson, C.; Szwed, M. Trend detection in river flow time series: 1. Annual maximum flow. *Hydrol. Sci. J.* 2005, *50*, 797–810. [CrossRef]
- Kundzewicz, Z.W.; Kanae, S.; Seneviratne, S.I.; Handmer, J.; Nicholls, N.; Peduzzi, P.; Mechler, R.; Bouwer, L.M.; Arnell, N.; Mach, K.; et al. Flood risk and climate change: Global and regional perspectives. *Hydrol. Sci. J.* 2014, *59*, 1–28. [CrossRef]
- Prudhomme, C.; Genevier, M. Can atmospheric circulation be linked to flooding in Europe? *Hydrol. Proc.* 2011, 25, 1180–1190. [CrossRef]
- 10. Burn, D.H.; Arnell, N.W. Synchronicity in global flood responses. J. Hydrol. 1993, 144, 381–404. [CrossRef]
- 11. Kundzewicz, Z.W.; Pińskwar, I.; Brakenridge, G.R. Large floods in Europe, 1985–2009. *Hydrol. Sci. J.* **2013**, *58*, 1–7. [CrossRef]
- 12. Kundzewicz, Z.W.; Pińskwar, I.; Brakenridge, G.R. Changes in river flood hazard in Europe—A review. *Hydrol. Res.* **2018**, *49*, 294–302. [CrossRef]
- Hartmann, H.; Snow, J.A.; Stein, S.; Su, B.; Zhai, J.; Jiang, T.; Krysanova, V.; Kundzewicz, Z.W. Predictors of precipitation for improved water resources management in the Tarim river basin: Creating a seasonal forecast model. *J. Arid Environ.* 2016, 125, 31–42. [CrossRef]

- 14. Lee, J.H.; Ramirez, J.A.; Kim, T.W.; Julien, P.Y. Variability, teleconnection, and predictability of Korean precipitation in relation to large scale climate indices. *J. Hydrol.* **2019**, *568*, 12–25. [CrossRef]
- 15. Quagraine, K.A.; Hewitson, B.; Jack, C.; Pinto, I.; Lennard, C. A methodological approach to assess the co-behavior of climate processes over Southern Africa. *J. Clim.* **2019**, *32*, 2483–2495. [CrossRef]
- 16. Wang, L.; Chen, W.; Huang, R. Interdecadal modulation of PDO on the impact of ENSO on the east Asian winter monsoon. *Geophys. Res. Lett.* **2008**, *35*, L20702. [CrossRef]
- 17. Chan, J.C.L.; Zhou, W. PDO, ENSO and the early summer monsoon rainfall over south China. *Geophys. Res. Lett.* **2005**, *32*, L08810. [CrossRef]
- 18. Chiew, F.H.S.; McMahon, T.A. Global ENSO-streamflow teleconnection, streamflow forecasting and interannual variability. *Hydrol. Sci. J.* **2002**, *47*, 505–522. [CrossRef]
- 19. Pociask-Karteczka, J. River hydrology and the North Atlantic Oscillation: A general review. *Ambio* **2006**, *35*, 312–314. [CrossRef]
- 20. Wrzesiński, D.; Paluszkiewicz, R. Spatial differences in the impact of the North Atlantic Oscillation on the flow of rivers in Europe. *Hydrol. Res.* **2011**, *42*, 30–39. [CrossRef]
- 21. Ruiz-Villanueva, V.; Stoffel, M.; Wyżga, B.; Kundzewicz, Z.W.; Czajka, B.; Niedźwiedź, T. Decadal variability of floods in the northern foreland of the Tatra Mountains. *Reg. Environ. Chang.* **2016**, *16*, 603–615. [CrossRef]
- 22. Sun, X.; Renard, B.; Thyer, M.; Westra, S.; Lang, M. A global analysis of the asymmetric effect of ENSO on extreme precipitation. *J. Hydrol.* **2015**, *530*, 51–65. [CrossRef]
- 23. Ward, P.J.; Eisner, S.; Flörke, M.; Dettinger, M.D.; Kummu, M. Annual flood sensitivities to El Niño–Southern Oscillation at the global scale. *Hydrol. Earth Sys. Sci.* **2014**, *18*, 47–66. [CrossRef]
- 24. Ward, P.J.; Jongman, B.; Kummu, M.; Dettinger, M.D.; Sperna Weiland, F.C.; Winsemius, H.C. Strong influence of El Niño Southern Oscillation on flood risk around the world. *PNAS* **2014**, *111*, 15659–15664. [CrossRef] [PubMed]
- 25. Emerton, R.; Cloke, H.L.; Stephens, E.M.; Zsoter, E.; Woolnough, S.J.; Pappenberger, F. Complex picture for likelihood of ENSO-driven flood hazard. *Nat. Comm.* **2017**, *8*, 14796. [CrossRef] [PubMed]
- Hodgkins, G.A.; Whitfield, P.H.; Burn, D.H.; Hannaford, J.; Renard, B.; Stahl, K.; Fleig, A.K.; Madsen, H.; Mediero, L.; Korhonen, J.; et al. Climate-driven variability in the occurrence of major floods across North America and Europe. J. Hydrol. 2017, 552, 704–717. [CrossRef]
- 27. Räsänen, T.A.; Kummu, M. Spatiotemporal influences of ENSO on precipitation and flood pulse in the Mekong River Basin. *J. Hydrol.* **2013**, 476, 154–168. [CrossRef]
- 28. Ward, P.J.; Kummu, M.; Lall, U. Flood frequencies and durations and their response to El Niño Southern Oscillation: Global analysis. *J. Hydrol.* **2016**, *539*, 358–378. [CrossRef]
- 29. Najibi, N.; Devineni, N. Recent trends in the frequency and duration of global floods. *Earth Sys. Dyn.* **2018**, *9*, 757–783. [CrossRef]
- 30. Dong, B.; Dai, A. The influence of the Interdecadal Pacific Oscillation on temperature and precipitation over the globe. *Clim. Dyn.* **2015**, 45, 2667. [CrossRef]
- 31. Lau, K.M.; Weng, H.Y. Recurrent teleconnection patterns linking summertime precipitation variability over East Asia and North America. *J. Meteorol. Soc. Jpn.* **2002**, *80*, 1309–1324. [CrossRef]
- 32. Kripalani, R.H.; Kulkarni, A. Climatic impact of El Niño/La Niña on the Indian monsoon: A new perspective. *Weather* **1997**, *52*, 39–46. [CrossRef]
- 33. Lv, A.F.; Qu, B.; Jia, S.F.; Zhu, W.B. Influence of three phases of El Nino-Southern Oscillation on daily precipitation regimes in China. *Hydrol. Earth Sys. Sci.* **2019**, *23*, 883–896. [CrossRef]
- Cheng, Q.P.; Gao, L.; Zuo, X.A.; Zhong, F.L. Statistical analyses of spatial and temporal variabilities in total, daytime, and nighttime precipitation indices and of extreme dry/wet association with large-scale circulations of Southwest China, 1961–2016. *Atmos. Res.* 2019, 219, 166–182. [CrossRef]
- 35. Wang, S.; Yuan, X. Extending seasonal predictability of Yangtze River summer floods. *Hydrol. Earth Sys. Sci. Discuss.* **2018**, *22*, 4201–4211. [CrossRef]
- 36. Zhang, Q.; Xu, C.; Jiang, T.; Wu, Y. Possible influence of ENSO on annual maximum streamflow of the Yangtze River, China. *J. Hydrol.* **2007**, *333*, 265–274. [CrossRef]
- 37. Tong, J.; Qiang, Z.; Deming, Z.; Yijin, W. Yangtze floods and droughts (China) and teleconnections with ENSO activities (1470–2003). *Quat. Int.* **2006**, *144*, 29–37. [CrossRef]
- 38. Ye, X.C.; Wu, Z.W. Contrasting impacts of ENSO on the interannual variations of summer runoff between the upper and mid-lower reaches of the Yangtze River. *Atmosphere* **2018**, *9*, 478. [CrossRef]

- 39. Iqbal, A.; Hassan, S.A. ENSO and IOD analysis on the occurrence of floods in Pakistan. *Nat. Haz.* **2018**, *91*, 879–890. [CrossRef]
- 40. Higashino, M.; Stefan, H.G. Variability and change of precipitation and flood discharge in a Japanese river basin. *J. Hydrol. Reg. Stud.* **2019**, *21*, 68–79. [CrossRef]
- 41. Chowdhury, M.R. The El Niño-Southern Oscillation (ENSO) and seasonal flooding–Bangladesh. *Theor. Appl. Climatol.* **2003**, *76*, 105–124. [CrossRef]
- 42. Zhang, Q.; Wanga, Y.; Singh, V.P.; Gu, X.; Kong, D.; Xiao, M. Impacts of ENSO and ENSO Modoki + A regimes on seasonal precipitation variations and possible underlying causes in the Huai River basin, China. *J. Hydrol.* **2016**, *533*, 308–319. [CrossRef]
- 43. Liu, J.; Zhang, Q.; Singh, V.P.; Gu, X.; Shi, P. Nonstationarity and clustering of flood characteristics and relations with the climate indices in the Poyang Lake basin, China. *Hydrol. Sci. J.* **2017**, *11.* [CrossRef]
- 44. Cao, Q.; Hao, Z.; Yuan, F.; Su, Z.; Berndtsson, R.; Hao, J.; Nyima, T. Impact of ENSO regimes on developingand decaying-phase precipitation during rainy season in China. *Hydrol. Earth Sys. Sci.* **2017**, *21*, 5415–5426. [CrossRef]
- 45. Saghafian, B.; Haghnegahdar, A.; Dehghani, M. Effect of ENSO on annual maximum floods and volume over threshold in the southwestern region of Iran. *Hydrol. Sci. J.* **2017**, *62*, 1039–1049. [CrossRef]
- Delgado, J.M.; Merz, B.; Apel, H. A climate-flood link for the lower Mekong River. *Hydrol. Earth Sys. Sci.* 2012, *16*, 1533–1541. [CrossRef]
- 47. Zhang, Q.; Gu, X.; Singh, V.P.; Xiao, M.; Chen, X. Evaluation of flood frequency under non-stationarity resulting from climate indices and reservoir indices in the East River basin, China. *J. Hydrol.* **2015**, 527, 565–575. [CrossRef]
- Deng, Y.; Jiang, W.; He, B.; Chen, Z.; Jia, K. Change in intensity and frequency of extreme precipitation and its possible teleconnection with large-scale climate index over the China from 1960 to 2015. *J. Geophys. Res. Atmos.* 2018, 123, 2068–2081. [CrossRef]
- 49. Xiao, M.; Zhang, Q.; Singh, V.P. Influences of ENSO, NAO, IOD and PDO on seasonal precipitation regimes in the Yangtze River basin, China. *Int. J. Climatol.* **2015**, *35*, 3556–3567. [CrossRef]
- 50. Cullen, H.M.; Kaplan, A.; Arkin, P.A.; Demenocal, P.B. Impact of the North Atlantic Oscillation on Middle Eastern climate and streamflow. *Clim. Chang.* **2002**, *55*, 315–338. [CrossRef]
- 51. Cayan, D.R.; Redmond, K.T.; Riddle, L.G. ENSO and hydrologic extremes in the western United States. *J. Clim.* **1999**, *12*, 2881–2893. [CrossRef]
- 52. Corringham, T.W.; Cayan, D.R. The effect of El Niño on flood damages in the Western United States. *Weather Clim. Soc.* 2019, 11, 489–504. [CrossRef]
- 53. Wang, S.-Y.S.; Huang, W.-R.; Hsu, H.-H.; Gillies, R.R. Role of the strengthened El Niño teleconnection in the May 2015 floods over the southern Great Plains. *Geophys. Res. Lett.* **2015**, *42*, 8140–8146. [CrossRef]
- 54. Munoz, S.E.; Dee, S.G. El Nino increases the risk of lower Mississippi River flooding. *Sci. Rep.* **2017**, *7*, 1772. [CrossRef] [PubMed]
- 55. Nakamura, J.; Lall, U.; Kushnir, Y.; Robertson, A.W.; Seager, R. Dynamical structure of extreme floods in the U.S. Midwest and the United Kingdom. *J. Hydrometeorol.* **2013**, *14*, 485–504. [CrossRef]
- 56. Andrews, E.D.; Antweiler, R.C.; Neiman, P.J.; Ralph, F.M. Influence of ENSO on flood frequency along the California Coast. *J. Clim.* **2004**, *17*, 337–348. [CrossRef]
- 57. Whan, K.; Zwiers, F. The impact of ENSO and the NAO on extreme winter precipitation in North America in observations and regional climate models. *Clim. Dyn.* **2017**, *48*, 1401–1411. [CrossRef]
- Wang, S.-Y.; Hakala, K.; Gillies, R.R.; Capehart, W.J. The Pacific Quasi-decadal Oscillation (QDO): An important precursor toward anticipating major flood events in the Missouri River Basin? *Geophys. Res. Lett.* 2014, 41, 991–997. [CrossRef]
- 59. Nasser, N.; Devineni, N.; Lu, M. Hydroclimate drivers and atmospheric teleconnections of long duration floods: An application to large reservoirs in the Missouri River Basin. *Adv. Water Resour.* **2017**, *100*, 153–167.
- 60. Villarini, G.; Goska, R.; Smith, J.A.; Vecchi, G.A. North Atlantic tropical cyclones and U.S. flooding. *Bull. Am. Meteorol. Soc. (BAMS)* **2014**, *95*, 1381–1388. [CrossRef]
- 61. Dai, A.G. The influence of the inter-decadal Pacific Oscillation on US precipitation during 1923-2010. *Clim. Dyn.* **2013**, *41*, 633–646. [CrossRef]
- 62. Gurrapu, S.; St. Jacques, J.-M.; Sauchyn, D.J.; Hodder, K.R. The influence of the Pacific Decadal Oscillation on annual floods in the rivers of western Canada. *J. Am. Water Resour. Assoc.* **2016**, *52*, 1031–1045. [CrossRef]

- 63. Burn, D.H.; Cunderlik, J.M.; Pietroniro, A. Hydrological trends and variability in the Liard River basin. *Hydrol. Sci. J.* **2004**, *49*, 53–67. [CrossRef]
- 64. Beebee, R.A.; Manga, M. Variation in the relationship between snowmelt runoff in Oregon and ENSO and PDO. *J. Am. Water Resour. Assoc.* **2004**, *40*, 1011–1024. [CrossRef]
- 65. Whited, D.C.; Lorang, M.S.; Harner, M.J.; Hauer, F.R.; Kimball, J.S.; Stanford, J.A. Climate, hydrologic disturbance, and succession: Drivers of floodplain pattern. *Ecology* **2007**, *88*, 940–953. [CrossRef] [PubMed]
- 66. Enfield, D.B.; Mestas-Nunez, A.M.; Trimble, P.J. The Atlantic Multidecadal Oscillation and its relation to rainfall and river flows in the continental U.S. *Geophys. Res. Lett.* **2001**, *28*, 2077–2080. [CrossRef]
- 67. Fortier, C.; Assani, A.; Mesfioui, M.; Roy, A.G. Comparison of the interannual and interdecadal variability of heavy flood characteristics upstream and downstream from dams in inversed hydrologic regime: Case study of Matawin River (Québec, Canada). *River Res. Appl.* **2011**, *27*, 1277–1289. [CrossRef]
- 68. Valdes-Manzanilla, A. Historical floods in Tabasco and Chiapas during sixteenth–twentieth centuries. *Nat. Hazards* **2016**, *80*, 1563–1577. [CrossRef]
- Munoz, S.E.; Giosan, L.; Therrell, M.D.; Remo, J.W.F.; Shen, Z.; Sullivan, R.M.; Wiman, C.; O'Donnell, M.; Donnelly, J.P. Climatic control of Mississippi River flood hazard amplified by river engineering. *Nature* 2018, 556, 95–98. [CrossRef]
- 70. Assani, A.A.; Charron, S.; Matteau, M.; Mesfioui, M.; Quessy, J.-F. Temporal variability modes of floods for catchments in the St. Lawrence watershed (Quebec, Canada). *J. Hydrol.* **2010**, *385*, 292–299. [CrossRef]
- 71. Hamlet, A.F.; Lettenmaier, D.P. Effects of 20th century warming and climate variability on flood risk in the western U.S. *Water Resour. Res.* 2007, *43*, W06427. [CrossRef]
- 72. McCabe, G.J., Jr.; Dettinger, M.D. Decadal variations in the strength of ENSO teleconnections with precipitation in the western United States. *Int. J. Climatol.* **1999**, *19*, 1399–1410. [CrossRef]
- 73. Gobena, A.K.; Gan, T.Y. Low-frequency variability in southwestern Canadian streamflow: Links to large-scale climate anomalies. *Int. J. Climatol.* **2006**, *26*, 1843–1869. [CrossRef]
- 74. Jain, S.; Lall, U. Floods in a changing climate: Does the past represent the future? *Water Resour. Res.* 2001, *37*, 3193–3205. [CrossRef]
- 75. Jain, S.; Lall, U. Magnitude and timing of annual maximum floods: Trends and large-scale climatic associations for the Blacksmith Fork River, Utah. *Water Resour. Res.* **2000**, *36*, 3641–3651. [CrossRef]
- Méndez, M.; Magaña, V. Regional aspects of prolonged meteorological droughts over Mexico and Central America. J. Clim. 2010, 23, 1175–1188. [CrossRef]
- Pascolini-Campbella, M.; Seagerb, R.; Pinsonc, A.; Cook, B.I. Covariability of climate and streamflow in the Upper Rio Grande from interannual to interdecadal timescales. *J. Hydrol. Reg. Stud.* 2017, 13, 58–71. [CrossRef]
- Valdes-Manzanilla, A. Effect of climatic oscillations on flood occurrence on Papaloapan river, México, during the 1550–2000 period. *Nat. Haz.* 2016, 94, 167–180. [CrossRef]
- Mazouz, R.; Assani, A.A.; Quessy, J.-F.; Légaré, G. Comparison of the interannual variability of spring heavy floods characteristics of tributaries of the St. Lawrence river in Quebec (Canada). *Adv. Water Resour.* 2012, 35, 110–120. [CrossRef]
- 80. Archfield, S.A.; Hirsch, R.M.; Viglione, A.; Blöschl, G. Fragmented patterns of flood change across the United States. *Geophys. Res. Lett.* **2016**, *43*. [CrossRef]
- 81. Mallakpour, I.; Villarini, G. Investigating the relationship between the frequency of flooding over the central United States and large-scale climate. *Adv. Water Resour.* **2016**, *92*, 159–171. [CrossRef]
- 82. Isla, F.I. ENSO-triggered floods in South America: Correlation between maximum monthly discharges during strong events. *Hydrol. Earth Syst. Sci. Discuss.* **2018**, 1–13. [CrossRef]
- Isla, F.I.; Toldo Junior, E.E. ENSO impacts on Atlantic watersheds of South America. *Quat. Environ. Geosci.* 2013, 4, 34–41. [CrossRef]
- Guevara-Murua, A.; Williams, C.A.; Hendy, E.J.; Imbach, P. 300 years of hydrological records and societal responses to droughts and floods on the Pacific coast of Central America. *Clim. Past* 2018, 14, 175–191. [CrossRef]
- 85. Depetris, P.J. The Parana River under extreme flooding: A hydrological and hydro-geochemical insight. *Interciencia* **2007**, *32*, 656–662.
- 86. Prieto, M.R. ENSO signals in South America: Rains and floods in the Paraná River region during colonial times. *Clim. Chang.* **2007**, *83*, 39–54. [CrossRef]

- 87. Arteaga, K.; Tutasi, P.; Jiménez, R. Climatic variability related to El Niño in Ecuador–a historical background. *Adv. Geosci.* **2006**, *6*, 237–241. [CrossRef]
- 88. Houston, J. Variability of precipitation in the Atacama Desert: Its causes and hydrological impact. *Int. J. Climatol.* **2006**, *26*, 2181–2198. [CrossRef]
- Borges, P.D.; Bernhofer, C.; Rodrigues, R. Extreme rainfall indices in Distrito Federal, Brazil: Trends and links with El Nino Southern Oscillation and Madden-Julian Oscillation. *Int. J. Climatol.* 2018, *38*, 4550–4567. [CrossRef]
- 90. Micevski, T.; Franks, S.W.; Kuczera, G. Multidecadal variability in coastal eastern Australian flood data. *J. Hydrol.* **2006**, 327, 219–225. [CrossRef]
- 91. Franks, S.W. Identification of a change in climate state using regional flood data. *Hydrol. Earth Sys. Sci.* 2002, *6*, 11–16. [CrossRef]
- 92. Franks, S.W.; Kuczera, G. Flood frequency analysis: Evidence and implications of secular climate variability, New South Wales. *Water Resour. Res.* **2002**, *38*, 1062. [CrossRef]
- 93. Verdon, D.C.; Wyatt, A.M.; Kiem, A.S.; Franks, S.W. Multidecadal variability of rainfall and streamflow: Eastern Australia. *Water Resour. Res.* **2004**, *40*, W10201. [CrossRef]
- 94. Cai, W.; van Rensch, P. The 2011 southeast Queensland extreme summer rainfall: A confirmation of a negative Pacific Decadal Oscillation phase? *Geophys. Res. Lett.* **2012**, *39*, L08702. [CrossRef]
- 95. Kiem, A.S.; Franks, S.W.; Kuczera, G. Multi-decadal variability of flood risk. *Geophys. Res. Lett.* 2003, 30, 1035. [CrossRef]
- 96. Zanardo, S.; Nicotina, L.; Hilberts, A.G.J.; Jewson, S.P. Modulation of economic losses from European floods by the North Atlantic Oscillation. *Geophys. Res. Lett.* **2019**, *46*, 2563–2572. [CrossRef]
- Benito, G.; Macklin, M.G.; Panin, A.; Rossato, S.; Fontana, A.; Jones, A.F.; Machado, M.J.; Matlakhova, E.; Mozzi, P.; Zielhofer, C. Recurring flood distribution patterns related to short-term Holocene climatic variability. *Sci. Rep.* 2015, *5*, 16398. [CrossRef] [PubMed]
- 98. Benito, G.; Macklin, M.G.; Zielhofer, C.; Jones, A.F.; Machado, M.J. Holocene flooding and climate change in the Mediterranean. *Catena* **2015**, *130*, 13–33. [CrossRef]
- Moreno, A.; Valero-Garcés, B.L.; González-Sampériz, P.; Rico, M. Flood response to rainfall variability during the last 2000 years inferred from the Taravilla Lake record (Central Iberian Range, Spain). *J. Paleolimnol.* 2008, 40, 943–961. [CrossRef]
- 100. Huntingford, C.; Marsh, T.; Scaife, A.A.; Kendon, E.J.; Hannaford, J.; Kay, A.L.; Lockwood, M.; Prudhomme, C.; Reynard, N.S.; Parry, S.; et al. Potential influences on the United Kingdom's floods of winter 2013/14. *Nat. Clim. Chang.* 2014, 4, 769–777. [CrossRef]
- 101. Hannaford, J.; Marsh, T.J. High-flow and flood trends in a network of undisturbed catchments in the UK. *Int. J. Climatol.* **2008**, *28*, 1325–1338. [CrossRef]
- Toonen, W.H.J.; Middelkoop, H.; Konijnendijk, T.Y.M.; Macklin, M.G.; Cohen, K.M. The influence of hydroclimatic variability on flood frequency in the Lower Rhine. *Earth Surf. Processes Landf.* 2016, 41, 1266–1275. [CrossRef]
- Schulte, L.; Pena, J.C.; Carvalho, F.; Schmidt, T.; Julia, R.; Llorca, J.; Veit, H. A 2600 year history of floods in the Bernese Alps, Switzerland: Frequencies, mechanisms and climate forcing. *Hydrol. Earth Sys. Sci. Discuss.* 2015, 12, 3391–3448. [CrossRef]
- 104. Wirth, S.B.; Glur, L.; Gilli, A.; Anselmetti, F.S. Holocene flood frequency across the Central Alps–solar forcing and evidence for variations in North Atlantic atmospheric circulation. *Quat. Sci. Rev.* 2013, 80, 112–128. [CrossRef]
- 105. Nobre, G.G.; Jongman, B.; Aerts, J.; Ward, P.J. The role of climate variability in extreme floods in Europe. *Environ. Res. Lett.* **2017**, *12*, 084012. [CrossRef]
- 106. Rimbu, N.; Dima, M.; Lohmann, G.; Stefan, S. Impacts of the North Atlantic Oscillation and the El Niño–Southern Oscillation on Danube river flow variability. *Geophys. Res. Lett.* **2004**, *31*, 1035. [CrossRef]
- 107. Ionita, M.; Lohmann, G.; Rimbu, N.; Chelcea, S. Interannual variability of Rhine River streamflow and its relationship with large-scale anomaly patterns in spring and autumn. *J. Hydrometeorol.* 2012, 13, 172–188. [CrossRef]
- 108. Ionita, M.; Rimbu, N.; Lohmann, G. Decadal variability of the Elbe River streamflow. *Int. J. Climatol.* **2011**, 31, 22–30. [CrossRef]

- 109. Nicholson, S.E.; Kim, J. The relationship of the El Niño–Southern Oscillation to African rainfall. *Int. J. Climatol.* **1997**, *17*, 117–135. [CrossRef]
- 110. Li, C.; Chai, Y.; Yang, L.; Li, H. Spatio-temporal distribution of flood disasters and analysis of influencing factors in Africa. *Nat. Haz.* **2016**, *82*, 721–731. [CrossRef]
- 111. Siderius, C.; Gannon, K.E.; Ndiyoi, M.; Opere, A.; Batisani, N.; Olago, D.; Pardoe, J.; Conway, D. Hydrological response and complex impact pathways of the 2015/2016 El Niño in Eastern and Southern Africa. *Earth's Futur.* 2018, 6, 2–22. [CrossRef]
- 112. Weldon, D.; Reason, C.J.C. Variability of rainfall characteristics over the South Coast region of South Africa. *Theor. Appl. Climatol.* **2014**, *115*, 177–185. [CrossRef]
- 113. Alemaw, B.F.; Chaoka, T.R. The 1950-1998 warm ENSO events and regional implications to river flow variability in Southern Africa. *Water* **2006**, *32*, 459–464.
- 114. Siam, M.S.; Eltahir, E.A.B. Explaining and forecasting interannual variability in the flow of the Nile River. *Hydrol. Earth Syst. Sci.* 2015, 19, 1181–1192. Available online: www.hydrol-earth-syst-sci.net/19/1181/2015 (accessed on 13 May 2019). [CrossRef]
- 115. Brigadier, L.; Ogwang, B.A.; Ongoma, V.; Ngonga, C.; Nyasa, L. Diagnosis of the 2010 DJF flood over Zambia. *Nat. Hazards* **2016**, *81*, 189–201. [CrossRef]
- Gamoyo, M.; Reason, C.; Obura, D. Rainfall variability over the East African coast. *Theor. Appl. Climatol.* 2015, 120, 311–322. [CrossRef]
- 117. Mamombe, V.; Kim, W.M.; Choi, Y.S. Rainfall variability over Zimbabwe and its relation to large-scale atmosphere–ocean processes. *Int. J. Climatol.* **2017**, *37*, 963–971. [CrossRef]
- 118. Jiang, P.; Yu, Z.; Gautam, M.R. Pacific and Atlantic Ocean influence on the spatiotemporal variability of heavy precipitation in the western United States. *Glob. Planet. Change* **2013**, *109*, 38–45.
- 119. Kundzewicz, Z.W.; Krysanova, V.; Benestad, R.E.; Hov, Ø.; Piniewski, M.; Otto, I.M. Uncertainty in climate change impacts on water resources. *Environ. Sci. Pol.* **2018**, *79*, 1–8. [CrossRef]
- 120. Gershunov, A.; Barnett, T.P. Interdecadal modulation of ENSO teleconnections. *Bull. Am. Meteorol. Soc.* **1998**, 79, 2715–2726. [CrossRef]



© 2019 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (http://creativecommons.org/licenses/by/4.0/).