

## Article

# Estimation of Reference Evapotranspiration during the Irrigation Season Using Nine Temperature-Based Methods in a Hot-Summer Mediterranean Climate

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**Abstract:** The FAO-56 Penman–Monteith (PM) equation is regarded as the most accurate equation to estimate reference evapotranspiration (E<sub>T0</sub>). However, it requires a broad range of data that may not be available or of reasonable quality. In this study, nine temperature-based methods were assessed for E<sub>T0</sub> estimation during the irrigation at fourteen locations distributed through a hot-summer Mediterranean climate region of Alentejo, Southern Portugal. Additionally, for each location, the Hargreaves–Samani radiation adjustment coefficient ( $k_{RS}$ ) was calibrated and validated to evaluate the appropriateness of using the standard value, creating a locally adjusted Hargreaves–Samani (HS) equation. The accuracy of each method was evaluated by statistically comparing their results with those obtained by PM. Results show that the calibration of the  $k_{RS}$ , a locally adjusted HS method can be used to estimate daily E<sub>T0</sub> acceptably well, with RMSE lower than 0.88 mm day<sup>-1</sup>, an estimation error lower than 4% and a R<sup>2</sup> higher than 0.69, proving to be the most accurate model for 8 (out of 14) locations. A modified Hargreaves–Samani method also performed acceptably for 4 locations, with a RMSE of 0.72–0.84 mm day<sup>-1</sup>, a slope varying from 0.95 to 1.01 and a R<sup>2</sup> higher than 0.78. One can conclude that, when weather data is missing, a calibrated HS equation is adequate to estimate E<sub>T0</sub> during the irrigation season.

**Keywords:** reference evapotranspiration; FAO Penman Monteith; Hargreaves–Samani; temperature-based ET methods; irrigation scheduling; hot summer Mediterranean climate



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## 1. Introduction

A simple method to estimate crop water requirements is through the computation of reference evapotranspiration (E<sub>T0</sub>). If accurate, these computations serve as a basis for several assessments such as water management, irrigation system design and management, irrigation scheduling, and crop modeling [1–9]. There are several methods for estimating E<sub>T0</sub>, being the FAO-56 application of the Penman–Monteith (PM) equation [4] widely regarded as the most accurate. The method provides consistent E<sub>T0</sub> values in many regions and climates [10,11] and it has long been accepted worldwide as a good E<sub>T0</sub> estimator when compared with other methods [12–18]. The PM equation presents certain advantages when compared with other E<sub>T0</sub> estimation methods. It can be used globally without the need for additional parameter estimations and it is well documented, has been implemented, and has been extensively validated. The major constraint of using the PM equation is the broad range of required data. Its physically based approach requires measurements of air temperature, windspeed, relative humidity, and solar radiation [4]. The number of stations where all this data is recorded is limited; an additional issue that could be brought into discussion is the data quality. One could also argue about the quality of the observed parameters.

The limitation of available and reliable data motivated Hargreaves and Samani [19] to develop a simpler method—known as the Hargreaves–Samani method (HS)—where only data of maximum and minimum air temperature, extraterrestrial radiation, and a radiation adjustment coefficient ( $k_{Rs}$ ) are required. Since extraterrestrial radiation is an astronomical data and can easily be estimated for a certain day and location, only observed maximum and minimum air temperatures are required. Similar approaches were also developed such as Hargreaves–Samani [19], modified Hargreaves–Samani [20,21], Schendel [22], Baier–Robertson [23], and Trajkovic [24]. These methods are widely compared with PM by different authors [15,18,25–29]. Having pragmatic ETo estimation methods helps to improve water use efficiency. It can be used on simple weather prediction in order to promote a forward-looking irrigation scheduling.

Tabari and Talaee [26] used three Hargreaves-based equations to estimate ETo in Iran (with Köppen–Geiger climates BWh, BWk, BSk, Dsa, and Dsb) and compared the results with PM ETo estimations, leading to root mean square error (RMSE) varying from 0.49 to 1.60 mm day<sup>-1</sup>. A similar study in Jordan was performed by Mohawesh and Talazi [15] where the accuracy of the original Hargreaves–Samani (HS) and the modified Hargreaves–Samani (MHS) equations, as proposed by Droogers and Allen [21], were compared with PM in a Köppen–Geiger BWh BSh climates. Results showed a RMSE ranging from 0.614 to 1.303 mm day<sup>-1</sup> for HS and varying from 0.557 to 2.033 mm day<sup>-1</sup> for the ModHS methods.

Raziei and Pereira [26] compared both HS and PM for 40 weather stations located across Iran, covering different climatic zones, from humid to hyperarid. For humid zones, ETo estimation RMSE varied from 0.18 to 0.37 mm day<sup>-1</sup>, while for semiarid locations it varied from 0.27 to 0.81 mm day<sup>-1</sup>; for hyperarid zones the RMSE ranged from 0.64 to 0.97 mm day<sup>-1</sup>, showing that the estimation error tends to be higher with the increase of aridity. Paredes et al. [18] also assessed the accuracy of the Hargreaves–Samani method after calibrating the  $k_{Rs}$  factor for Azores, Portugal. ETo estimations using this method led to a RMSE varying for 0.47–0.86 mm day<sup>-1</sup> and a coefficient of determination ranging from 0.57 to 0.79 based on a radiation adjustment coefficient that varied from 0.14 to 0.23 °C<sup>-0.5</sup>.

Two complementary studies were performed by Valipour and Eslamian [27] and by Valipour [29] where, for 31 locations from Iran, the accuracy to estimate ETo of various temperature-based methods were compared with PM, including the original HS, four MHS (including those proposed by [20,21]), Schendel (SCH), and Baier–Robertson (B&R) methods. Results showed that the MHS equations proved to be the most accurate for 27 (out of 31) locations with a coefficient of determination ( $R^2$ ) varying from 0.9762 to 0.9990 proving its effectiveness to estimate ETo when limited data is available.

Akhavan et al. [29] compared nine temperature-based methods with PM to estimate actual evapotranspiration of maize in Karaj, Iran—a Köppen–Geiger climates Csa, using two different crop coefficient approaches—single and dual. Both these methods differ since dual crop coefficient methodology consists on a separate computation of the two components of crop evapotranspiration, plant transpiration, which is represented by the basal crop coefficient, and soil evaporation, represented by the soil evaporation coefficient [4]. This study including the equation proposed by Baier and Robertson [23], Schendel [22], Droogers and Allen [20], Trajkovic [24], and Berti et al. [21]. Results showed that the RMSE varied from 1.97 to 5.80 mm day<sup>-1</sup> and from 0.88 to 8.51 mm day<sup>-1</sup> for single and dual crop coefficients approaches, respectively. For both scenarios a modified Hargreaves–Samani method performed the best while the Schendel method performed the worst.

These studies have shown that temperature-based methods can be as nearly as accurate as PM in estimating ETo, suggesting their use where reliable full dataset is lacking. However, and since each climatic condition contains a wide range of magnitude of each weather parameter (e.g., temperature, relative humidity, wind speed, solar radiation, etc.), the results of previous studies may not be applicable for other climates without further validation of each equation. Additionally, they only provide information when using each equation for the whole year, lacking the evaluation of each equation specifically during

the irrigation season, when estimations of ETo are really required for irrigation scheduling and management.

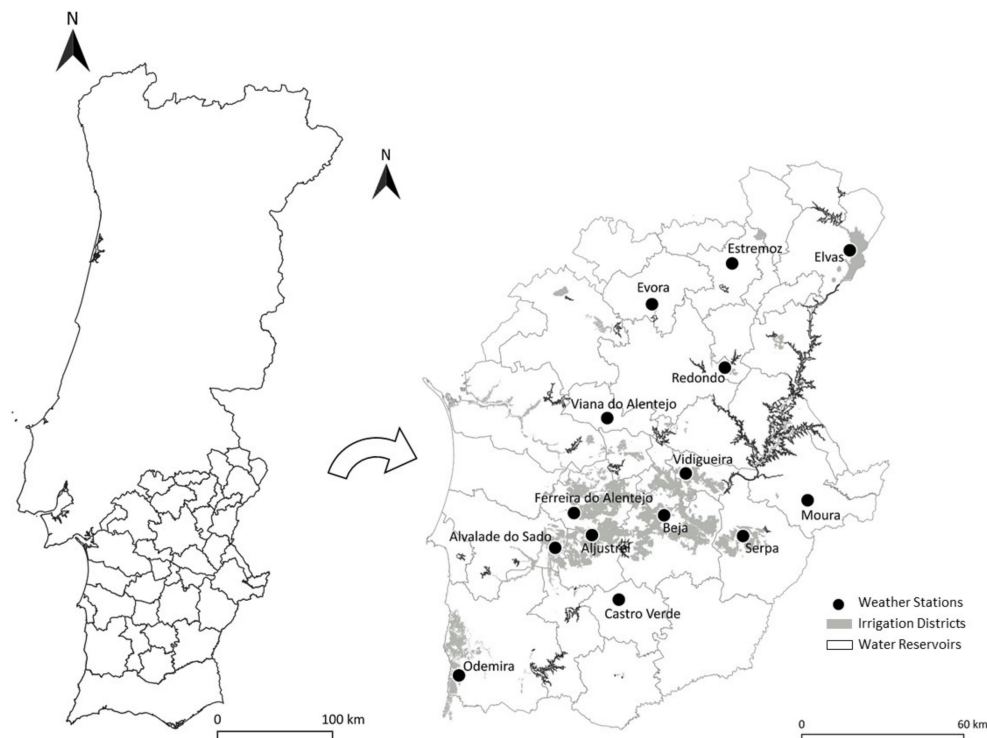
The main objective of this study was to evaluate the performance of nine temperature-based methods to estimate PM reference evapotranspiration for 14 locations of Alentejo Region, Southern Portugal, a hot summer and Mediterranean climate region Csa, during the irrigation season (April to October). The specific objectives are: (1) to evaluate the estimation performance of temperature-based methods by comparison with the PM method during the irrigation season; (2) to calibrate and validate the Hargreaves model for each location to further improve its performance; (3) evaluate the necessity of location-by-location calibration vs. whole region calibration; and (4) to determine the best model based on the weather conditions of each location with the practical purpose of irrigation scheduling.

## 2. Materials and Methods

### 2.1. Study Area

This study was conducted in Alentejo Region, Southern Portugal, and used meteorological data from 14 locations across the region where the Irrigation Operation and Technology Center (COTR) has a network of full weather stations operating and collecting incoming maximum and minimum air temperature, relative humidity, wind speed, rainfall, and solar radiation. All data is daily validated by a team of experienced technicians, assuring its quality and feasibility.

Figure 1 and Table 1 respectively present the geographical position of the weather data locations, and their coordinates and period of observation. The region has a Csa climate according to Köppen–Geiger classification and is characterized by a semiarid Mediterranean climate of hot and dry season in the summer and mild temperature associated to annual rainfall in winter. Irrigation is crucial to achieve farming sustainability and profitability in the Alentejo region [8]. Additionally, and due to recurrent water scarcity, one way to achieve these goals is through the improvement of water use efficiency. Table 2 presents the yearly mean and standard deviation of main weather variables at each assessed location, including for the peak month of July.



**Figure 1.** Weather stations location in the Alentejo region of Portugal with Csa hot summer Mediterranean climate. County boundaries are shown on the left side and specific weather station designation corresponding to location on the right.

**Table 1.** Weather stations coordinates, elevation, distance to the sea, and date ranges of the weather data series.

Weather Station	Code	Latitude (N)	Longitude (W)	Elevation (m)	Distance to the Sea (km)	Date Range	Number of Days <sup>1</sup>
Aljustrel	Alj	37°58'17"	08°11'25"	104	55	Sep/2001–Aug/2019	3828
Alvalade do Sado	Alv	37°55'44"	08°20'45"	79	40	Set/2001–Aug/2019	3837
Beja	Bej	38°02'15"	07°53'06"	206	79	Sep/2001–Aug/2019	3847
Castro Verde	CV	37°45'21"	08°04'35"	200	64	Oct/2007–Aug/2019	2531
Elvas	Elv	38°54'56"	07°05'56"	202	160	Sep/2001–Aug/2019	3840
Estremoz	Est	38°52'20"	07°35'49"	404	120	Feb/2006–Aug/2019	2929
Évora	Evo	38°44'16"	07°56'10"	246	85	Feb/2002–Aug/2019	3699
Ferreira do Alentejo	FdA	38°02'42"	08°15'59"	74	47	Sep/2001–Aug/2019	3843
Moura	Mou	38°05'15"	07°16'39"	172	100	Sep/2001–Aug/2019	3838
Odemira	Ode	37°30'06"	08°45'12"	92	4	Jul/2002–Aug/2019	3681
Redondo	Red	38°31'41"	07°37'40"	236	105	Sep/2001–Aug/2019	3836
Serpa	Ser	37°58'06"	07°33'03"	190	90	May/2004–Aug/2019	3316
Viana do Alentejo	Via	38°21'39"	08°07'32"	138	57	Mar/2006–Aug/2019	2925
Vidigueira	Vid	38°10'37"	07°47'35"	155	86	Nov/2007–Aug/2019	2518

<sup>1</sup> During the irrigation season (April to October).

**Table 2.** Yearly mean and standard deviation of maximum (Tmax) and minimum (Tmin) temperatures, reference evapotranspiration (ETo), and rainfall at each assessed location.

Station	Tmax (°C)	pTmax (°C)	Tmin (°C)	pTmin (°C)	ETo (mm Day <sup>-1</sup> )	pETo (mm Day <sup>-1</sup> )	Rainfall (mm Year <sup>-1</sup> )
Alj	24.4 (±7.5)	33.2 (±4.1)	9.9 (±5.2)	15.1 (±2.2)	3.4 (±2.1)	6.4 (±1.1)	525
Alv	24.7 (±7.3)	33.1 (±4.2)	10.3 (±5.1)	15.4 (±1.9)	3.6 (±2.1)	6.4 (±1.0)	488
Bej	23.9 (±7.8)	33.6 (±4.0)	10.3 (±4.8)	15.2 (±2.5)	3.6 (±2.2)	6.8 (±1.0)	512
CV	24.1 (±7.7)	33.5 (±4.0)	9.8 (±4.9)	14.9 (±2.2)	3.9 (±2.4)	7.3 (±1.2)	393
Elv	24.5 (±8.4)	35.1 (±3.8)	9.4 (±5.7)	15.8 (±2.6)	3.5 (±2.2)	6.8 (±0.9)	504
Est	22.4 (±8.1)	32.3 (±4.1)	9.3 (±5.0)	14.3 (±2.8)	3.0 (±1.9)	5.7 (±0.8)	640
Evo	23.7 (±7.9)	33.1 (±4.1)	8.9 (±5.3)	14.6 (±2.2)	3.3 (±2.0)	6.1 (±1.0)	567
FdA	24.7 (±7.4)	33.4 (±4.1)	9.8 (±5.2)	15.1 (±2.1)	3.3 (±1.9)	6.0 (±1.0)	514
Mou	24.9 (±8.2)	35.5 (±3.8)	8.5 (±6.0)	14.6 (±2.8)	3.2 (±1.9)	6.1 (±0.8)	482
Ode	21.2 (±4.7)	24.8 (±3.3)	11.1 (±3.9)	14.4 (±2.0)	3.0 (±1.4)	4.4 (±0.9)	568
Red	24.1 (±8.1)	34.4 (±3.9)	10.4 (±5.3)	15.9 (±2.5)	3.7 (±2.3)	7.0 (±1.1)	484
Ser	25.2 (±8.1)	35.0 (±3.9)	10.5 (±5.2)	15.9 (±2.5)	3.5 (±2.1)	6.5 (±0.9)	497
Via	23.6 (±7.8)	32.9 (±4.1)	9.9 (±4.7)	14.7 (±2.2)	3.5 (±2.1)	6.4 (±1.1)	625
Vid	24.9 (±7.9)	34.6 (±3.9)	10 (±5.2)	15.6 (±2.2)	3.5 (±2.1)	6.5 (±0.9)	501

p—peak month (July).

## 2.2. Temperature-Based ETo Estimation Methods

All of the data collected were used to estimate the reference evapotranspiration using nine temperature-based models and were compared with PM equation to evaluate the accuracy of each method (Table 3). Those methods were selected based on its low data and demonstrated capacity to estimated ETo in other climates.

ETo is the reference crop evapotranspiration (mm day<sup>-1</sup>), Rn is the net radiation (MJ m<sup>-2</sup> day<sup>-1</sup>), G is the soil heat flux (MJ m<sup>-2</sup> day<sup>-1</sup>),  $\gamma$  is the psychrometric constant (kPa °C<sup>-1</sup>),  $e_s$  is the saturation vapor pressure (kPa),  $e_a$  is the actual vapor pressure (kPa),  $\Delta$  is the slope of the saturation vapor pressure-temperature curve (kPa °C<sup>-1</sup>),  $u_2$  is the mean daily wind speed at 2 m (m s<sup>-1</sup>), H is the elevation (m),  $\phi$  is the latitude (rad), Tmax is the maximum air temperature (°C), Tmin is the minimum air temperature (°C), Tavg is the average air temperature (°C), RH is the average relative humidity (%), Ra is the extra-terrestrial radiation (MJ m<sup>-2</sup> day<sup>-1</sup>), Rs is the solar radiation (MJ m<sup>-2</sup> day<sup>-1</sup>), P is the monthly precipitation (mm), and  $k_{RS}$ , n, and k are experimental coefficients.

**Table 3.** Method used and the parameters applied in each equation.

Method	Code	Reference	Equation	Parameters
FAO Penman-Monteith	PM	[4]	$ET_o = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{T + 273} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34 u_2)}$	H, $\phi$ , Tav <sub>g</sub> , Tmax, Tmin, RH, u, n
Hargreaves-Samani	HS	[19]	$ET_o = 0.0135 \times k_{RS} \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$	Tmax, Tmin, k <sub>RS</sub> , $\phi$
Modified Hargreaves-Samani 1	MHS1	[20]	$ET_o = 0.0030 \times 0.408Ra \times (T_{avg} + 20) \times (T_{max} - T_{min})^{0.4}$	Tmax, Tmin, $\phi$
Modified Hargreaves-Samani 2	MHS2	[20]	$ET_o = 0.0025 \times 0.408Ra \times (T_{avg} + 16.8) \times (T_{max} - T_{min})^{0.5}$	Tmax, Tmin, $\phi$
Modified Hargreaves-Samani 3	MHS3	[20]	$ET_o = 0.0013 \times 0.408Ra \times (T_{avg} + 17.0) \times (T_{max} - T_{min} - 0.0123P)^{0.76}$	Tmax, Tmin, P, $\phi$
Modified Hargreaves-Samani 4	MHS4	[21]	$ET_o = 0.00193 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.517}$	Tmax, Tmin, $\phi$
Schendel	SCH	[21]	$ET_o = 16 \frac{T_{avg}}{RH}$	Tmax, Tmin, RH
Baier and Robertson	B&R	[23]	$ET_o = 0.157T_{max} + 0.158(T_{max} - T_{min}) + 0.109Ra - 5.39$	Tmax, Tmin, $\phi$
Trajkovic	TR	[24]	$ET_o = 0.0023 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.424}$	Tmax, Tmin, $\phi$
Enku and Melesse	E&M	[30]	$ET_o = \frac{(T_{max})^n}{k}$	Tmax, n, k

For the Hargreaves–Samani equation, the empirical coefficient  $k_{RS}$  was originally considered to range from 0.16 to 0.19 °C<sup>−0.5</sup>, respectively for “interior” or “coastal” regions [4]. In the original version of the Hargreaves–Samani (HS) equation [19], a bulk constant term of 0.023, known as the Hargreaves coefficient, is used. It corresponds to the product  $0.0135 \times k_{RS}$ , with  $k_{RS} = 0.17$  °C<sup>−0.5</sup>, and 0.0135 representing a conversion of the units’ constant. However,  $k_{RS}$  is supposed to vary with altitude, reflecting the changes of air pressure and volumetric heat capacity of the atmosphere [31]. Therefore,  $k_{RS}$  should vary spatially, internalizing the effects of the site elevation and distance to sea [4]. For a similar climatic region, Moratiel et al. [32] evaluated the effectiveness of  $k_{RS}$  calibration, both regionally and seasonally, and concluded that, when calibrating this factor,  $ET_o$  estimations may improve, leading to a decrease of the RMSE from 0.62 to 0.54 mm day<sup>−1</sup>. Thus, one should argue that a constant  $k_{RS} = 0.17$  or the established range of 0.16–0.19 °C<sup>−0.5</sup> for “interior” or “coastal” regions may not be suitable for all locations, requiring its calibration to reflect the site-specific conditions; however, when insufficient data is available, the original value may be used.

As for this study, two approaches were used for the Hargreaves–Samani equation. One where a standard  $k_{RS}$  of 0.17 °C<sup>−0.5</sup> was used for all locations as proposed by [19]. Another, where  $k_{RS}$  was adjusted for each location. Adopting a trial and error procedure, being calibrated using 50% of the years, randomly chosen, from the dataset, and validated for the independent dataset obtained for the remainder of the years.

Along with the Hargreaves–Samani equation, five HS modified methods were also used—modified Hargreaves 1 (ModHS1), modified Hargreaves 2 (ModHS2), modified Hargreaves 3 (ModHS3), modified Hargreaves 4 (ModHS4), and Trajkovic (TR). These models vary from the original method since some parameters were modified, namely the  $k_{RS}$ , the empirical temperature coefficient (HT) and the empirical Hargreaves exponent (HE). As stated, for the original HS,  $k_{RS} = 0.17$  °C<sup>−0.5</sup>, HT = 17.8, and HE = 0.5. For ModHS1, the bulk constant term of 0.023 was altered, resulting from the adoption of a  $k_{RS} = 0.22$  °C<sup>−0.5</sup>. Additionally, both HT and HE were adjusted, resulting on HT = 20 and HE = 0.4. Both ModHS2 and ModHS4 were also adjusted, with the former having a HT = 16.8 and the latter a HE = 0.517; as for radiation adjustment coefficient, a  $k_{RS} = 0.19$  and 0.14 °C<sup>−0.5</sup> was used, respectively. For ModHS3 both HT and HE were also modified to 17 and 0.76, respectively, and a  $k_{RS} = 0.10$  °C<sup>−0.5</sup> was used. This method includes rainfall as an added parameter to estimate  $ET_o$  since, according to [20], the precipitation can in

some regards represent relative levels of humidity, improving the method accuracy. As for TR, HE was adjusted, equaling 0.424.

Three other equations were also evaluated — Schendel (SCH), Baier and Robertson (B&R), and Enku and Melesse (E&M). B&R requires the same data as the original HS but it consists in a different approach (vd. Table 2). As for SCH an additional parameter is used—average relative humidity (RH)—and Ra is not considered to estimate ETo. E&M presents itself as the less data demanding method since it only requires the maximum and average temperature; the latter was used to locally adjust the experiment coefficient k.

### 2.3. Evaluation Criteria

The accuracy of each method was assessed by comparing their results with those of the PM equation through the indicators listed below:

- (1) The coefficients of regression and determination relating the PM and temperate-based ETo, b and  $R^2$  respectively, are defined as:

$$b = \frac{\sum_{i=1}^n ET_{PMi} ET_{TBi}}{\sum_{i=1}^n ET_{PMi}^2} \quad (1)$$

$$R^2 = \left\{ \frac{\sum_{i=1}^n (ET_{PMi} - \overline{ET_{PM}}) (ET_{TBi} - \overline{ET_{TB}})}{\left[ \sum_{i=1}^n (ET_{PMi} - \overline{ET_{PM}})^2 \right]^{0.5} \left[ \sum_{i=1}^n (ET_{TBi} - \overline{ET_{TB}})^2 \right]^{0.5}} \right\}^2 \quad (2)$$

- (2) The root mean square error, RMSE, which characterizes the variance of the estimation error:

$$RMSE = \left[ \frac{\sum_{i=1}^n (ET_{TBi} - ET_{PMi})^2}{n} \right]^{0.5} \quad (3)$$

where  $ET_{PMi}$  and  $ET_{TBi}$  ( $i = 1, 2, \dots, n$ ) represent pairs of values of ETo estimated using PM equation and other temperature-based method, respectively, for a given variable and  $\overline{ET_{PM}}$  and  $\overline{ET_{TB}}$  are the respective mean values and n is the number of days used in the assessment.

## 3. Results

### 3.1. Calibration and Validation of Radiation Factor ( $k_{RS}$ )

As previously stated, estimations of ETo may be performed using the Hargreaves–Samani equation with a locally calibrated and validated radiation adjustment coefficient  $k_{RS}$ , thus resulting in an adjusted HS equation. Table 4 shows the calibrated and validated  $k_{RS}$  values for each one of the 14 locations and for the entire Alentejo region. Table 4 also presents the statistical summary (b,  $R^2$ , and RMSE) of ETo estimates for each location, that resulted from adopting the Adjusted HS equation, for calibration, validation and all years, and the ETo estimations that resulting from using the original HS equation.

Results (Table 4) show that the estimation accuracy of ETo when using the Hargreaves–Samani equation, after the calibration of  $k_{RS}$  for each site, improved for 10 locations. When using the standard  $k_{RS}$ , b varied from 0.96 to 1.23 (with only CV resulting in a b < 1.00),  $R^2$  values range from 0.69 to 0.83, and RMSE values varying from 0.65 to 1.38 mm day<sup>-1</sup>. When adopting the calibrated  $k_{RS}$  for each location, these indicators improved significantly, with b ranging from 0.96 to 1.01 and a RMSE varying from 0.64 to 0.88 mm day<sup>-1</sup>, while the correlation ( $R^2$ ) remains the same as expected. Some locations stand out from the set: the RMSE for Moura decreased from 1.38 to 0.71 mm day<sup>-1</sup> (with b decreasing from 1.23 to 1.01), while for Estremoz it decreased from 1.10 to 0.70 mm day<sup>-1</sup> (with b decreasing from 1.16 to 0.96), representing a decrease of 19 and 36%, respectively. As for the Alentejo region, the calibration of  $k_{RS}$  led to acceptable results with a slope of 1.0 and a RMSE equal to 0.86 mm day<sup>-1</sup>, both in line with the accuracy found for each individual location.

**Table 4.** Accuracy of daily reference evapotranspiration estimations using the Hargreaves equation after  $k_{RS}$  factor calibration/validation.

Station	Adjusted HS Equation									Original HS Equation ( $k_{RS} = 0.17 \text{ } ^\circ\text{C}^{-0.5}$ )			
	Adjusted $k_{RS}$ ( $^\circ\text{C}^{-0.5}$ )	Calibration			Validation			All			b	R <sup>2</sup>	RMSE
		b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE			
Alj	0.16	1.02	0.80	0.78	1.01	0.80	0.79	1.01	0.80	0.78	1.08	0.80	0.92
Alv	0.16	1.00	0.80	0.74	0.99	0.84	0.68	0.99	0.82	0.71	1.05	0.82	0.81
Bej	0.17	1.00	0.81	0.81	1.02	0.85	0.73	1.01	0.83	0.77	1.01	0.83	0.77
CV	0.17	0.96	0.83	0.84	0.96	0.84	0.79	0.96	0.83	0.82	0.96	0.83	0.82
Elv	0.16	1.02	0.79	0.85	1.01	0.78	0.87	1.01	0.78	0.86	1.08	0.78	0.99
Est	0.14	0.97	0.80	0.65	0.95	0.78	0.73	0.96	0.79	0.70	1.16	0.79	1.10
Evo	0.15	0.99	0.77	0.79	0.98	0.75	0.82	0.99	0.76	0.80	1.12	0.76	1.07
FdA	0.15	1.01	0.79	0.74	1.01	0.81	0.69	1.01	0.80	0.72	1.14	0.80	1.06
Mou	0.14	1.03	0.80	0.72	1.00	0.81	0.70	1.01	0.81	0.71	1.23	0.81	1.38
Ode	0.17	1.03	0.65	0.69	1.00	0.74	0.59	1.01	0.69	0.64	1.01	0.69	0.65
Red	0.17	1.00	0.80	0.86	1.00	0.78	0.90	1.00	0.79	0.88	1.00	0.79	0.88
Ser	0.15	1.00	0.78	0.79	0.96	0.79	0.79	0.98	0.78	0.79	1.11	0.78	1.06
Via	0.16	1.00	0.81	0.74	0.98	0.80	0.79	0.99	0.81	0.77	1.05	0.81	0.86
Vid	0.15	0.98	0.81	0.73	0.97	0.80	0.77	0.97	0.80	0.75	1.10	0.80	1.00
ALENTEJO	0.16	1.02	0.76	0.86	1.01	0.78	0.83	1.01	0.77	0.84	1.08	0.77	0.98

Figures 2 and 3 try to show the effect of site-specific conditions over  $k_{RS}$ . As discussed previously, and according to Allen [31] and Allen et al. [4], the  $k_{RS}$  is supposed to reflect the volumetric heat capacity of the atmosphere, leading to lower  $k_{RS}$  values with higher elevation and higher distance to the sea. Analyzing Figure 2 it can be concluded that “coastal” locations tended to show a higher  $k_{RS}$ , while the “interior” site led to lower values; exception made for Beja ( $k_{RS} = 0.17 \text{ } ^\circ\text{C}^{-0.5}$ ), Castro Verde ( $k_{RS} = 0.17 \text{ } ^\circ\text{C}^{-0.5}$ ), Redondo ( $k_{RS} = 0.17 \text{ } ^\circ\text{C}^{-0.5}$ ), and Elvas ( $k_{RS} = 0.16 \text{ } ^\circ\text{C}^{-0.5}$ ) where lower values would be expected. One can assume that due to the presence of irrigation districts and water reservoirs close to each site the influence of air moisture on the volumetric heat capacity of the atmosphere can affect the radiation constant.

### 3.2. Estimating ETo by Temperature-Based Methods

The nine temperature-based methods were used to estimate daily ETo ( $\text{mm day}^{-1}$ ) for the 14 locations and were conducted from April to October, the period that covers the growing season of the main irrigated crops in Alentejo. The ETo values estimated by the all temperature-based equations were compared with estimates by the standard FAO-56 PM equation. The statistical summary (b, R<sup>2</sup>, and RMSE) of ETo estimates for the 14 locations are presented in Table 5.

The original Hargreaves–Samani (HS) equation tended to overestimate ETo for all 14 locations (b varying from 1.00 to 1.23), except for Castro Verde where ETo was underestimated (b = 0.96). R<sup>2</sup> values ranged from 0.69 to 0.83, with Odemira performing the worse, and Beja and Castro Verde achieving the best correlation. RMSE values for all fourteen locations varied from 0.65 to 1.38  $\text{mm day}^{-1}$ , with nine stations not exceeding a RMSE higher than 1.00  $\text{mm day}^{-1}$ . When calibrating  $k_{RS}$  for each location (vd. Table 4), this accuracy can be improved, one can conclude that the recurrent overestimation and related estimation error may be explained by using a standard  $k_{RS} = 0.17 \text{ } ^\circ\text{C}^{-0.5}$ .

When using the MHS1, MHS2, and MHS3 equations, ETo tends to be overestimated for all locations (b varying from 1.02 to 1.49); only with MHS1 for Castro Verde ETo is underestimated (b = 0.99). For the same equations, R<sup>2</sup> ranged from 0.65 to 0.84, with the combination Odemira/MHS3 performing worse and Castro Verde/MHS2 leading to the best correlation. RMSE varied from 0.76 to 2.58  $\text{mm day}^{-1}$ . The MHS2 at Évora led to the lowest RMSE; contrarily, when using the MHS3 to estimate ETo at Moura it led to a RMSE of 2.58  $\text{mm day}^{-1}$ . These results show that if a higher a  $k_{RS}$  (= 0.22 and  $19 \text{ } ^\circ\text{C}^{-0.5}$  for MHS1

and MHS2, respectively) than the ones calibrated for each location tend to overestimate ETo. Additionally, when including rainfall to estimate ETo (as in MHS3) does not improve ETo estimation but quite the opposite. MH3 led to the lowest performance from all the four modified Hargreaves–Samani equation, suggesting that during the irrigation season, when rainfall is scarce, the inclusion of this parameter is unadvisable.

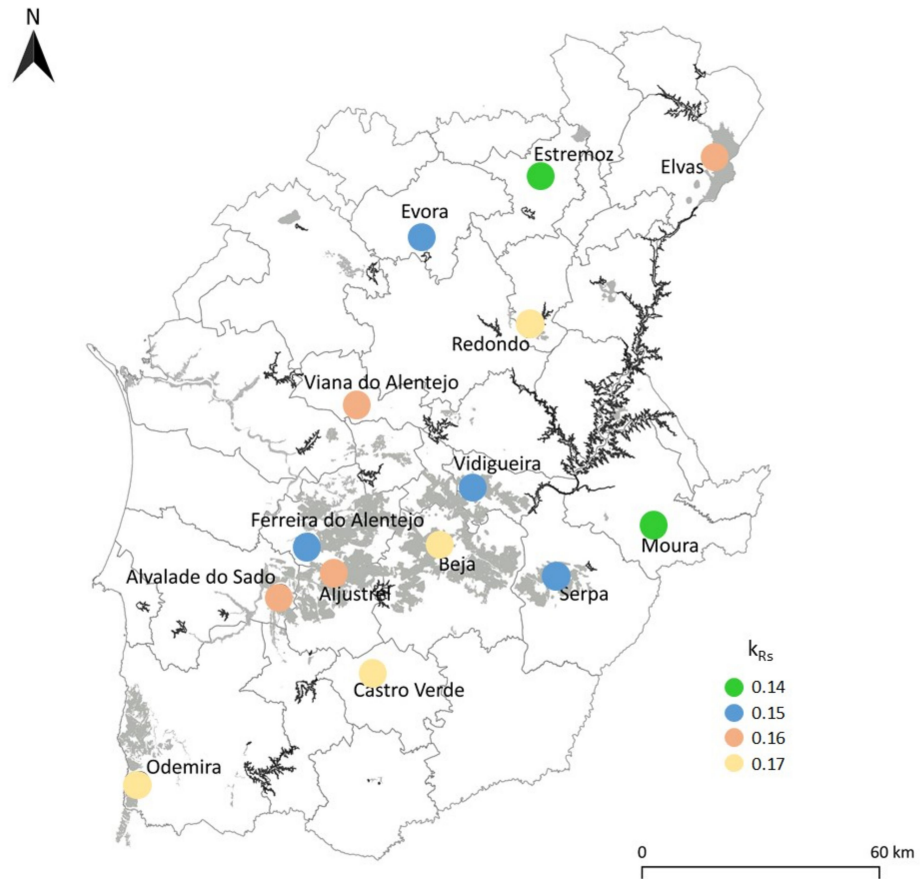


Figure 2. Calibrated and validated radiation adjustment coefficient  $k_{Rs}$  for each location.

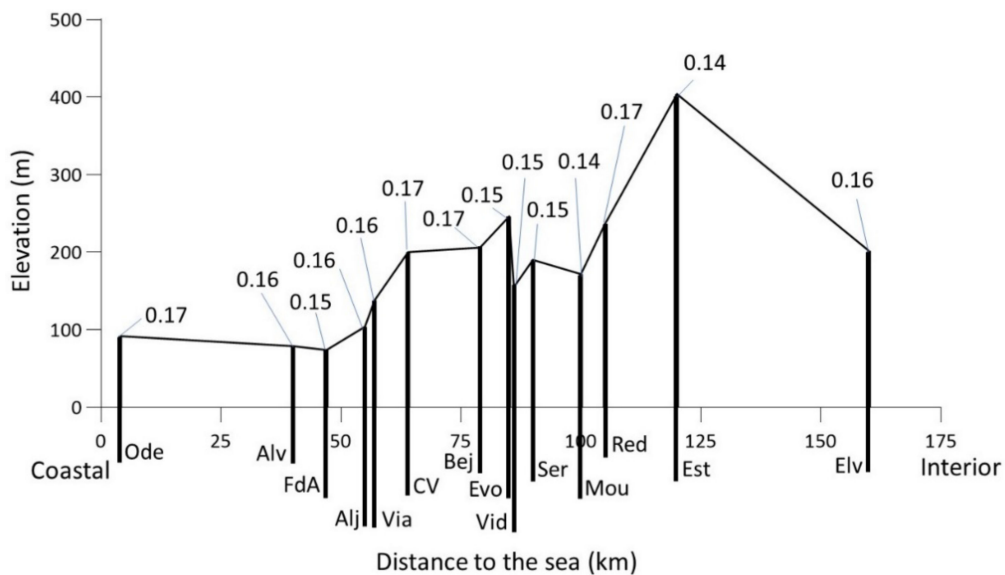


Figure 3. Calibrated and validated radiation adjustment coefficient  $k_{Rs}$  for each longitude and land elevation.



**Table 5.** Slope (b), coefficient of determination ( $R^2$ ), and root mean square error (RMSE) for the relationship between daily ETo estimated by the Penman–Monteith equation and by nine temperature-based equations.

Equation	Station	Aljustrel			Alvalade do Sado			Beja			Castro Verde			Elvas			Estremoz			Évora			Ferreira do Alentejo		
		b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE
Original Hargreaves		1.08	0.80	0.92	1.05	0.82	0.81	1.01	0.83	0.77	0.96	0.83	0.82	1.08	0.78	0.99	1.16	0.79	1.10	1.12	0.76	1.07	1.14	0.80	1.06
Adjusted Hargreaves		1.01	0.80	0.78	0.99	0.82	0.71	1.01	0.83	0.77	0.96	0.83	0.82	1.01	0.78	0.86	0.96	0.79	0.70	0.99	0.76	0.80	1.01	0.80	0.72
Modified Hargreaves 1		1.12	0.80	1.04	1.09	0.82	0.91	1.05	0.83	0.84	0.99	0.82	0.85	1.11	0.79	1.09	1.21	0.80	1.26	1.15	0.77	1.19	1.18	0.80	1.21
Modified Hargreaves 2		1.15	0.80	1.14	1.12	0.82	1.01	1.08	0.83	0.90	1.02	0.84	0.83	1.15	0.78	1.23	1.24	0.79	1.36	1.19	0.76	1.32	1.21	0.80	1.34
Modified Hargreaves 3		1.26	0.77	1.73	1.23	0.79	1.59	1.18	0.82	1.38	1.13	0.83	1.20	1.28	0.75	1.89	1.35	0.76	1.91	1.32	0.73	1.98	1.35	0.77	1.99
Modified Hargreaves 4		0.95	0.80	0.77	0.93	0.82	0.76	0.89	0.83	0.87	0.85	0.84	1.11	0.95	0.78	0.84	1.03	0.79	0.74	0.99	0.76	0.81	1.01	0.80	0.72
Schendel		1.14	0.54	1.64	1.09	0.58	1.38	1.14	0.51	1.85	1.04	0.61	1.55	1.27	0.50	2.41	1.28	0.37	2.37	1.17	0.49	1.82	1.19	0.55	1.69
Baier and Robertson		1.16	0.75	1.28	1.13	0.76	1.17	1.08	0.81	0.96	1.03	0.83	0.86	1.16	0.73	1.39	1.24	0.75	1.44	1.21	0.72	1.47	1.23	0.74	1.51
Trajkovic		0.87	0.80	0.94	0.85	0.82	0.97	0.82	0.83	1.15	0.77	0.83	1.45	0.87	0.79	1.00	0.94	0.79	0.70	0.90	0.77	0.87	0.92	0.80	0.75
Enku and Melesse		1.47	0.53	3.12	1.43	0.53	3.00	1.42	0.59	2.99	1.33	0.64	2.61	1.52	0.52	3.54	1.59	0.50	3.46	1.53	0.52	3.34	1.56	0.50	3.39

Equation	Station	Moura			Odemira			Redondo			Serpa			Viana			Vidigueira			ALENTEJO		
		b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE	b	R <sup>2</sup>	RMSE
Original Hargreaves		1.23	0.81	1.38	1.01	0.69	0.65	1.00	0.79	0.88	1.11	0.78	1.06	1.05	0.81	0.86	1.10	0.80	1.00	1.08	0.77	0.98
Adjusted Hargreaves		1.01	0.81	0.71	1.01	0.69	0.64	1.00	0.79	0.88	0.98	0.78	0.79	0.99	0.81	0.77	0.97	0.80	0.75	1.01	0.77	0.84
Modified Hargreaves 1		1.26	0.80	1.51	1.10	0.69	0.79	1.03	0.79	0.94	1.15	0.79	1.18	1.09	0.80	0.99	1.14	0.81	1.12	1.12	0.77	1.08
Modified Hargreaves 2		1.31	0.81	1.71	1.08	0.68	0.76	1.21	0.80	1.34	1.18	0.78	1.33	1.12	0.81	1.05	1.17	0.80	1.26	1.14	0.77	1.19
Modified Hargreaves 3		1.49	0.78	2.58	1.06	0.65	0.88	1.17	0.77	1.45	1.31	0.75	1.98	1.23	0.81	1.55	1.31	0.78	1.91	1.26	0.74	1.78
Modified Hargreaves 4		1.09	0.80	0.87	0.89	0.68	0.73	0.88	0.79	1.01	0.98	0.78	0.80	0.93	0.81	0.80	0.98	0.80	0.76	0.95	0.77	0.84
Schendel		1.39	0.53	2.60	1.04	0.35	1.12	1.19	0.55	2.04	1.27	0.47	2.34	1.14	0.57	1.70	1.26	0.51	2.24	1.22	0.47	2.24
Baier and Robertson		1.35	0.76	1.94	1.02	0.66	0.77	1.07	0.77	1.05	1.19	0.73	1.46	1.12	0.79	1.11	1.19	0.75	1.40	1.15	0.73	1.31
Trajkovic		0.98	0.81	0.68	0.85	0.69	0.80	0.81	0.79	1.27	0.89	0.79	0.89	0.85	0.80	1.02	0.89	0.81	0.88	0.87	0.77	0.97
Enku and Melesse		1.74	0.56	4.00	1.34	0.31	2.22	1.36	0.59	2.83	1.49	0.49	3.35	1.46	0.64	2.95	1.50	0.52	3.34	1.53	0.53	3.62

From all the modified Hargreaves equations, MHS4 proved to be the one that led to better results. *b* varied from 0.85 to 1.09 (for Castro Verde and Moura, respectively), and with a *R*<sup>2</sup> ranging from 0.68 to 0.84. Additionally, and comparing to the other modified HG equations, MHS4 led to smaller RMSE varying from 0.72 to 1.11 mm day<sup>-1</sup>, for Ferreira do Alentejo and Castro Verde, respectively. These results can be explained by the lower than standard *k*<sub>RS</sub> value (0.14 °C<sup>-0.5</sup>) and a slightly higher HE (0.517) that are adopted by this method, leading to a performance close to the one achieved by the adjusted HS equation.

Differently to the other equations, TR underestimated ETo for all 14 locations, with *b* varying from 0.77 to 0.98. Similarly, TR showed *R*<sup>2</sup> values ranging from 0.69 to 0.83, with Odemira and Beja/Castro Verde having the lowest and highest correlations, respectively. As for HG, TR led to similar RMSE ranging from 0.68 to 1.45 mm day<sup>-1</sup>, with only four locations having a RMSE greater than 1.00 mm day<sup>-1</sup>. Despite adopting the standard *k*<sub>RS</sub> and HT values (0.17–0.5 and 17.8 °C, respectively), TR uses a lower HE (0.424) than the original HS equation. One can assume that the underestimation of ETo by TR may be explained by the empirical values used by this method.

Despite of having similar *R*<sup>2</sup> as for TR (varying from 0.66 to 0.83), the B&R equation overestimates ETo for all 14 locations (*b* ranging from 1.02 to 1.35). Additionally, B&R led to higher RMSE than TR, varying from 0.77 to 1.94 mm day<sup>-1</sup>, with only three locations (Beja, Castro Verde and Odemira) not exceeding a RMSE higher than 1.00 mm day<sup>-1</sup>, disapproving its adoption for the region. Similarly, SCH and E&M proved to be ineffective to estimate ETo. Both equations overestimated ETo (SCH—1.04 ≤ *b* ≤ 1.39; E&M—1.33 ≤ *b* ≤ 1.74) and led to *R*<sup>2</sup> lower than 0.64. It resulted that the RMSE for SCH ranged from 1.12 to 2.60 mm day<sup>-1</sup>, and for E&M it varied from 2.22 to 4 mm day<sup>-1</sup>. It can be concluded that these methods are not suitable to estimate ETo for any of the locations under study.

When upscaling the approach for the whole Alentejo region, results show that the adjusted HS method performed the best, with a slight overestimation (*b* = 1.01) and a RMSE equal to 0.84 mm day<sup>-1</sup>. If no *k*<sub>RS</sub> calibration was performed MH4 would be the most accurate model; despite leading to the same RMSE as for the adjusted HS method, this equation tends to underestimate ETo (*b* = 0.95). The least accurate method for the whole region is the E&M equation, with a RMSE = 3.62 mm day<sup>-1</sup>.

Table 6 and Figure 4 present the recommended methods to be used for each location. For all 14 locations the selected equation led to RMSE lower than 0.88 mm day<sup>-1</sup> (12 locations show a RMSE ≤ 0.80 mm day<sup>-1</sup>), and an estimation error smaller than 5% (*b* varied from 0.95 to 1.02). Table 6 also compiles the recommended equations for each location to be adopted when estimating ETo when only temperature data is available. It can be concluded that the adjusted HS method, when weather data is missing, is the most adequate approach to estimate ETo during the irrigation season.

**Table 6.** Accuracy of ETo estimations using temperature-based methods (*k*<sub>RS</sub> calibrated).

Station	Most Adequate Method	<i>b</i>	<i>R</i> <sup>2</sup>	RMSE	Recommended Equation
Alj	MHS4	0.95	0.80	0.77	$E_{To} = 0.00193 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.517}$
Alv	Adjusted HS	0.99	0.82	0.71	$E_{To} = 0.00216 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
Bej	Original/ Adjusted HS	1.01	0.83	0.77	$E_{To} = 0.00230 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
CV	MHS2	1.02	0.84	0.83	$E_{To} = 0.00250 \times 0.408Ra \times (T_{avg} + 16.8) \times (T_{max} - T_{min})^{0.5}$
Elv	MHS4	0.95	0.78	0.84	$E_{To} = 0.00193 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.517}$
Est	Adjusted HS	0.96	0.79	0.70	$E_{To} = 0.00189 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
Evo	Adjusted HS	0.99	0.76	0.80	$E_{To} = 0.00203 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
FdA	MHS4	1.01	0.80	0.72	$E_{To} = 0.00193 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.517}$
Mou	TR	0.98	0.81	0.68	$E_{To} = 0.00230 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.424}$
Ode	Original/ Adjusted HS	1.01	0.69	0.64	$E_{To} = 0.00230 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
Red	Original/ Adjusted HS	1.00	0.79	0.88	$E_{To} = 0.00230 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
Ser	Adjusted HS	0.98	0.78	0.79	$E_{To} = 0.00203 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
Via	Adjusted HS	0.99	0.81	0.77	$E_{To} = 0.00216 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.5}$
Vid	MHS4	0.98	0.80	0.76	$E_{To} = 0.00193 \times 0.408Ra \times (T_{avg} + 17.8) \times (T_{max} - T_{min})^{0.517}$

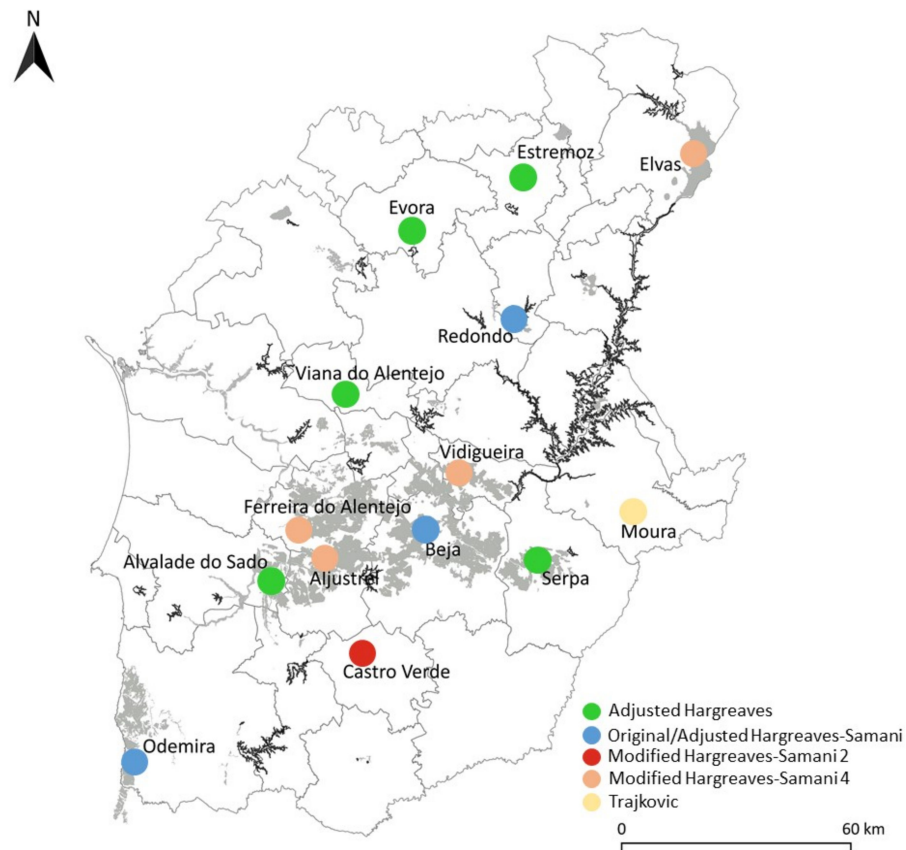


Figure 4. Most suitable method for each location.

$ETo$  is the reference crop evapotranspiration ( $\text{mm day}^{-1}$ ),  $T_{\text{max}}$  is the maximum air temperature ( $^{\circ}\text{C}$ ),  $T_{\text{min}}$  is the minimum air temperature ( $^{\circ}\text{C}$ ),  $T_{\text{avg}}$  is the average air temperature ( $^{\circ}\text{C}$ ), and  $R_a$  is the extraterrestrial radiation ( $\text{MJ m}^{-2} \text{day}^{-1}$ ).

#### 4. Discussion

Results show that, after the calibration of radiation adjustment coefficient ( $k_{RS}$ ), the estimation of  $ETo$  tends to be more accurate than when adopting the standardized  $k_{RS}$  ( $= 0.17 \text{ }^{\circ}\text{C}^{-0.5}$ ) value as proposed by [19]. The accuracy indicators are in accordance with the ones found on previous studies [18,26] where, after  $k_{RS}$  calibration,  $ETo$  led to similar results.

Figure 3 seems to show that, and contrarily to what was proposed by Allen [30] and Allen et al. [5], there is no close relation with elevation and distance to the sea and lower  $k_{RS}$ ; since locations with higher distance to the sea and higher elevations do not reflect this assumption. In fact, and comparing the calibrated  $k_{RS}$  for Moura and Redondo, and for Estremoz and Elvas, where one would guess higher and lower  $k_{RS}$ , a lower (higher) radiation adjustment coefficient can be found. Additionally, and differently from what is proposed by Allen et al. [4], results show that “coastal” locations present a  $k_{RS}$  close to  $0.17 \text{ }^{\circ}\text{C}^{-0.5}$  (lower than the proposed standard value of  $0.19 \text{ }^{\circ}\text{C}^{-0.5}$ ), with “interior” locations showing a  $k_{RS}$  that varies from  $0.14$  to  $0.16 \text{ }^{\circ}\text{C}^{-0.5}$  (slightly lower than standardized value of  $0.16 \text{ }^{\circ}\text{C}^{-0.5}$ ).

For five locations (Alvalade do Sado, Estremoz, Évora, Serpa, and Viana do Alentejo) the adjusted Hargreaves–Samani equation proved to be the most efficient, while the original HS is the most adequate for three locations (Beja, Odemira, and Redondo); this is due to when calibrating the  $k_{RS}$  factor, the standard value of  $0.17 \text{ }^{\circ}\text{C}^{-0.5}$  led to the best results. For Aljustrel, Elvas, Moura, and Vidigueira, MHS4 proved to be the most suitable to estimate  $ETo$ . For these locations, the combined effect of lowering the  $k_{RS}$  value from  $0.17$

to  $0.14\text{ }^{\circ}\text{C}^{-0.5}$  and increasing the HE value from 0.5 to 0.517 led to better results than only adjusting the radiation adjustment coefficient as in for the adjusted HS method. Table 6 also shows that MHS2 proves to be the most suitable to estimate ETo for Castro Verde and TR for Moura. As for MHS4, MHS2 the combinations of increasing the  $k_{RS}$  from 0.17 to  $0.19\text{ }^{\circ}\text{C}^{-0.5}$  and decreasing the HT value from 17.8 to 16.8 led to better results than to adopt the original HS equation for CV, mainly due to lowering the average temperature counterpart when estimating ETo. As for Moura, results show that decreasing the HE from 0.5 to 0.424 proved to be more effective than calibrating the  $k_{RS}$  factor for this location.

## 5. Conclusions

Reference evapotranspiration (ETo) estimations, if accurate, may serve as a decision support indicator for water management, irrigation system design and management, and irrigation scheduling. However, some methods require a wide range of data. Simpler methods, where only data of maximum and minimum air temperature and extraterrestrial radiation is required, prove to be useful where data collection is limited.

Results show that after calibration of the radiation adjustment coefficient,  $k_{RS}$ , an adjusted calibrated Hargreaves–Samani (HS) method can be used to estimate daily ETo acceptably well in most subregions of Alentejo (Alvalade do Sado, Beja, Estremoz, Évora, Odemira, Redondo, Serpa, and Viana do Alentejo).

This study has evaluated the accuracy of nine temperature-based equations to estimate ETo across fourteen different locations in Alentejo, Southern Portugal, in relation to the standard method of Penman–Monteith. From the nine equations adopted in this study, the locally adjusted Hargreaves–Samani method, proved to be most accurate to estimate ETo (for 8 out of 14 locations).

Thus, for the Alentejo region, accuracy results indicated the appropriateness of using the Hargreaves–Samani method for most of its subregions since it leads to acceptable ETo estimations when enough data is not available. Nonetheless, additional studies are recommended to better assess the use of temperature-based methods in Southern Portugal, when limited data is available, since results show that the calibration of the empirical temperature coefficient and the empirical Hargreaves exponent is advisable.

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**Conflicts of Interest:** The authors declare no conflict of interest.

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