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# **Geochemical Characteristics of Late Ordovician Shales in the Upper Yangtze Platform, South China: Implications for Redox Environmental Evolution**

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**Abstract:** Changes to the redox environment of seawater in the Late Ordovician affect the process of organic matter enrichment and biological evolution. However, the evolution of redox and its underlying causes remain unclear. This paper analyzed the vertical variability of main, trace elements and  $\delta^{34}$ Spy from a drill core section (well ZY5) in the Upper Yangtze Platform, and described the redox conditions, paleoproductivity and paleoclimate variability recorded in shale deposits of the *P. pacificus* zone and *M. extraordinarius* zone that accumulated during Wufeng Formation. The results showed that shale from well ZY5 in Late Ordovician was deposited under oxidized water environment, and there are more strongly reducing bottom water conditions of the *M. extraordinarius* zone compared with the *P. pacificus* zone. Excess silica ( $SiO_{2(exc)}$ ) and substitution index of paleoproductivity (Y) indicated that the *P. pacificus* zone had higher paleoproductivity whereas the *M. extraordinarius* zone was lower. The high productivity level controlled  $O<sub>2</sub>$  release in the shallow water area as well as the oxidation degree of the *P. pacificus* zone. The decrease of productivity and the relatively stagnant water mass of the inner Yangtze Sea controlled the formation of relatively reduced water conditions in the *M. extraordinarius* zone. The chemical index of alteration (CIA) results suggested that palaeoclimatic conditions changed from warm and humid to cold and dry climate from the *P. pacificus* to the *M. extraordinarius* zones in the study area. A comparative analysis of the published Fe-S-C data for the Xiushan Datianba section showed that in the *P. pacificus* zone of the inner Yangtze Sea, warm and humid climate conditions drove high productivity, sulphate flux and low reactive iron flux, which promoted the expansion of oxic ocean-surface waters and mid-depth euxinic waters. In the *M. extraordinarius* zone, the cold and dry climate with significant uplift of the Xiang'e Submarine High led to the relative sea level decline, resulting in low productivity, sulfate flux and high reactive iron flux, which promoted the expansion of the mid-depth ferruginous waters and the shrinkage of oxic ocean-surface waters. The results offered new insights into the co-evolution of continents and oceans, and explained the role of continental weathering and uplift of the Xiang'e Submarine High in the exchange of sulfate flux and nutrients in the redox environment change of inner Yangtze Sea during the Late Ordovician.

**Keywords:** inner Yangtze sea; Late Ordovician; paleoclimatic; paleoredox; paleoproductivity

## **1. Introduction**

The Late Ordovician is an important transition period in geological history, accompanied by large-scale glaciation, volcanism, sea level change, local tectonic activities and other geological and biological extinction events [\[1](#page-14-0)[–3\]](#page-14-1). The Wufeng-Longmaxi Formation is widely distributed in the Yangtze platform, and is an important stratigraphic record

of geological events and an exceptional hydrocarbon source for commercial oil and gas exploration and production in South China [\[4\]](#page-14-2). The change of redox environment in the Late Ordovician had a significant effect on the accumulation of organic matter and the biotic extinction event. Therefore, the evolution of the organic-rich shale sedimentary environment during the Ordovician-Silurian transition has become a focus of recent research efforts [\[5](#page-14-3)[–7\]](#page-15-0).

Recent studies have suggested that the Ordovician-Silurian oceans may have been made up of mid-depth euxinic waters dynamically sandwiched between surface oxic and deep ferruginous waters [\[3,](#page-14-1)[8](#page-15-1)[–10\]](#page-15-2). The spatial heterogeneity of the redox environment in the Yangtze Sea during the Late Ordovician is largely related to the global climate change (e.g., the Hirnantian Glaciation) and tectonic activity (e.g., the Kwangsian Orogeny) [\[5,](#page-14-3)[8\]](#page-15-1). It has been confirmed that the expansion of mid-depth marine euxinia was related to the uneven spatial distribution of sulfate contents in the inner Yangtze sea during the late Ordovician [\[10\]](#page-15-2). However, the spatiotemporal evolution mechanism of the redox environment is still unclear, and many important factors including continental weathering have not been taken into consideration. The change in oxic ocean-surface and mid-depth euxinic waters is believed to have influenced the change of palaeoproductivity and sulfate flux  $[1,11,12]$  $[1,11,12]$  $[1,11,12]$  in response to both tectonic and palaeoclimate changes.

To analyze the redox environment evolution and its controlling factors from the inner Yangtze Sea in the Late Ordovician, this study will focus on Wufeng Formation shale from Well ZY5 deposited in the southwest margin of the Xiang'e Submarine High as the research object. The Si, Cu, Ni and Zn proxies are used for oceanic primary productivity evaluation [\[13\]](#page-15-5). A combination of independent geochemical proxies (i.e., redox-sensitive trace elements Mo and U, pyrite  $\delta^{34}S$  compositions ( $\delta^{34}S_{\text{py}}$ )) and pyrite framboid size distribution data are used to reconstruct the redox conditions of the bottom water [\[14](#page-15-6)[–16\]](#page-15-7). Then, we made a comprehensive analysis of the paleoclimate conditions. By comparing Fe-S-C results with published data from the Datianba section [\[8\]](#page-15-1), we aimed to provide a new insight of the heterogeneity of climate change and tectonic activity on the redox environment evolution in the inner Yangtze Sea during the Late Ordovician.

## **2. Geologic Setting**

The South China Plate was an isolated microcontinent attached to the northwestern margin of Gondwana in Late Ordovician (Figure [1a](#page-2-0)) [\[17\]](#page-15-8). The Yangtze Platform was located in the middle of the South China Plate and was covered by a vast continental margin sea. Due to the influence of the Kwangsian Orogeny, the Cathaysia plate collided with the Yangtze plate from southeast to northwest, resulting in the uplift of the Qianzhong Oldland and Xiang'e Submarine High [\[18\]](#page-15-9). The Yangtze Sea was divided into inner Yangtze Sea and outer Yangtze Sea [\[19\]](#page-15-10), forming a paleogeographic pattern with "one uplift and three depressions" surrounded by oldlands on three sides (Figure [1b](#page-2-0)). Due to the relatively weak connectivity between the inner and outer seas, the inner Yangtze Sea was semi-restricted with anoxic and low-energy bottom waters [\[1,](#page-14-0)[8,](#page-15-1)[10\]](#page-15-2).

Well ZY5 and the Xiushan Datianba section are located in the inner Yangtze Sea. Datianba was near the northeastern margin of the Qianzhong oldland, and well ZY5 was located in the southwest margin of the Xiang'e Submarine High (Figure [1b](#page-2-0)). Two sections contain the Upper Ordovician Wufeng Formation, which are dominantly black shale containing a large number of spongy spicules and radiolarian fossils [\[20\]](#page-15-11). The Wufeng Formation of the Datianba section was continuously deposited, whereas well ZY5 near the Xiang'e Submarine High was missing the late Hirnantian stage strata. Under the influence of Kwangsian Orogeny, the Cathaysia oldland gradually advanced to the northwest during the late Ordovician, which resulted in the intensification of the uplift of the Xiang'e Submarine High. The Submarine High and surrounding areas were subjected to different degrees of denudation [\[21\]](#page-15-12).

Chen et al. [\[22\]](#page-15-13) used graptolites fossils from Late Ordovician-Early Silurian transition profiles in South China and their stratigraphic correlation to establish regional and global

biological zoning programmes. The *Dicellograptus Complexus* graptolite zone, *Paraorthograptus pacificus* graptolite zone and *Metabolograptus extraordinarius* graptolite zone developed *thograptus pacificus* graptolite zone and *Metabolograptus extraordinarius* graptolite zone dein the Wufeng Formation of the Late Ordovician. Well ZY5 is based on graptolite character-istics and the GR curve recognition method proposed by Xi et al. [\[23\]](#page-15-14) can identify the range from 2016 m to 2010 m as the *Paraorthograptus pacificus* graptolite zone (hereafter referred to as the *P. pacificus* zone). From 2010 m to 2096 m is the *Metabolograptus extraordinarius*<br>constable zone (the *M. minordinarius* zone). (Figure 1.) graptolite zone (the *M. extraordinarius* zone) (Figure [1c](#page-2-0)). *dinarius* graptolite zone (the *M. extraordinarius* zone) (Figure 1c). profiles in South China and their stratigraphic correlation to establish regional and global biological zoning programmes. The *Dicellograptus Complexus* graptolite zone, *Paraor-*

<span id="page-2-0"></span>

Figure 1. (a) Late Ordovician global paleogeography in South China (Modified after [\[2\]](#page-14-4)); (b) Paleogeography of Yangtze Sea during the Late Ordovician (Modified after [\[1\]](#page-14-0)); (c) Biostratigraphic division of Well ZY5 (Modified after [\[24\]](#page-15-15)). Paraorthograp*thograptus pacificus = P. pacificus, M. extraordinarius* = *Metabolograptus extraordinarius. tus pacificus* = *P. pacificus*, *M. extraordinarius* = *Metabolograptus extraordinarius*.

#### **3. Materials and Analytical Methodology 3. Materials and Analytical Methodology**

# *3.1. Samples and Experimental Methods 3.1. Samples and Experimental Methods*

A total of 16 fresh core samples were collected from Well ZY5 and weathered surfaces A total of 16 fresh core samples were collected from Well ZY5 and weathered surfaces of samples were removed prior to analysis, and then washed and dried. Core samples of samples were removed prior to analysis, and then washed and dried. Core samples were analyzed for TOC, sulfur contents of pyrite  $(S_{py})$ ,  $\delta^{34}S_{py}$ , major and trace elements, reactive  $\lim_{\text{from}}$  (Fe<sub>HR</sub>) and pyrite particle size statistics. All analyses were carried out at the State Key Laboratory of Biogeology and Environmental Geology, China University of Geosciences  $(Wuhan)$ .

For TOC analysis, approximately 100 g sample finer than 80 mesh size was treated For TOC analysis, approximately 100 g sample finer than 80 mesh size was treated with hydrochloric acid (1:7 HCl: water) at 50 $\degree$ C to remove the carbonates, neutralized until to a pH of 7.0 by adding distilled water, then the sample was transferred to a porcelain lain crucible and dried in an oven for testing and analysis using a Jena Eltra 4000 carbon-crucible and dried in an oven for testing and analysis using a Jena Eltra 4000 carbon-sulfur analyzer (Jena, Germany). The analytical error of TOC measurements was better than 0.2%. Pyrite framboids were measured using a scanning electron microscope (Hitachi-SU8010) (Hitachi, Tokyo, Japan). A size population more than 100 pyrite framboids were selected to measure diameters. Our methodology of pyrite particle size analysis is described in Xi et al. [\[24](#page-15-15)[,25\]](#page-15-16).

For major-element analysis (e.g., Al, P, and Si), the predried sample powders of smaller than 200 mesh were selected calcined for 4 h in a 1000  $\degree$ C high-temperature furnace to remove organics. Approximately 0.5 g of calcined sample was treated with anhydrous lithium tetraborate (eight times the sample mass), 0.4 g lithium fluoride and 0.3 g ammonium nitrate, accurately weighed into a porcelain crucible. After being completely mixed, the mixture was transferred to a platinum crucible, treated with a lithium bromide solution and placed on an electric hot plate for drying. The crucible was subsequently fused at 1150  $\degree$ C with an automatic flame melting to form a borate glass disks, and then measured by X-ray fluorescence spectrometry with a test accuracy deviation of less than 3%.

For trace elements analysis (e.g., Mo and U), samples  $(0.5 g)$  of smaller than 200 mesh particle size were placed in crucibles and calcined for 4 h in a 700  $\degree$ C high-temperature furnace to remove organics. Approximately 40 mg of calcined sample were dissolved using a standard multi-acid digestion (HF-HNO<sub>3</sub> $-$ HCl). The solution was diluted in 2% HNO<sub>3</sub> and measured using inductively coupled plasma mass spectrometry (ICP~MS) (Nu Instruments, Wrexham, UK) with a test accuracy deviation of less than 5%.

Highly reactive Fe (Fe<sub>HR</sub>) is contained in carbonate iron (Fe<sub>carb</sub>), oxide/hydroxide iron (Fe<sub>ox</sub>), magnetite iron (Fe<sub>mag</sub>), and pyrite (Fe<sub>py</sub>). The pyrite iron concentration (Fe<sub>py</sub>) and pyrite sulfur  $(S_{py})$  was extracted by using the chromium reduction method described by Canfield et al. [\[26\]](#page-15-17). Powdered samples were treated with an acidified  $CrCl<sub>2</sub>$  solution, the liberated hydrogen sulfide was trapped by silver-nitrate solution through precipitation as Ag<sub>2</sub>S. Fe<sub>py</sub> and S<sub>py</sub> was then gravimetrically calculated. Fe<sub>carb</sub>, Fe<sub>ox</sub> and Fe<sub>mag</sub> were determined using the sequential extraction procedure of Poulton and Canfield [\[27\]](#page-15-18). All concentrations of extracted iron speciation were diluted to a suitable concentration with 2%  $HNO<sub>3</sub>$  and tested by atomic absorption spectrometry (AAS); analytical errors were under 5%.

For  $\delta^{34}S_{\text{py}}$  analysis, the Ag<sub>2</sub>S precipitates from pyrite extractions were combusted with excess of  $V_2O_5$  in preparation for measurement of its S-isotopic composition ( $\delta^{34}S_{py}$ ). Isotope measurements were made using an elementary analyzer-stable isotope ratio mass spectrometer (EA-IRMS) (Finnigan MAT, Bremen, Germany), and the sulfur isotope adopted the V-CDT standard with accuracy better than  $\pm 0.2$ %. based on replicate analyses of standards IAEA S1 ( $\delta^{34}S = -0.3\%$ ), IAEA S2 ( $\delta^{34}S = +22.65\%$ ) and IAEA S3 ( $\delta^{34}S = -32.5\%$ ).

## *3.2. Background: Proxies Used in this Study*

## 3.2.1. Paleoredox Proxies

The concentrations and enrichment characteristics of Mo and U are often used to reconstruct the redox environment. Mo exists in molybdate anion  $(MO_4^2)$  form in oxic water columns, and is converted to particle-reactive thiomolybdates (MoS<sub>4-x</sub>O<sub>x</sub><sup>2-</sup>, x = 1–4) in euxinic environments. In the latter form, it can be taken up by Fe-Mo-S nano-minerals or organo-metallic ligands in organic matter, resulting in high abundances in the sediment [\[28\]](#page-15-19). Therefore, strong enrichments of Mo are a salient characteristic of euxinic depositional environment. U is dissolved in water as  $U^{6+}$  ( $UO_2(CO_3)_{3}^{4-}$ ) in oxic and suboxic water environments, and is converted to  $U^{4+}$  around the Fe(II)-Fe(III) redox boundary [\[13,](#page-15-5)[14\]](#page-15-6). Reduced U is usually insoluble and easily adsorbed by sediments. Consequently, strong enrichments of U indicate an anoxic environment, whereas lower enrichments of these trace metals are consistent with an oxic-suboxic water environment [\[13\]](#page-15-5). Trace-metal enrichment factors (EF) were calculated as:  $X_{EF} = (X/AI)_{sample} / (X/AI)_{PAdS}$ ;  $X_{EF} > 3$  was interpreted to reflect detectable enrichment relative the average shale, while  $X_{EF}$  < 1 indicates the depletion concentrations. In the modern ocean, Mo concentrations greater than 100 ppm represent an euxinia environment, concentrations between 25 and 100 ppm represent an intermittent euxinia environment, and concentrations less than 25 ppm represent a non-euxinic environment [\[29\]](#page-15-20).

U/Th and  $C_{org}/P_{total}$  indexes are positively correlated with the redox environment, which always increase with the increase of water anoxic degree. The U/Th ratio is usually greater than 1.25 in an anoxic environment, between 0.75 and 1.25 in a suboxic environment, and less than 0.75 in an oxic water column environment [\[30\]](#page-15-21). Algeo and Ingall [\[31\]](#page-15-22) proposed

that  $\frac{C_{org}}{P_{total}}$  less than 50 reflects an oxic water column environment, between 50 and 200 represents a suboxic water column environment, and greater than 200 suggests anoxic.

The size distribution of pyrite framboids has been used to determine water-column redox conditions in modern and ancient sedimentary basins, and is largely controlled by growth time and rate. The mean size and standard deviation of pyrite can reflect the redox environment of the ancient ocean [\[16](#page-15-7)[,32\]](#page-15-23). In an oxic water environment, pyrite can only be formed under the redox interface. Sulfate in the water column is continuously reduced by sulfate-reducing bacteria to form hydrogen sulfide (H<sub>2</sub>S) under the redox interface. Since hydrogen sulfide (H2S) and elemental Fe are continuously supplied in small amounts, they have a relatively slow growth rate and longer growth time, resulting in framboids with sizes greater than 7.7  $\pm$  4.1 µm. In anoxic water environments, framboids may form within the water column near the redox interface. Since the environment is rich in ferrous ions and (H2S) and there is no sediment support, it will soon sink into the ocean floor. Therefore, framboids are unable to achieve diameters much greater than  $5.0 \pm 1.7 \,\mu m$  [\[32\]](#page-15-23).

### 3.2.2. Paleoproductivity Proxies

Nutrient elements (Cu, Ni and Zn) in sediments can be used to reconstruct paleoproductivity. This is because oceanic primary productivity in the euphotic zone is significantly influenced by the availabilities of nutrient elements in seawater, and their concentrations decrease from the surface to the deep sea [\[33](#page-16-0)[,34\]](#page-16-1), which is finally buried in the sediments together with the sinking organic matter. Therefore, their sedimentary enrichments can also be used to infer the oceanic primary productivity. Generally speaking, the synthetic substitution of multiple elements is more reliable than that of a single element as a productivity indicator. Chen et al. [\[35\]](#page-16-2) used principal component analysis to construct the substitution index of paleoproductivity including excess copper  $(Cu_{Xs})$ , excess nickel  $(Ni_{Xs})$ and excess zinc  $(Zn<sub>Xs</sub>)$  to indicate the paleoproductivity strength. The calculation is as follows:

$$
Y = 0.593955w(Cu_{XS}) + 0.523396w(Ni_{XS}) + 0.610962w(Zn_{XS})
$$
 (1)

where  $X_{XS}$  represents the excess value of trace elements (e.g., Cu and Ni), the correction formula is written as:  $X_{\text{Xs}} = X_{\text{total}} - T_{\text{total}} \times (X/T_{\text{I}})_{\text{PAAS}}$ ,  $X_{\text{total}}$  and  $T_{\text{total}}$  represent the elements to X and Ti sample values, respectively, and  $(X/Ti)_{PAAS}$  represents the ratio of element to Ti in post-Archean Australian Shale (PAAS).

Excess silica ( $SiO<sub>2(exc)</sub>$ ) refers to the silicon of the non-clay origin and is used to characterize the biogenic silicate portion [\[36\]](#page-16-3). The calculation is as follows:

$$
SiO2(exc) = [SiO2]sample - [Al2O3]sample \times (SiO2/Al2O3)debris
$$
 (2)

where  $(SiO_2/Al_2O_3)_{\text{debris}} = 2.28$  is based on the content of major elements relationships.  $[SiO2]_{sample}$  and  $[A1_2O_3]_{sample}$  represent sample  $SiO_2$  and  $Al_2O_3$  values, respectively. Specifically, CaO and  $Al_2O_3$  have a moderate negative correlation when the intercept of the x-axis is 22.3%, which reflects the  $Al_2O_3$  content in the pure shale portion of the study area (Figure [2a](#page-5-0)). When  $\text{Al}_2\text{O}_3$  is 22.3%, based on the SiO<sub>2</sub> vs.  $\text{Al}_2\text{O}_3$  crossplot the pure shale fraction of  $SiO<sub>2</sub>$  content is estimated to be 50.8% (Figure [2b](#page-5-0)).

<span id="page-5-0"></span>

Figure 2. Correlations between CaO vs  $Al_2O_3$  (a), and SiO<sub>2</sub> vs  $Al_2O_3$  (b) from well ZY5. One outlier with high carbonate content (CaO = 11.1%) was not viewed as representative. content (CaO = 11.1%) was not viewed as representative.

# 3.2.3. Paleoclimate Proxies 3.2.3. Paleoclimate Proxies

The chemical alteration index (CIA) was developed to reflect provenance chemical The chemical alteration index (CIA) was developed to reflect provenance chemical weathering intensity [\[37,](#page-16-4)[38\]](#page-16-5). In recent years, it has been applied to study paleoclimate weathering intensity [37,38]. change in the process of sedimentation [39]. change in the process of sedimentation [\[39\]](#page-16-6).

$$
CIA = Al_2O_3/(Al_2O_3 + CaO^* + Na_2O + K_2O) \times 100
$$
 (3)

In the formula, the unit of each oxide is the mole fraction, and  $CaO^*$  represents  $CaO$ of composition silicate mineral fraction  $[40]$ . Therefore, it is necessary to calibrate CaO in carbonate minerals (calcite and dolomite) and phosphate minerals (apatite). This paper used phosphate for CaO calibration according to the formula CaO\* = CaO  $-$  P<sub>2</sub>O<sub>5</sub>  $\times$  10/3. If the number of moles of CaO\* obtained was less than the number of moles of Na<sub>2</sub>O, If the number of moles of  $CaO^*$  obtained was less than the number of moles of Na<sub>2</sub>O, the calculated value was used to replace  $CaO^*$ , If the mole number of  $CaO^*$  was greater than the mole number of Na<sub>2</sub>O, the mole number of Na<sub>2</sub>O was used instead of CaO\* [\[41\]](#page-16-8). Additionally,  $K_2O$  was further corrected for K-metasomatism and metamorphism of the source rocks: K<sub>2</sub>O<sub>corr</sub> = molar[m × Al<sub>2</sub>O<sub>3</sub> + m × (CaO\* + Na<sub>2</sub>O)]/(1 – m), where m =  $K_2O/(Al_2O_3 + CaO* + Na_2O + K_2O)$  [\[42\]](#page-16-9). Generally speaking, a CIA of 50~65 represents a low degree of chemical weathering intensity in cold or arid climate conditions. CIA values of 65~85 indicate moderate chemical weathering intensity in warm and humid climate conditions. CIA values of 85~100 suggest strong chemical weathering intensity in hot and humid climate conditions [\[37\]](#page-16-4).

# **4. Results 4. Results**

## *4.1. Paleoredox Data 4.1. Paleoredox Data*

The U/Th and  $\rm{C_{org}/P_{total}}$  ratios have a similar vertical trend in well ZY5. U/Th values ranged from 0.15~0.31(mean 0.22) in the *P. pacificus* zone and increased to 0.21~0.48 (mean ranged from 0.15~0.31(mean 0.22) in the *P. pacificus* zone and increased to 0.21~0.48 (mean 0.33) in the *M. extraordinarius* zone. C<sub>org</sub>/P<sub>total</sub> ratios are bteween 5.50~13.79 (mean 9.94) in the *P. pacificus* zone and increased with significant fluctuations to 7.20~15.20 (mean 11.13) the *P. pacificus* zone and increased with significant fluctuations to 7.20~15.20 (mean 11.13) in the *M. extraordinarius* zone. in the *M. extraordinarius* zone.

The concentrations and enrichment factors of Mo and U have broadly similar trends The concentrations and enrichment factors of Mo and U have broadly similar trends as the U/Th and C<sub>org</sub>/P<sub>total</sub> ratios (Figure [3\)](#page-6-0). The *P. pacificus* zone exhibited extremely low values: Mo =  $0.39~0.84$  ppm (mean  $0.57$  ppm), U =  $1.12~5.26$  ppm (mean = 2.78 ppm), and Mo<sub>EF</sub> = 0.87~2.35 (mean 1.34), U<sub>EF</sub> = 0.42~0.87 (mean 0.63). The *M. extraordinarius* zone had relatively higher concentrations and enrichment factors compared with the *P. pacificus* zone: Mo =  $0.53~6.11$  ppm (mean 3.47 ppm), U =  $4.09~8.72$  ppm (mean 6.07 ppm), and  $M_{\text{OEF}} = 0.73 \sim 9.98$  (mean 6.23),  $U_{\text{EF}} = 0.62 \sim 1.56$  (mean 1.17) (Table S1).

<span id="page-6-0"></span>

 $F_{\text{R}}$   $\sim$   $F_{\text{R}}$ ,  $\sim$ **Figure 3.** Chemostratigraphic profiles of TOC, Fe<sub>HR</sub>,  $\delta^{34}$ Spy, Mo-U concentrations, Mo<sub>EF</sub> and U<sub>EF</sub>, C<sub>org</sub>/P<sub>totl</sub> and U/Th of the well ZY5. Lithological legend and abbreviations as in Figure [1.](#page-2-0) well ZY5. Lithological legend and abbreviations as in Figure 1.

## The TOC contents vary between 0.17%~2.28% (mean 1.29%) in the *P. pacificus* zone, *4.2. Productivity Data 4.2. Productivity Data*

The TOC contents vary between 0.17%~2.28% (mean 1.29%) in the  $P.$   $\emph{pacificus zone,}$  $\frac{1}{270}$  increased, ranging from 20.8% to 75.4% (mean 54.8%) in the *P. pacificus zone, and the P. pacificus* zone, and the *P. pacificus* zone, and the *P. pacificus*  $\frac{1}{270}$  and the *P. pacificus in the P. pacifi* and  $0.73\% \sim 1.35\%$  (mean  $1.07\%$ ) in the *M. extraordinarius* zone. The  $SiO_{2(exc)}$  contents first increased, ranging from 29.8% to 75.4% (mean 54.8%) in the P. pacificus zone, and then decreased ranging from 26.7% to 41.8% (mean 33.2%) (Figure 4). The Y values displayed similar trends as SiO<sub>2(exc)</sub>, gradually increasing from 77.4 to 216.6 (mean 155.6) in the *P*. *pacificus* zone, and then decreasing from 63.4 to 150.7 (mean 120.6) in the *M. extraordinarius pacificus* zone, and then decreasing from 63.4 to 150.7 (mean 120.6) in the *M. extraordinarius* zone (Table S2).

<span id="page-6-1"></span>

**Figure 4.** Profiles of Y,  $SiO_{2(exc)}$  and chemical index of alteration (CIA) values from the well ZY5. Vertical dashed lines show the approximate ranges between cold or arid climate conditions, and warm and humid climate conditions. Lithological legend and abbreviations as in Figure [1.](#page-2-0)

## *4.3. Sulfur Content and δ <sup>34</sup>Spy and Framboidal Pyrite Size Distribution 4.3. Sulfur Content and δ34Spy and Framboidal Pyrite Size Distribution*

The sulfur contents of pyrite  $(S_{py})$  range between 0.11%~5.31% (mean 1.12%) in well ZY5. Deposits of the *P. pacificus* zone are bteween 0.11%~0.72% (mean 0.45%), gradually increasing to 0.78%~5.31% (mean 1.89%) in the *M. extraordinarius* zone (Table S1).  $\delta^{34}S_{py}$ values range between  $-24.83\%$  to  $-3.61\%$  in the *P. pacificus zone*, in the *M. extraordinarius* zone they decreased from  $-0.36\,\%$  to  $-25.19\,\%$  in the early stage and increased to  $-10.57\,\%$ in the middle stage (Figure 3).  $\frac{1}{2}$ 

 $V_{\rm eff}$  dashed lines show the approximate ranges between cold or arid climate conditions, and  $\alpha$ 

Size distribution analysis for in the early of the *P. pacificus* zone was not done due to lack of pyrite. In the middle and late stages of the P. Pacificus zone pyrite framboids were generally large with mean sizes ranging from 5.6  $\mu$ m to 6.5  $\mu$ m (standard deviations =  $1.8 \mu m$ ). It is noteworthy that irregular, large pyrite is present in these deposits (Figure 5a). The  $M$ . *extraordinarius* zone pyrite framboids had smaller grain sizes with mean sizes ranging from 5.5  $\mu$ m to 7.2  $\mu$ m (standard deviations = 3.2  $\mu$ m) (Figure 6). Relatively small pyrite framboids and massive pyrite are present in *M. extraordinarius* zone deposits (Figure 5b,c). 5b,c). rite fra[m](#page-7-0)boids and massive pyrite are present in *M. extraordinarius* zone deposits (Figure

<span id="page-7-0"></span>

Figure 5. Morphological characteristics of pyrite in the shale samples of Well ZY5: (a) euhedral pyrite aggregate, sample of P. pacificus zone; (b,c) Pyritohedral microcrystals in tiny framboidal pyrite, sample of M. extraordinarius zone. of *P. pacificus* zone; (**b**,**c**) Pyritohedral microcrystals in tiny framboidal pyrite, sample of *M. extraordinarius* zone.

<span id="page-7-1"></span>

Stratigraphy Depth Lithology		(m)	Number Pyrite-framboidal size distribution						Average	Standard deviation	Sample	
			$\overline{0}$	Ś.	10	15	20	25	$30$ (um)			amount
Hirnantian		2096		CETT	-cà	$\bigcirc$			ZY3-26 $\circ$		$6.9 \pm 4.8$	$N=40$
		2097		$\bigoplus$	000 000				$O: OZY3-28$		$7.2 \pm 5.3$	$N = 65$
					COCHIDROD: 0 0				$ZY5-29$		$5.8 \pm 2.7$	$N = 57$
	M.extraordinarius.	2098		Gee D					ZY5-30		$5.5 \pm 1.4$	$N = 66$
		2099		<b>OCCI</b>	HO- $\oplus$				$ZY5-31$		$6.3 \pm 3.2$	$N=20$
		2100										
Late Katian		2101		December 0					$ZY5-34$		$6.2 \pm 2.0$	$N=19$
		2102										
	P. pacificus	2103										
		2104		$\oplus$	$\mathbb{D}$				$ZY5-41$		$6.5 \pm 1.6$	$N = 28$
		2105								First quartile CCX	<b>COOO</b>	Median $h$ <sup>Maximum</sup> Outliers
		2106		$\sim$						Minimum	Third quartile Mean	

**Figure 6.** Pyrite framboid size distribution and 'box and whisker' plots from Late Ordovician shale in well ZY5. Lithological legend and abbreviations as in Figure [1.](#page-2-0)

## *4.4. Paleoclimatic Data 4.4. Paleoclimatic Data*

A triangle map of A-CN-K can indicate the chemical weathering trend and the change A triangle map of A-CN-K can indicate the chemical weathering trend and the of original composition and mineral characteristics during weathering [\[37\]](#page-16-4). In the process of weathering, plagioclase is weathered more easily than potassium feldspar, resulting in a higher loss rate of Ca and Na elements than the K elements. Therefore, the weathering trend lines of rocks without potassium metasomatism and diagenetic transformation are parallel to the A-C-N line in the ternary diagram.

**Figure 6.** Pyrite framboid size distribution and 'box and whisker' plots from Late Ordovician shale

The projection points of the samples in the study area deviated from the original weathering trend line (center line a in Figure [7d](#page-8-0)), indicating that shale in Well ZY5 experienced a certain amount of potassium metasomatism. After correction, the CIA<sub>corr</sub> values were higher and relatively stable in the *P. pacificus* zone, fluctuating between 68.7<br>were higher and 75.9 (mean = 72.9). Further upwards, CIA values declined, ranging from 59.9 to 72.2 (mean = 67.6) in the *M. extraordinarius* zone (Figure [4,](#page-6-1) Table S2). 67.6) in the *M. extraordinarius* zone (Figure 4).

<span id="page-8-0"></span>

Figure 7. (a) Crossplot of TiO<sub>2</sub> versus Zr; (b) Crossplot of Th versus Sc; (c) La-Sc-Th ternary diagram showing source rock types from Well ZY5 and (d) A-CN-K (Al<sub>2</sub>O<sub>3</sub>-(CaO<sup>\*</sup> + Na<sub>2</sub>O) – K<sub>2</sub>O) ternary diagram showing diagenetic K-metasomatism and the chemical index of alteration (CIA) ranges for different chemical weathering degrees. and the chemical index of alteration (CIA) ranges for different chemical weathering degrees.

## **5. Discussion**

## <span id="page-9-1"></span>*5.1. Reconstruction of Redox Environments at Well ZY5 during the Late Ordovician*

The U/Th and C<sub>org</sub>/P<sub>total</sub> data suggested that the shale from the *P. pacificus* zone were deposited in an oxic water column environment. An interpretation consistent with changes of pyrite framboid sizes, the dearth of pyrite framboids in early stage (at depth 2105–2106 m), and then the presence of abundant large pyrite framboids in late stage (at depth 2101–2104 m) (Figure [6\)](#page-7-1). The enrichment characteristics and concentrations of Mo and U were similar to crustal values ( $M$ <sub>OEF</sub> = 1.34, U<sub>EF</sub> = 0.63; Mo = 0.57 ppm, U = 2.78 ppm; Figure [3\)](#page-6-0), which further implicated an oxic water column environment during the *P. pacificus* zone deposition.

The distribution of *M. extraordinarius* zone data on U/Th and C<sub>org</sub>/P<sub>total</sub> ratios showed that deposition occurred under an oxic-suboxic environment (Figure [3\)](#page-6-0), and it was a more reducing bottom water environment compared with deposition of the *P. pacificus* zone. This was consistent with the presence of abundant relatively small pyrite framboids (Figure [6\)](#page-7-1). Mo and U showed obvious enrichment characteristics relative to crustal values in the *M. extraordinarius* zone (Mo<sub>EF</sub> = 6.23, U<sub>EF</sub> = 1.17; Figure [3\)](#page-6-0), Overall, well ZY5 shifted toward a more reduction water column environment from the *P. pacificus* zone to the *M. extraordinarius* zone.

### *5.2. Mechanisms Controlling the Paleoredox Environment*

### <span id="page-9-0"></span>5.2.1. Paleoproductivity

The significant increase of ocean surface productivity promoted the downward migra-tion of the redox interface and intensification of bottom-water anoxia [\[12\]](#page-15-4). Y and  $SiO_{2(exc)}$ content variability from the *P. pacificus* to *M. extraordinarius* zones in well ZY5 indicated that the productivity level of the *P. pacificus* zone was significantly higher than the *M. extraordinarius* zone. This was basically consistent with the vertical paleoproductivity variability characterized by the organic matter accumulation rate (OCAR) in well PY1 [\[10\]](#page-15-2) and Datianba section [\[8\]](#page-15-1) (Figure [7\)](#page-8-0). The higher productivity in the *P. pacificus* zone indicated that higher nutrient inputs (e.g., P, Si, Cu and Zn) led to the growth of primary producers (e.g., algae) in the inner Yangtze Sea, and consumer plankton (such as radiolarians) began to flourish. The atmospheric  $O_2$  levels were relatively low during the Late Ordovician [\[31\]](#page-15-22). Biogeochemical modeling has shown that, at low atmospheric  $O_2$  levels, dissolved  $O_2$ concentrations in surface seawater are controlled principally by local productivity through photosynthetic  $O_2$  release, with higher dissolved  $O_2$  contents in regions of higher productivity [\[12,](#page-15-4)[43\]](#page-16-10). Therefore, this produced better-oxygenated conditions in the *P. pacificus* zone.

The decline of productivity in the *M. extraordinarius* zone was induced by a combination of: (1) formation of the Gondwana continental glacier, which led to a significant drop in global sea level [\[44](#page-16-11)[,45\]](#page-16-12), and (2) the sharp uplift of the Xiang'e Submarine High due to the increasing compression of Cathaysian plate to Yangtze plate [\[46\]](#page-16-13), which resulted in a weak exchange rate between the inner and outer Yangtze Sea during this period [\[10\]](#page-15-2). This led to a decline of nutrient input. However, the relatively restricted water mass was conducive to the formation of relatively reduced water conditions [\[5\]](#page-14-3).

### 5.2.2. Provenance Composition of Shale and Paleoclimate Characteristics

Nutrient and sulfate levels have important effects on the redox environment of the water column, and paleoclimate characteristics and provenance composition are the key factors controlling nutrient and sulfate fluxes. The source rock composition of fine-grained clastic rocks can be distinguished using main and trace elements in sediments. Elements such as Th, Sc and Zr are difficult to dissolve during geological processes such as migration, diagenesis and weathering, and are reliable indicators to distinguish the composition of source rocks  $[47,48]$  $[47,48]$ . When TiO<sub>2</sub>/Zr > 200 it indicates mafic igneous rocks, intermediate igneous rocks range between 55 and 195, and values lower than 55 are felsic igneous rocks [\[49\]](#page-16-16).

The TiO<sub>2</sub>/Zr ratios for all shale samples ranged from 9.1 to 58.3 (mean 44.6), indicating that the source rocks were felsic igneous rocks (Figure [7a](#page-8-0)), such as granite and granodiorite. Correlations between the La-Te-Sc trigram and Th/Sc diagram of black shale samples in the study area (Figure [7b](#page-8-0),c) also indicated that the source rocks were mostly granodiorite, and a few samples were granite. Therefore, the source rocks in the study area were mainly felsic igneous rocks, consistent with findings that the bedrock of the Yangtze craton is composed of granitic gneiss and amphibolite [\[50\]](#page-16-17).

CIAcorr values in the *P. pacificus* zone were relatively high and stable in the study area (Figure [7\)](#page-8-0), ranging from 68.7 to 75.9 (mean 72.9), and suggest that source rocks underwent a moderate degree of chemical weathering under a warm and humid climate.  $CIA<sub>corr</sub>$ values declined, varying between 59.9 and 72.2 (mean = 67.6) in the *M. extraordinarius* zone, indicating a gradual evolution from warm and humid climate to cold and dry climate conditions, and weaker chemical weathering intensity of the source area under relatively cooler and more arid conditions.

Within local and short time scales, water redox state is mainly controlled by the supply balance of organic matter and various oxidants within the region [\[51](#page-16-18)[–53\]](#page-16-19). The composition of the source rock undergoing weathering will cause differences in the migration and enrichment of nutrient elements and oxidants to the ocean [\[12\]](#page-15-4). Late Ordovician shale in the study area was sourced from felsic igneous rocks, which would have required chemical weathering of the silicate mineral matrix to expose the contained sulfides to the atmosphere [\[52\]](#page-16-20). Therefore, we inferred that the difference in chemical weathering intensity was the main factor controlling sulfate weathering variability and nutrient flux in the inner Yangtze Sea during the *P. pacificus* zone and the *M. extraordinarius* zone.

As recently shown, volcanism plays an important role in global climate change [\[54\]](#page-16-21). The *P. pacificus* zone has experienced a warm and humid climate, it may be related to the global warming around the Katian-Hirnantian boundary caused by the greenhouse gases generated by the potential LIP-scale eruptions [\[55\]](#page-16-22). This is confirmed by recent reports of mercury enrichment and K-bentonites in South China, Laurencia and Poland during the late Ordovician [\[56](#page-16-23)[–58\]](#page-16-24). The *P. pacificus* zone experienced a moderate degree of chemical weathering intensity in warm and humid climates. Sulfides such as pyrite accounted for a higher proportion than paste salts and other sulfates in the felsic source rocks. The sulfides were oxidized to produce sulfates in the process of chemical weathering during this time, which acted as a medium to accelerate the weathering of source rocks and were transported into the ocean by rivers. On the other hand, the abundance of nutrients from volcanic [\[58\]](#page-16-24) and weathering sources entering the ocean is significantly enhanced, and the suitable temperature provides necessary conditions for the mass reproduction of hydrocarbon forming organisms (e.g., sponge spicule and graptolite) in surface sea-water [\[57\]](#page-16-25), which stimulated the increase of oceanic primary productivity and promoted the expansion of surface oxic water environments. However, the *M. extraordinarius* zone was under cold and dry climatic conditions. Chemical weathering was weak due to low temperatures and limited sediment, and this cooling was accompanied by a reduction of the hydrological cycle, which caused the input of nutrients in the Inner Yangtze Sea to be significantly lower than in the *P. pacificus* zone, resulting in a decline in productivity. In addition, due to the coupling of global cooling and the uplift of the Xiang'e Submarine High, the inner Yangtze Sea became a stagnant and low-energy environment, which resulted in more reducing surface water conditions compared with the *P. pacificus* zone.

#### *5.3. Regional Comparison of Redox Environments*

The spatial patterns of redox variation in the Late Ordovician can be assessed within the biostratigraphic framework proposed by Chen et al. and Xi et al. [\[22](#page-15-13)[,23\]](#page-15-14). This framework allows us to compare redox variation and explore further the mechanism of the difference in the spatial development of redox environment in the Inner Yangtze Sea during the *P. pacificus* zone and the *M. extraordinarius* zone.

## 5.3.1. The *P. pacificus* Zone 5.3.1. The *P. pacificus* Zone

The Xiushan Datianba section was deposited in the northeastern margin of the The Xiushan Datianba section was deposited in the northeastern margin of the QiQianzhong oldland (Figure 1[b\).](#page-2-0) The relatively high  $Fe_{py}/Fe_{HR}$  profile in the Datianba section of the *P. pacificus* zone indicated deposition in an euxinic-dominated environment [\[8\]](#page-15-1) (Figure [8\)](#page-11-0). Water-column euxinic environment is generally controlled by the relative fluxes of C, S and Fe in mid-depth waters, which is related to oceanic primary productivity, sulfate concentrations, and  $Fe_{HR}$  supply  $[53,59,60]$  $[53,59,60]$  $[53,59,60]$ . The connection between inner and outer Yang-outer Yangtze seas (Figures [1b](#page-2-0) and [9a](#page-12-0)) was beneficial to the relatively high productivity level in its le level in its sedimentary history, and the pattern of the relatively high TOC content distribution through the upper *P. pacificus* zone of the Datianba section is similar to that from the other to the other worldwide sites (e.g., Zbrza section in Poland, Arctic Canada) [\[61](#page-16-28)[,62\]](#page-17-0), which the conmay be controlled by the "Boda" warming event. Marine productivity is sufficiently high they be combined by the "Boda" warming event. Marine productivity is sufficiently high (see Section [5.2.1\)](#page-9-0), and a relatively low Fe<sub>HR</sub> fluxes (Datianba sections mean =  $1.4\%$ , [\[8\]](#page-15-1)) for the *P. pacificus* zone. Ample sulfate availability may be the main factor affecting the for the *P. pacificus* zone. Ample sulfate availability may be the main factor affecting the development of euxinia at mid-depth waters. *culture is party and such a might summer availability may be the main factor affecting the* 

<span id="page-11-0"></span>

Figure 8. Comparison of productivity indicators, S<sub>py</sub>,  $\delta^{34}$ S<sub>py</sub> in the Katian-Hirnantian shales of Well ZY5 in the study area, Fe-speciation data and  $\delta^{34}S_{py}$  data for the Datianba sections are from [\[8\]](#page-15-1), OCAR (mass accumulation rates of organic carbon) values and  $\delta^{34}S_{py}$  data for the Well PY1 are from [\[10\]](#page-15-2). *M. extra.* = *M. extraordinarius*, HF = Hirnantia Fauna.

<span id="page-12-0"></span>

Figure 9. Schematic of a stratified redox model for the Late Ordovician inner Yangtze Sea. (a) oceanic redox structure in the the *P. pacificus* zone; (**b**) oceanic redox structure in early the *M.extraordinariu*s zone, (**c**) oceanic redox structure during the P. pacificus zone; (b) oceanic redox structure in early the *M.extraordinarius* zone, (c) oceanic redox structure during the late [8,10]. *M.extraordinarius* zone. The sedimentary model is based on geochemical data from the Datianba section and Well PY1 [\[8,](#page-15-1)[10\]](#page-15-2).

bundle availability in a marine system can be investigated by using δ<sup>3</sup> p<sub>py</sub> [59,60]. The availability of sulfate and organic matter controls the rate of bacterial sulfate reduction (BSR) and affects the composition of  $\delta^{34}S_{py}$  in the BSR process [\[49\]](#page-16-16). Increasing productivity levels may consume sulfate and reduce the sulfur isotope fractionation between sulfate and sulfide ( $\epsilon_{BSR} = \delta^{34}S_{SO4-} \delta^{34}S_{Py}$ ), which causes the  $\delta^{34}S_{Py}$  to become heavier [\[63\]](#page-17-1). Laboratory experiments have suggested that when sulfate concentrations fall below a Laboratory experiments have suggested that when sulfate concentrations fall below a experiments have suggested that when sulfate concentrations fall below a biological biological threshold of 200 µM, the fractionation between sulphate and sulfide during threshold of 200 μM, the fractionation between sulphate and sulfide during BSR is re-BSR is reduced [\[64\]](#page-17-2), leading to small or zero values for *ε*BSR and thus isotopically heavier duced [64], leading to small or zero values for ɛBSR and thus isotopically heavier pyrites. pyrites. The δ <sup>34</sup>Spy values formed within the water column in the *P. pacificus* zone (e.g., PY1 and Datianba section) were relatively light (Figure [8\)](#page-11-0). Wang et al. [\[65\]](#page-17-3) suggested that the lighter conversion of  $\delta^{34}S_{py}$  may have been controlled by a decrease in sulfate reduction rate or an increase in sulfate concentrations. The clear negative correlation between OCAR and  $\delta^{34}S_{\text{max}}$  in well PY1 [10] indicated that high biological productivity was enough to drive  $\mathbf{r}_1$  in well productivity was enough to drive was enough Sulfate availability in a marine system can be investigated by using  $\delta^{34} \mathrm{S_{py}}$  [\[59](#page-16-26)[,60\]](#page-16-27). The and  $\delta^{34}S_{\rm py}$  in well PY1 [\[10\]](#page-15-2) indicated that high biological productivity was enough to drive

sulfate reduction. Therefore, the most reasonable explanation may be that the seawater sulfate concentrations increased during BSR. An interpretation consistent with an increase of  $S_{\text{pv}}$  in the mid-depth waters of the study area (PY1 mean= 1.4%, Datianba sections mean = 1.6% [\[10\]](#page-15-2)). So ample sulfate accelerated the consumption of organic matter by BSR and produced  $H_2S$  accumulation in mid-depth water column, resulted to the formation and expansion of the euxinic water mass in the Datianba section [\[8\]](#page-15-1). Although  $\delta^{34}S_{\text{py}}$  values in the *P. pacificus* zone were light in well ZY5 (average = −14.5‰), it is always in an oxic water environment. Obviously, the sulfate availability may not be the main factor affecting the redox of Well ZY5.

According to the reconstructed paleotopography, well ZY5 was deposited in areas of shallow seas. Under the warm and humid climate conditions of the *P. pacificus* zone, the input of nutrients with a rise in sea level promoted the growth of algae and other organisms on the surface of the ocean, while plankton released a large amount of  $O_2$ . In shallow seas,  $O_2$  takes precedence over  $SO_4^2$ <sup>-</sup> oxidize the organic matter in the sinking process [\[66\]](#page-17-4), and  $O_2$  has a reactive absorption effect on  $H_2S$  and inhibites BSR [\[67\]](#page-17-5). The consequent inadequated production of  $H_2S$  was not conducive to the persistently formation of the euxinic water mass to shallow areas, so well ZY5 water mass was an oxic environment (Figures [3](#page-6-0) and [9a](#page-12-0)). Biogeochemical models have found that active biological activities caused the total respiratory  $O_2$  demand to exceed the supply of  $O_2$  to deep oceans through vertical mixing and diffusion [\[68\]](#page-17-6), which caused sulfates to play a leading role in the process of organic matter degradation in mid-depth waters [\[67\]](#page-17-5). As a result, the euxinic waters environment of the inner Yangtze was sandwiched between the oxic ocean-surface layer and the deep ferruginous waters conditions, showing a tapering wedge towards the deep ocean (Figure [9a](#page-12-0)).

### 5.3.2. The *M. extraordinarius* Zone

The *M. extraordinarius* zone, Fe<sub>py</sub>/Fe<sub>HR</sub> profile of the Xiushan Datianba section from relatively high in the early stage to low in middle and late stage, suggest that it was deposited in the euxinic to ferruginous transitions (Figures [8](#page-11-0) and [9b](#page-12-0),c). Generally speaking, ferruginous conditions may be formed in marine systems characterized by low marine sulfate, low productivity and high Fe $_{HR}$ . Due to the extensive euxinic environment formed in the *P. pacificus* zone of the inner Yangtze Sea [\[10\]](#page-15-2) the sulfate concentrations were gradually consumed. In addition, climate cooling resulted in a significant drop in global sea level [\[45,](#page-16-12)[47\]](#page-16-14), and the intense uplift of the Xiang'e Submarine High during this period, resulting in a weak exchange rate between the inner and outer Yangtze Sea, thus limiting sulfate and nutrient resupply from both riverine and outer Yangtze seas sources. The inference was consistent with the heavy values of  $\delta^{34}S_{\text{py}}$  and a decrease of  $S_{\text{py}}$  (PY1 mean = 1.1%, Datianba sections mean = 1.4%) in the study, in addition,  $SiO<sub>2(exc)</sub>$  and OCAR exhibits low values at *M. extraordinarius* zone (Figure [8\)](#page-11-0). Therefore, The H2S produced by BSR in mid-depth waters was insufficient to react  $Fe_{HR}$  and allow it to accumulate [\[54\]](#page-16-21). This was consistent with the relatively high of Fe<sub>HR</sub> ratios during the *M. extraordinarius* zone (well ZY5 mean =  $3.1\%$ , Datianba section mean =  $2.2\%$ ) (Figure [3\)](#page-6-0). It was beneficial to the formation and expansion of ferruginous waters conditions in the Datianba section. However, well ZY5 was located on the flank of the inner Yangtze Sea, along the southwest margin of the the Xiang'e Submarine High (Figure [1b](#page-2-0)), and significant uplift of the Xiang'e Submarine High during this time resulted in relatively shallow water compared to coeval deposits of the Datianba section, which was key to the formation of an oxic water environment in well ZY5. However, the enhancement of water mass restriction caused the expansion of ferruginous waters conditions, which was one of the reasons for the enhanced of reduction degree of surface water compared with the *P. pacificus* zone. On the other hand, nutrients from continental chemical weathering sources were cut off during climate cooling and relative sea level decline. The low nutrient input resulted in the decline of paleoproductivity of bottom waters in shallow areas (see Section [5.2.1\)](#page-9-0), which decreased  $O_2$  release intensity in the ocean surface waters, so water column environment of well ZY5 evolved from an

oxic environment to enhanced water reduction degree (Figure [3;](#page-6-0) see Section [5.1\)](#page-9-1). Therefore, with the shrinkage of euxinic and oxic water column conditions (Figure [9b](#page-12-0)), the water column of the inner Yangtze Sea evolved into an environment characterized by surface oxic waters and deep ferruginous waters conditions (Figure [9c](#page-12-0)).

## **6. Conclusions**

Shale in well ZY5 was deposited in an oxic water column environment, and the reduction degree of water in the *M. extraordinarius* zone was stronger than in the *P. pacificus* zone. High productivity level controlled  $O<sub>2</sub>$  release in the surface layer of the *P. pacificus* zone as well as the oxidation degree in the shallow water area of the inner Yangtze Sea. The decrease in productivity and the relatively restricted water mass of the inner Yangtze Sea controlled the formation of relatively reduced water conditions in the *M. extraordinarius* zone.

The change in CIA values from the *P.pacificus* zone to the *M. extraordinarius* zone showed that paleoclimate conditions changed from warm and humid to cold and dry in the study area. The geochemistry characteristics of the Late Ordovician Datianba sections and well ZY5 on the Upper Yangtze platform were useful in illustrating the redox evolution of the inner Yangtze Sea. In the *P. pacificus* zone, warm and humid climatic conditions drove the inner Yangtze Sea's high productivity levels, sulfate fluxes and low reactive iron flux, which promoted the spread of surface oxic water conditions and mid-depth euxinic water column conditions. The *M. extraordinarius* zone, the cold and dry climate and the uplift of the Xiang'e Submarine High co-contributed to the relative sea level decline, causing low productivity and sulfate flux, and high reactive iron flux, which promoted the expansion of mid-depth ferruginous water and the shrinkage of oxic surface water. The regional paleogeographic framework is an important reason for the spatial difference of redox environment.

**Supplementary Materials:** The following are available online at [https://www.mdpi.com/article/](https://www.mdpi.com/article/10.3390/min11070710/s1) [10.3390/min11070710/s1,](https://www.mdpi.com/article/10.3390/min11070710/s1) Table S1: Key geochemical data for marine redox evaluation and trace element at the Well ZY5, Table S2: Key geochemical data for calculation of marine productivity fluxes, CIA and S isotopic compositions at the Well ZY5.

**Author Contributions:** D.L., Z.X., Y.Y., S.T. and B.Z. conceived and designed the experiments; D.L., Z.X., B.Z. and Y.Y. performed the experiments and data curation; D.L., Z.X. and S.T. wrote and revised the paper. All authors have read and agreed to the published version of the manuscript.

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